

Propriedades Físicas da Água do Mar

- Temperatura
 - Variação - amplitude
 - Métodos de determinação
 - Distribuição horizontal superficial
 - Distribuição vertical
 - Temperatura potencial
- Salinidade
 - Conceito
 - Variação - amplitude
 - Métodos de determinação
 - Distribuição horizontal superficial
 - Distribuição vertical
- Densidade
 - Conceito, anomalias, volume específico
 - Variação – amplitude
 - Métodos de determinação
 - Distribuição horizontal superficial
 - Distribuição vertical
 - Estabilidade/Instabilidade

Presença do sal na água do mar: impactos

- Altera:
 - Viscosidade
 - Compressibilidade – propagação do som
 - Índice de refração da luz
 - Densidade
 - Temperatura de máxima densidade
 - Ponto de congelamento

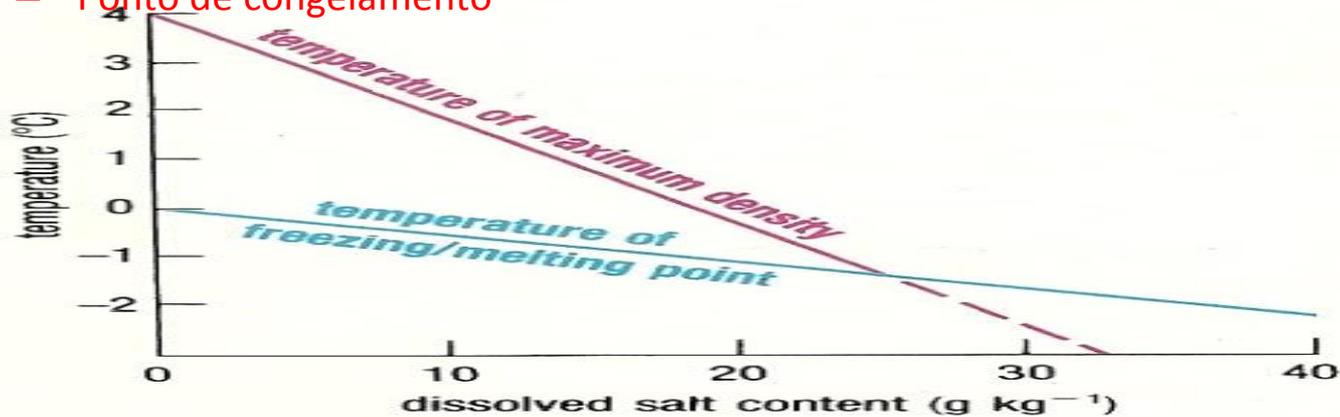


Figure 1.2 Temperatures of freezing and melting point and maximum density of liquid water as functions of dissolved salt content.

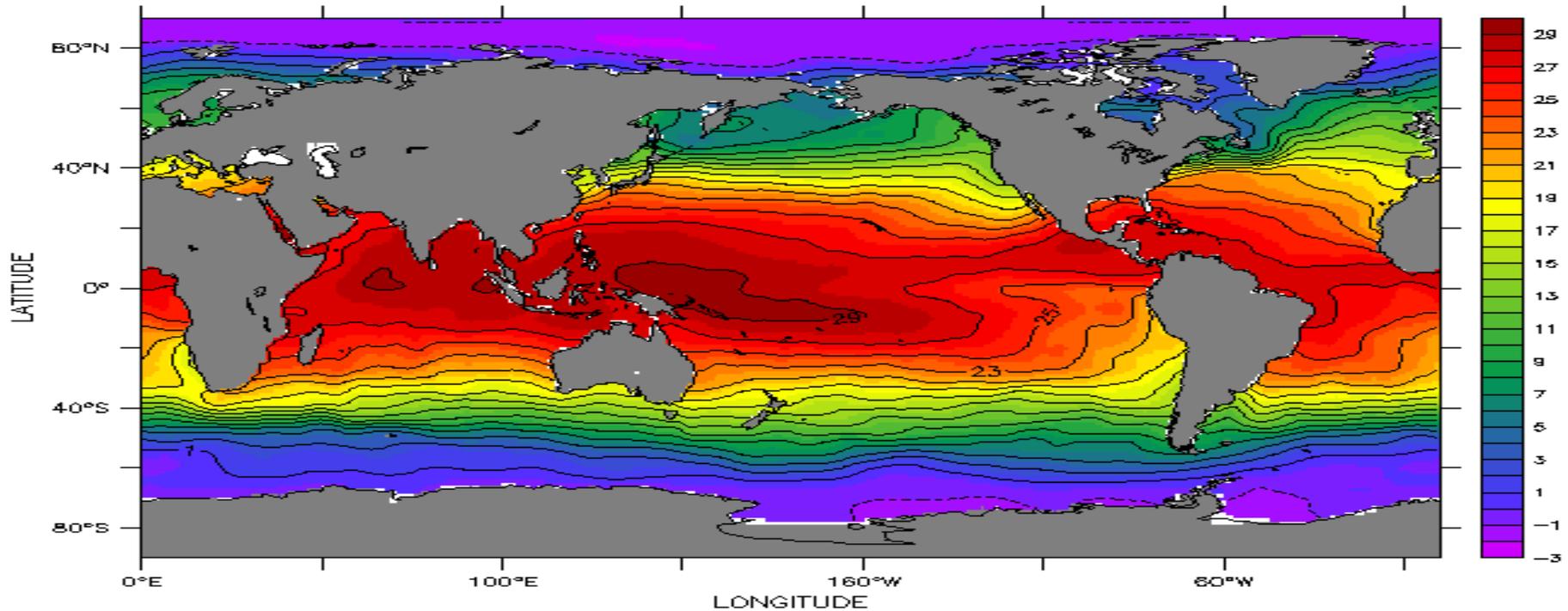
Temperatura da Água do Mar

NOAA/PMEL TMAP  FERRET Ver 4.0

DEPTH (m) : 0

DATA SET: ocean-atlas-annual

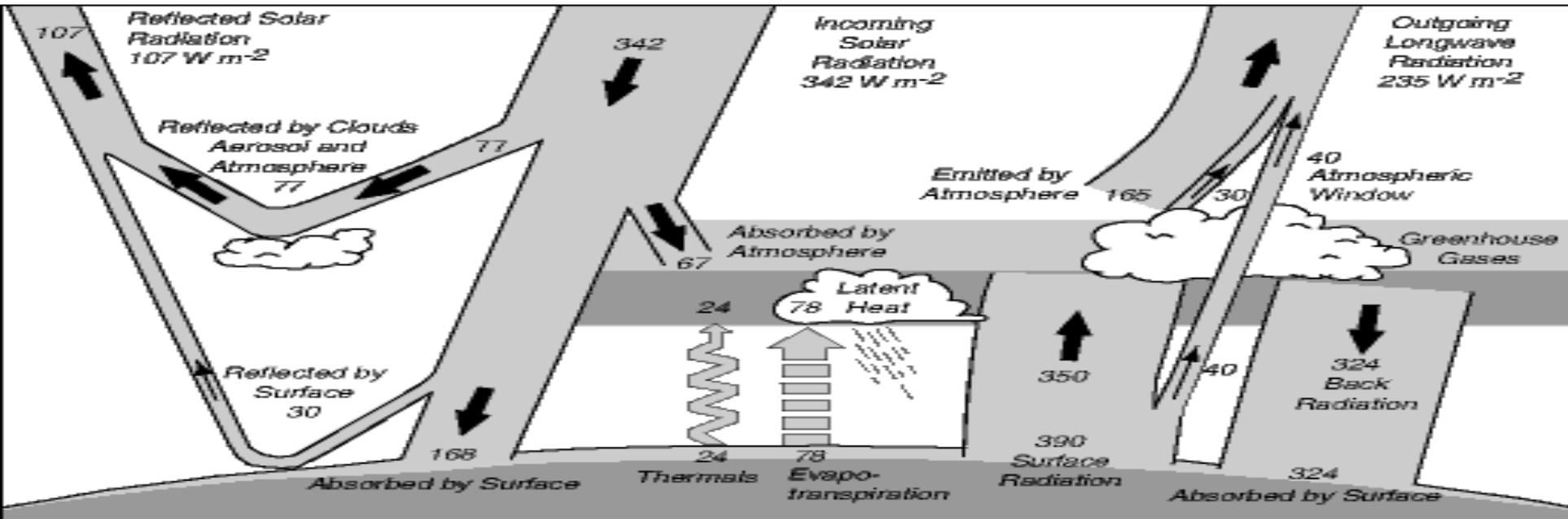
World Ocean Atlas 1994 * 1x1 Degree Annual Means



Temperature (Deg C)

Mean sea-surface temperature calculated from the optimal interpolation technique (Reynolds and Smith, 1995) using shipreports and AVHRR measurements of temperature. Contour interval is 1° C with heavy contours every 5° C. Shaded areas exceed 29° C.

Balanço de Calor



The mean annual radiation and heatbalance of the earth. From Houghton et al. (1996: 58), which used data from Kiehl and Trenberth (1996).

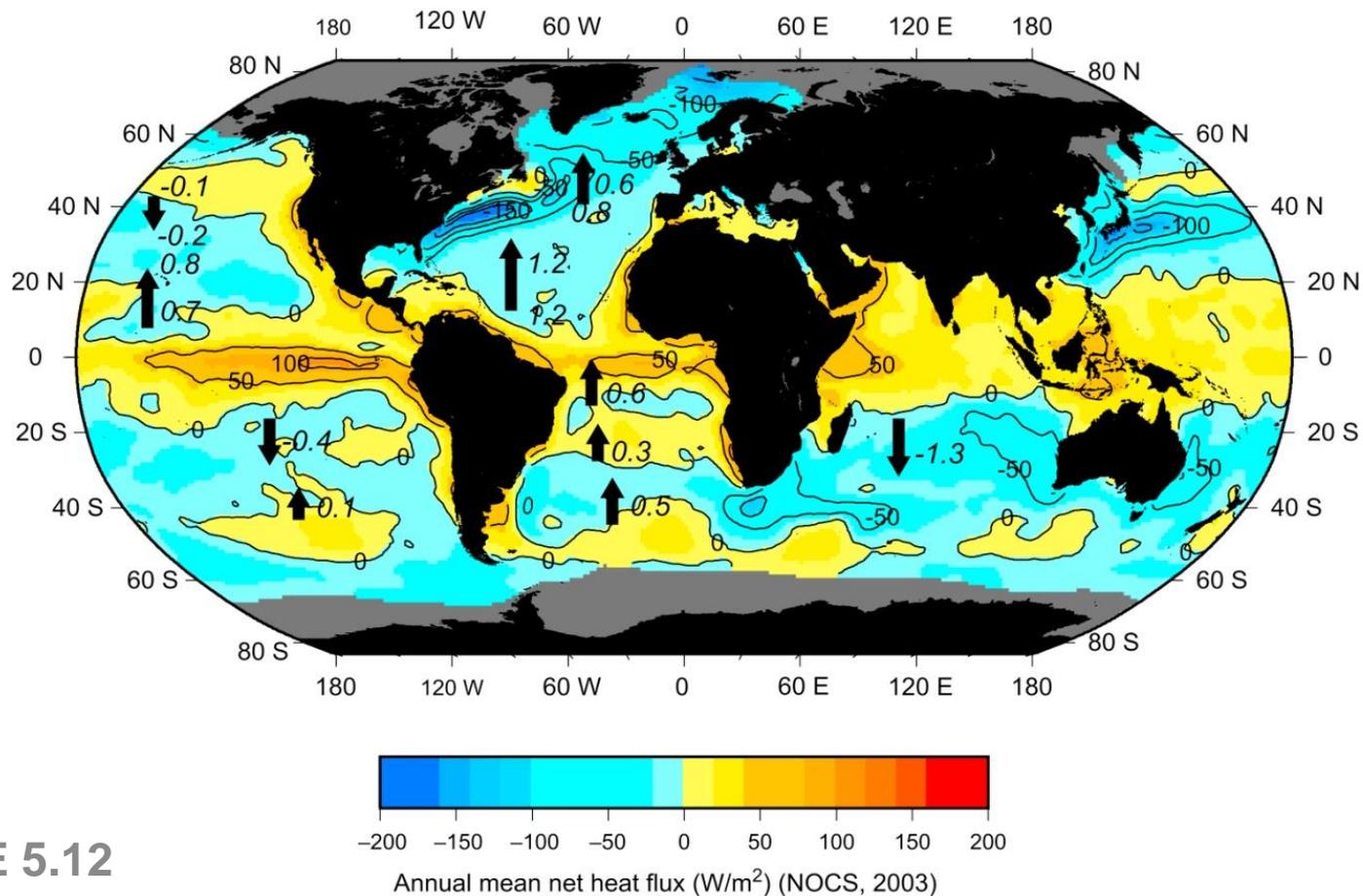


FIGURE 5.12

Annual average net heat flux (W/m^2). Positive: heat gain by the sea. Negative: heat loss by the sea. Data are from the NOCS climatology (*Grist and Josey, 2003*). Superimposed numbers and arrows are the meridional heat transports (PW) calculated from ocean velocities and temperatures, based on *Bryden and Imawaki (2001)* and *Talley (2003)*. Positive transports are northward. The online supplement to Chapter 5 (Figure S5.8) includes another version of the annual mean heat flux, from *Large and Yeager (2009)*. This figure can also be found in the color insert.

Garrafas de Nansen e de Niskin



**Nansen water bottles
before (I), during (II), and
after (III) reversing.
(From Dietrich et al. 1980)**

Fig. 6.16. Water sampling bottles (Nansen bottles), circa 1960 for mounting individually on a wire with reversing thermometer racks.



Fig. 6.17. Single Niskin bottle

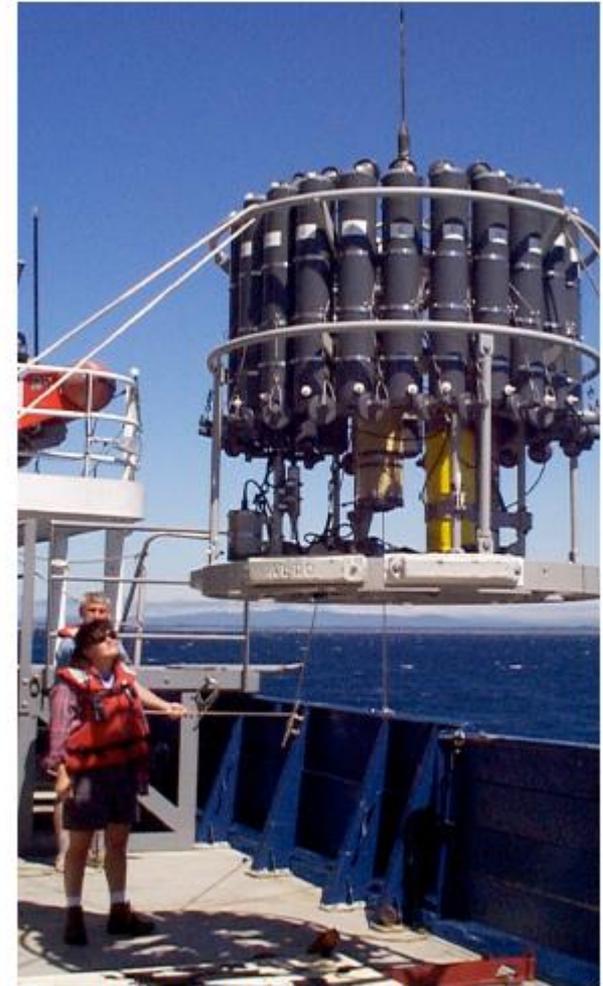


Fig. 6.18. Rosette sampler. Large sampler used in the World Ocean Circulation Experiment, with 36 10-liter Niskin bottles, an acoustic pinger (lower left), an LADCP (center, yellow long), a CTD (bottom, horizontal), and transmissometer (center, yellow short).

Termômetros de inversão



Nansen water bottles before (I), during (II), and after (III) reversing.
(From Dietrich et al. 1980)

Fig. 6.16. Water sampling bottles (Nansen bottles), circa 1960 for mounting individually on a wire with reversing thermometer racks.

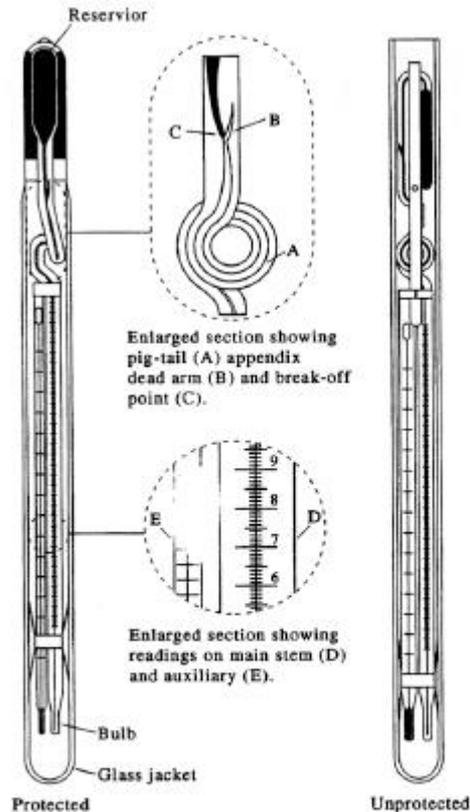


Fig. 6.21. Protected and unprotected reversing thermometers.



Fig. 6.22. Reading reversing thermometers

Métodos de medição

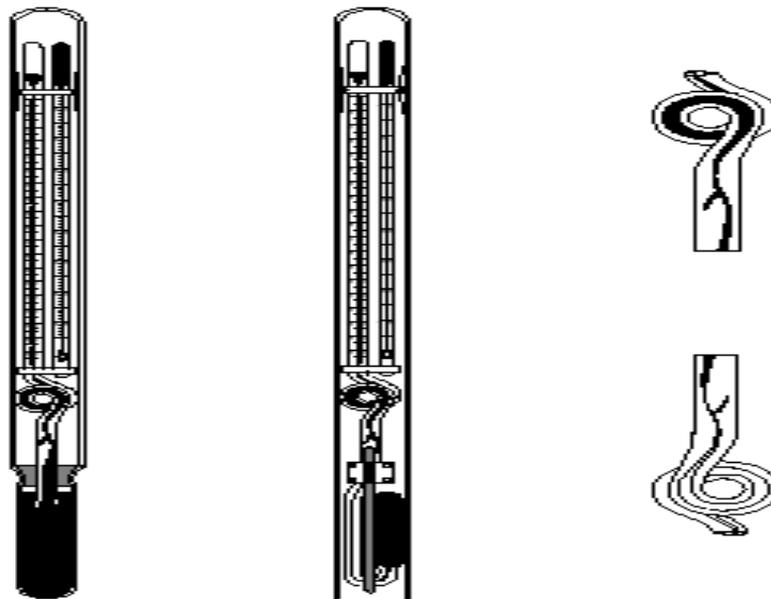
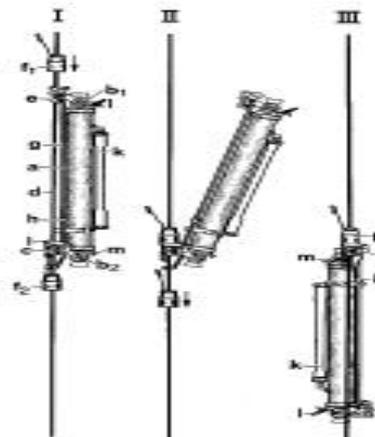
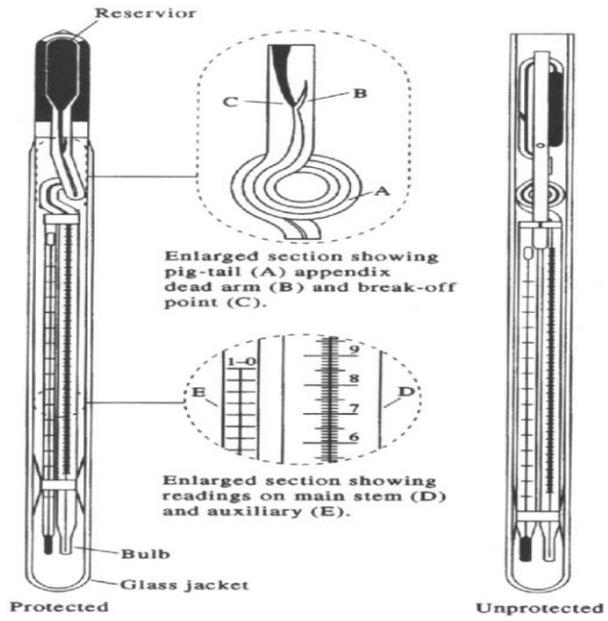


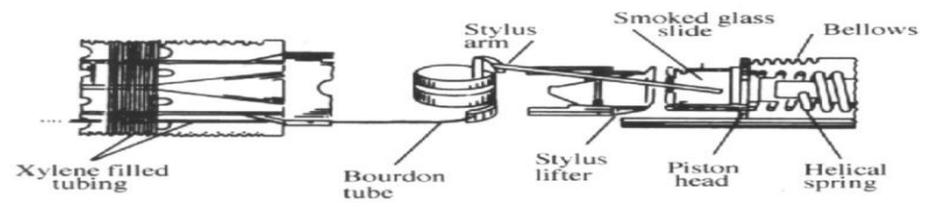
Figure 6.13 **Left:** Protected and unprotected reversing thermometers in set position, before reversal. **Right:** The constricted part of the capillary in set and reversed positions (From von Arx, 1962).





Temperature element

Pressure element



Bati-Termógrafo mecânico

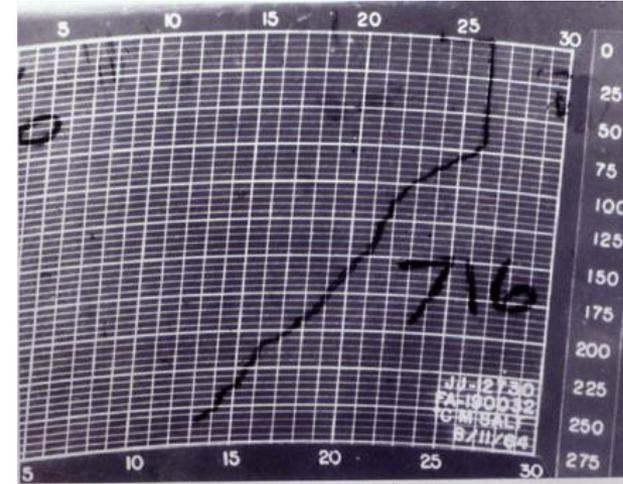
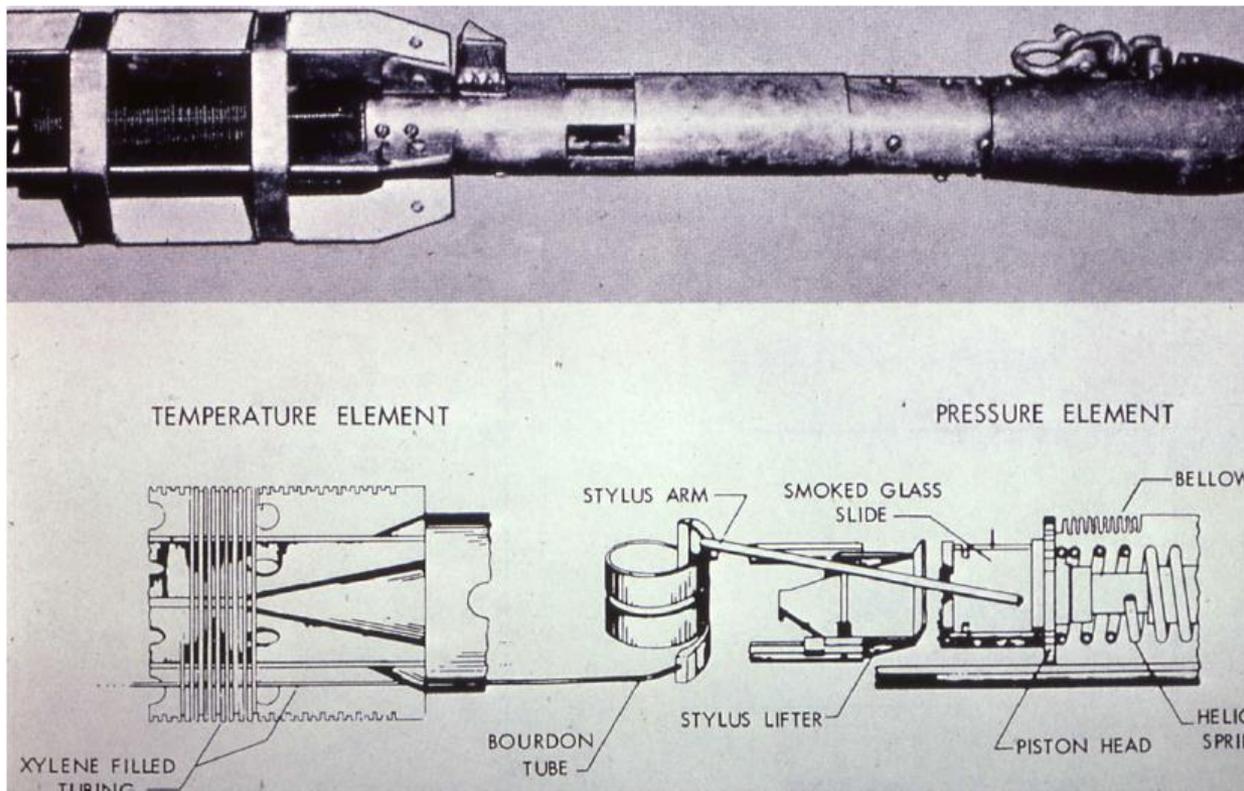
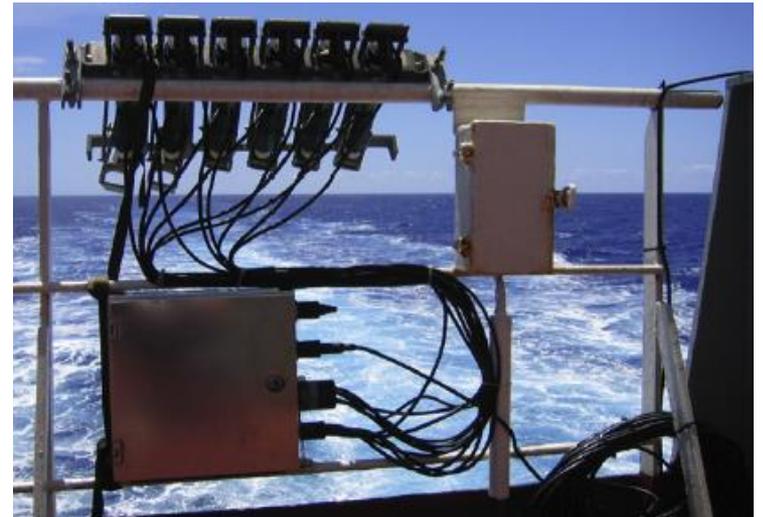
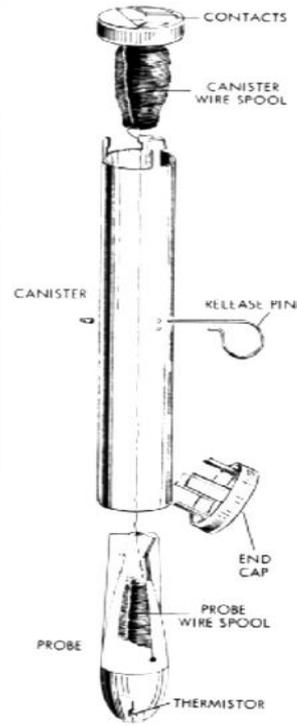
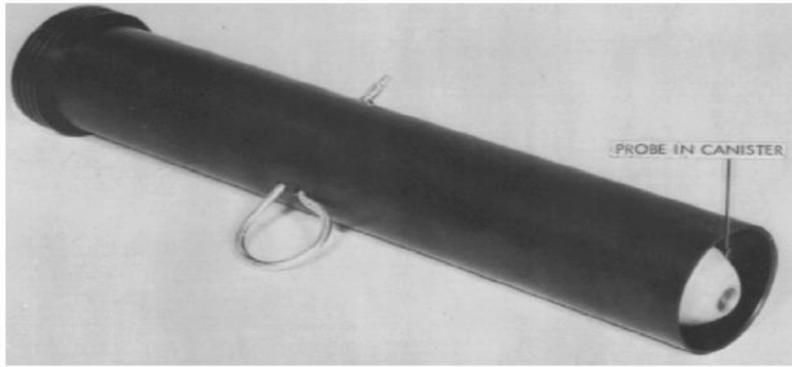


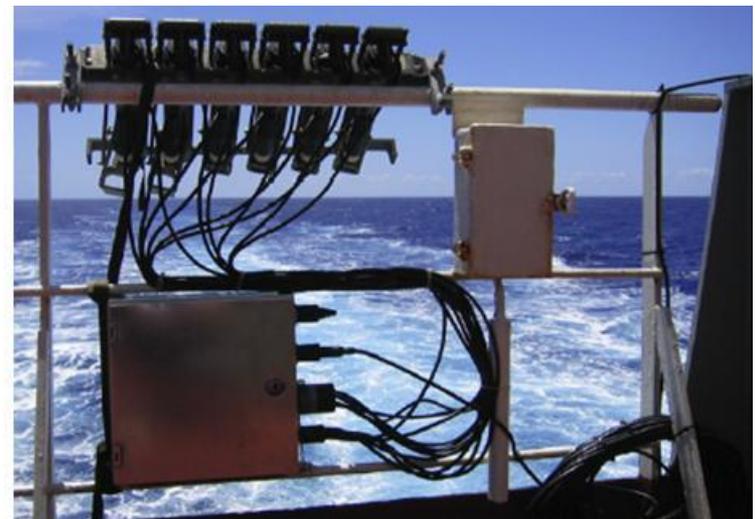
FIGURE S16.16 Mechanical bathythermograph (MBT), in use from 1951 to 1975. Source: Neumann and Pierson (1966).

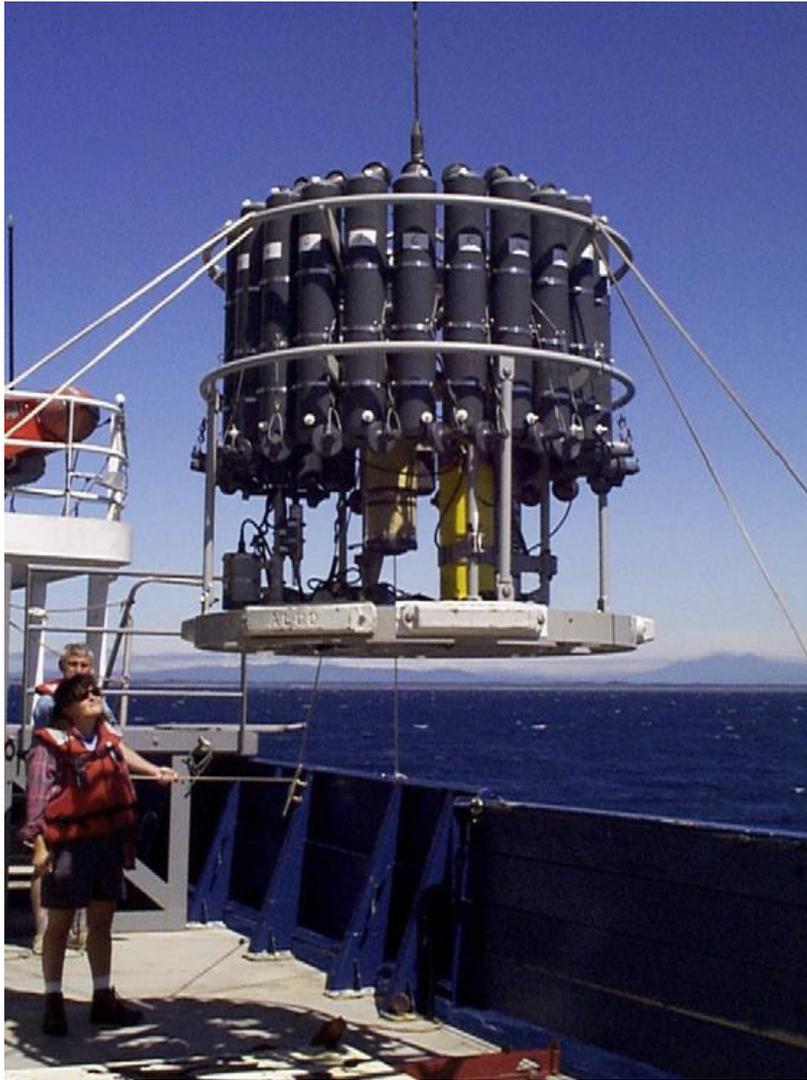




eXpendable-BT

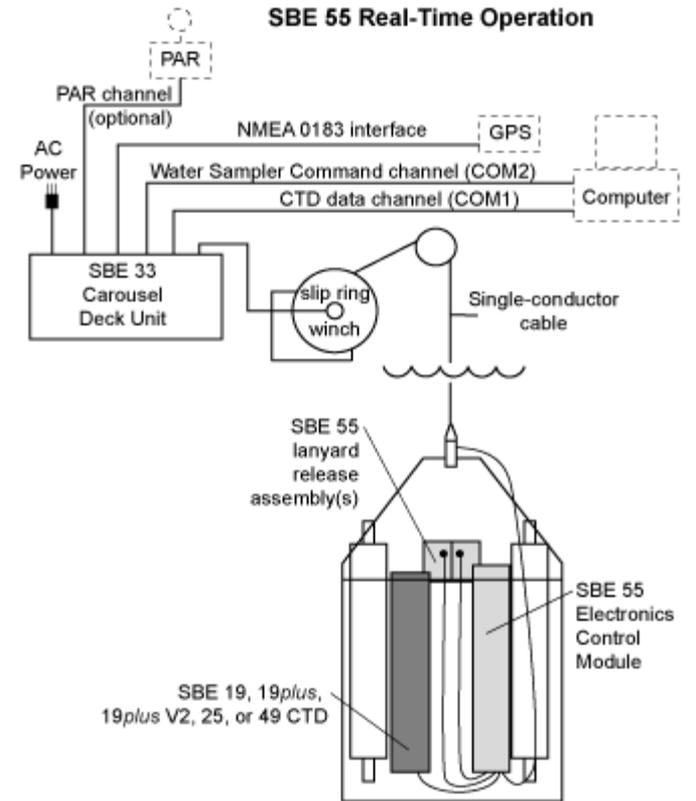
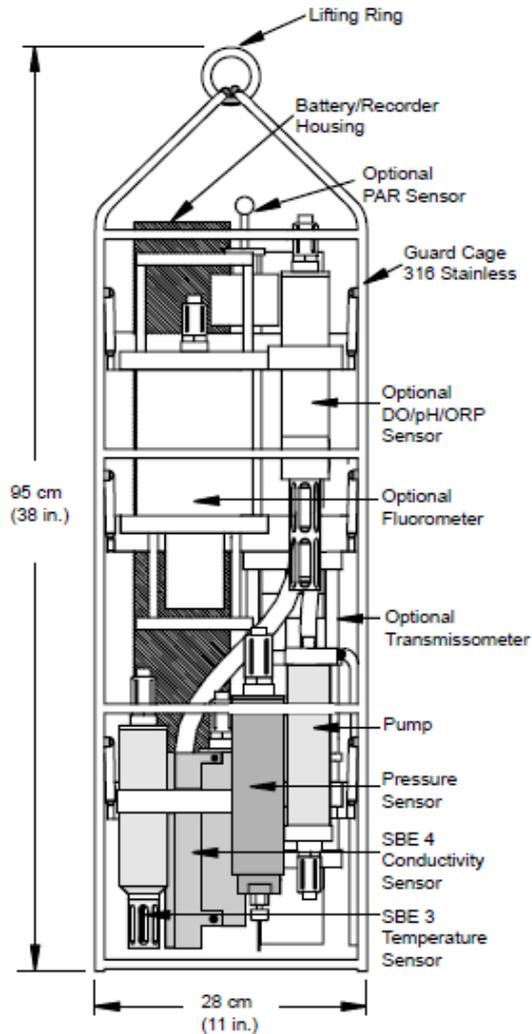
FIGURE S16.18 An expendable bathythermograph (XBT). Source: From NOAA UOTC (2009).

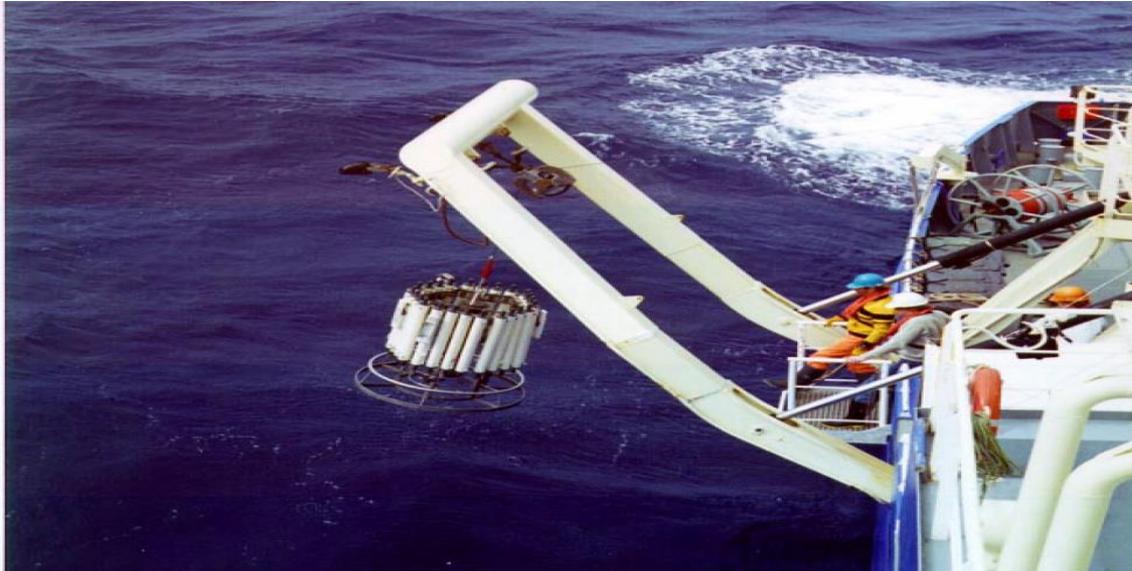




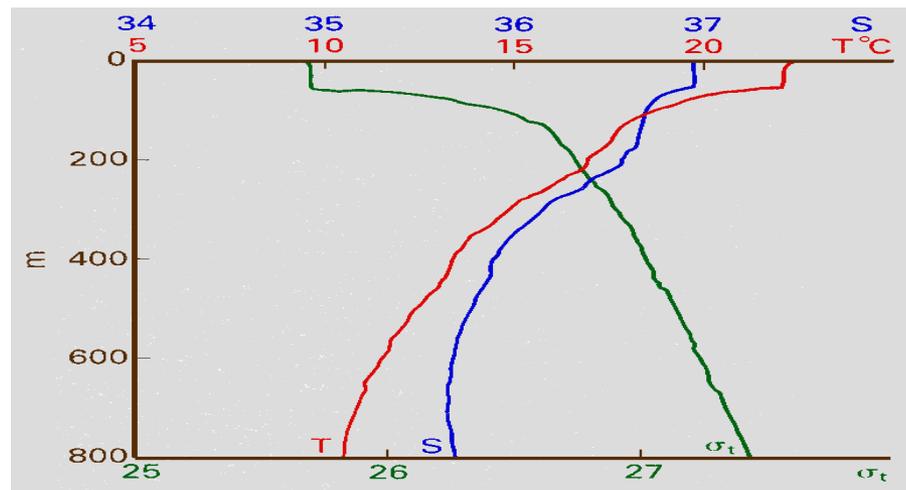
Conductividade Temperatura Depth
(Profundidade)

CTD

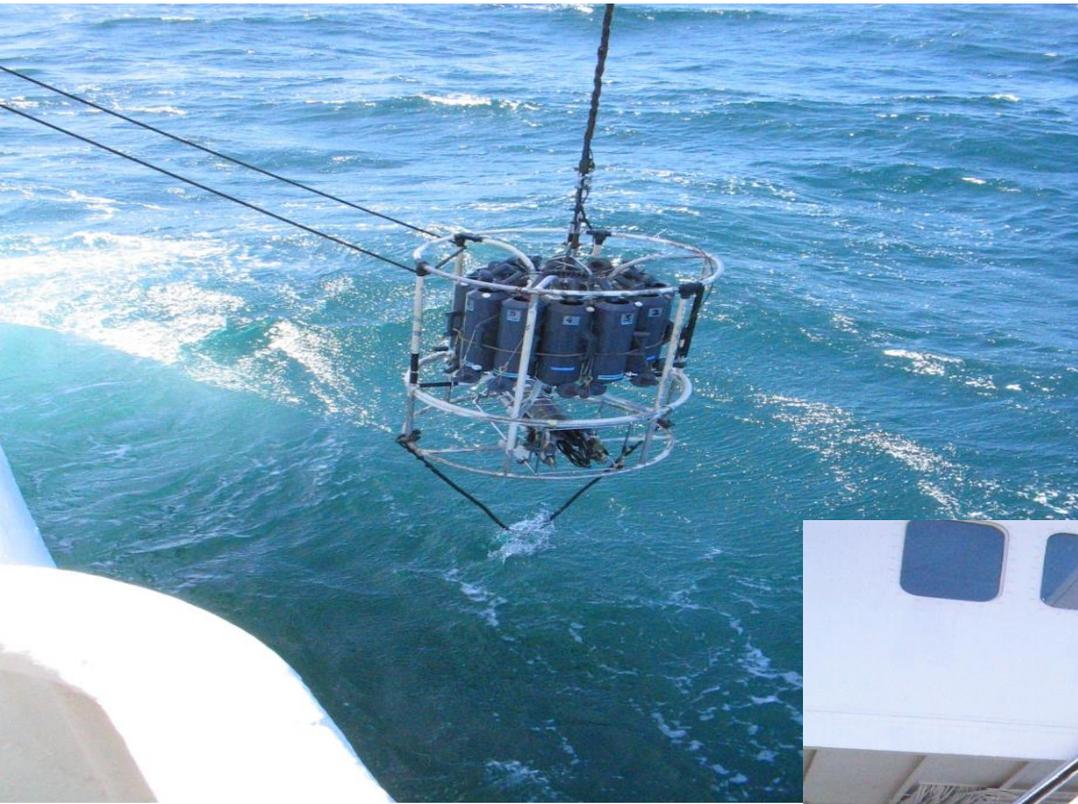




A medida hidrográfica : Roseta + CTD.



O perfil de T/S/Sigma-t resultante



Barcos de oportunidade

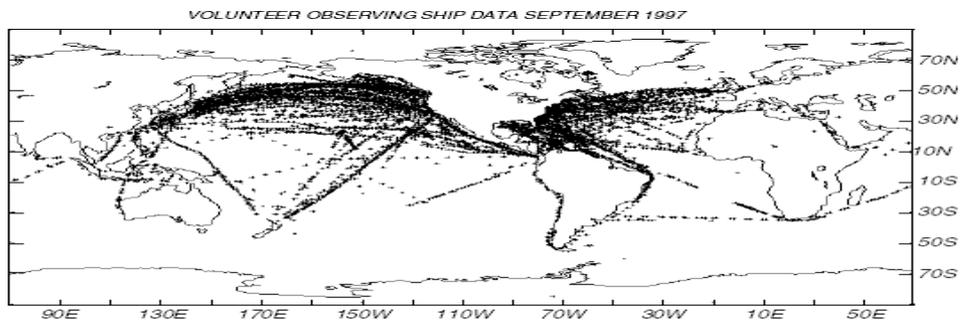
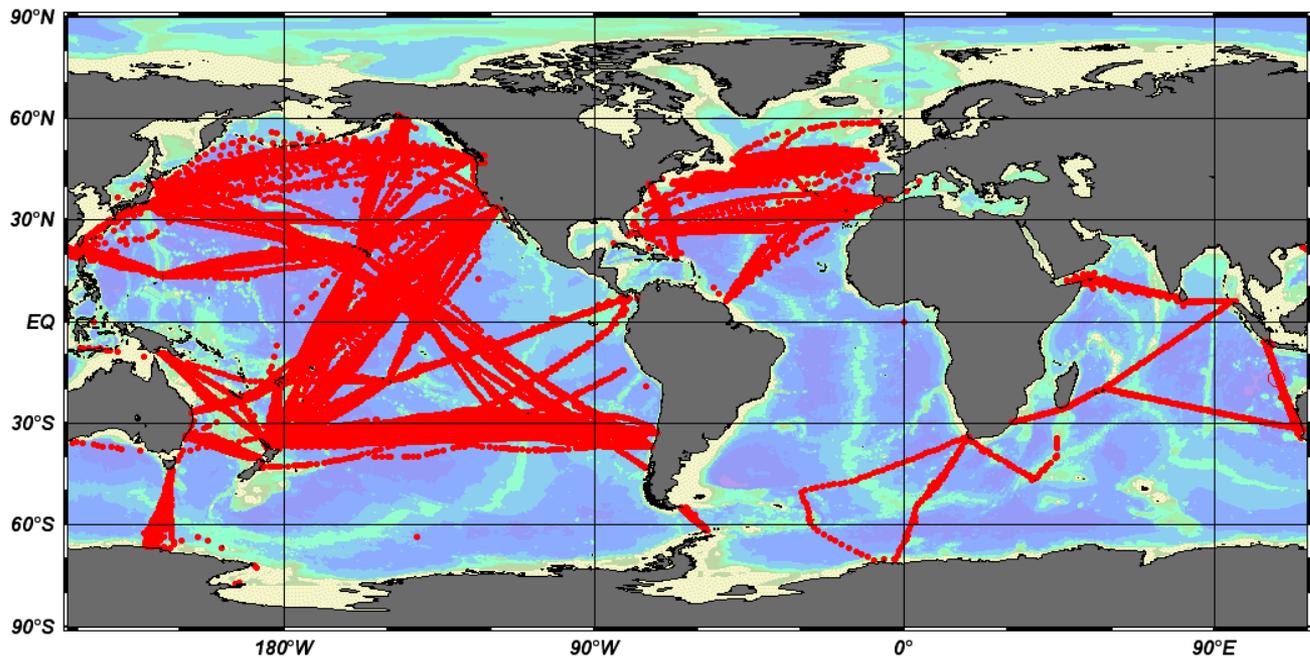
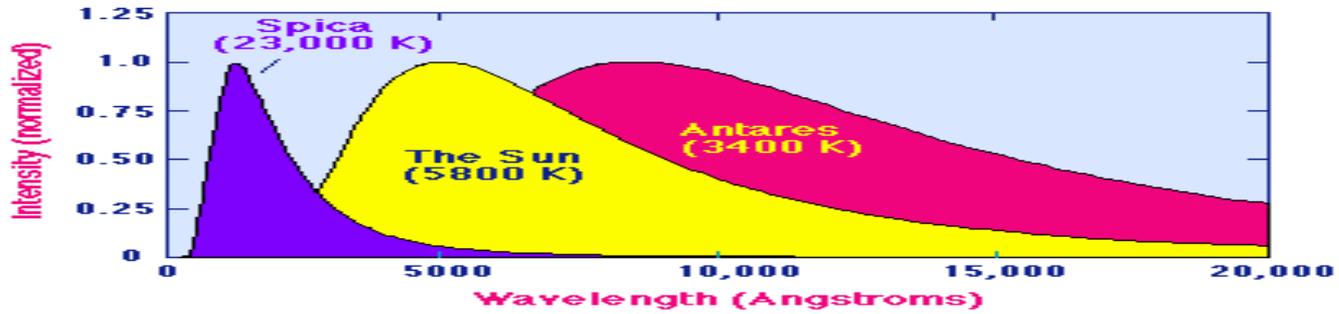


Figure 4.5 Location of surface observations made from volunteer observing ships and reported to national meteorological agencies. (From NOAA, National Ocean Service)

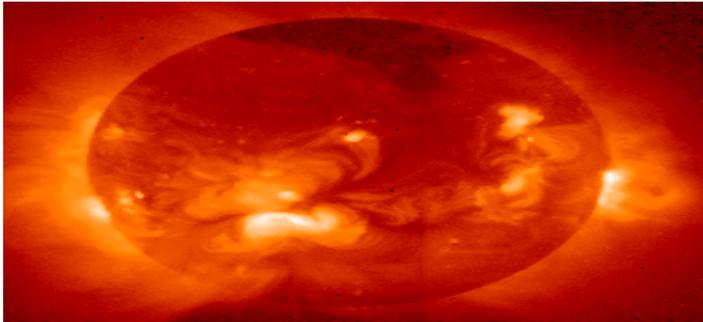
eWOCE UOT - High Density Lines



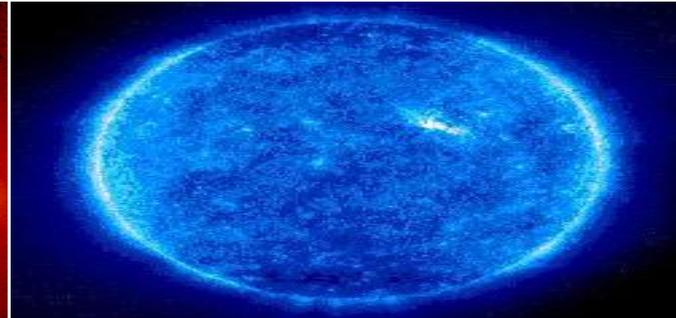
Sensoriamento Remoto Espectros de Radiação E.M.



Espectro Visível: 4000 a 7000 Å ou 400 a 750 μm.

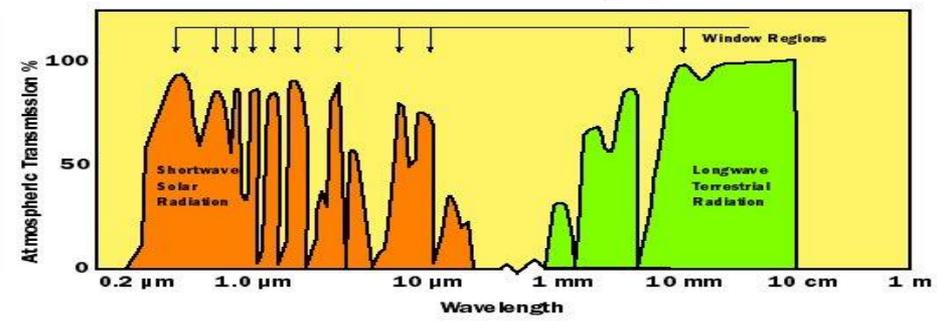
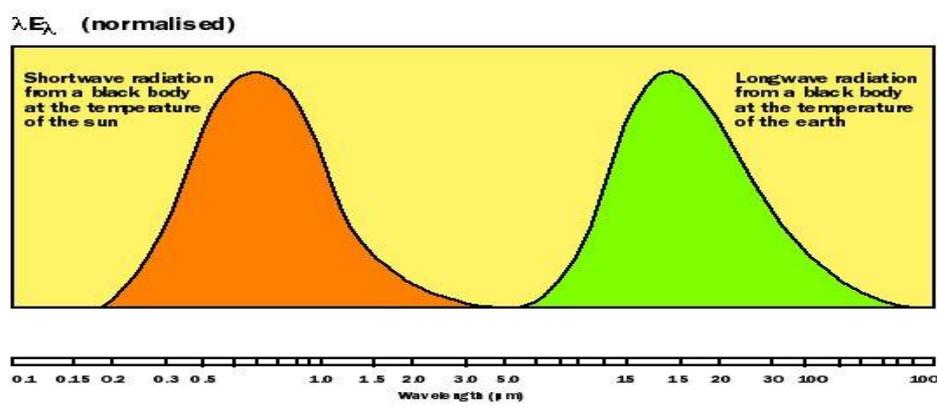
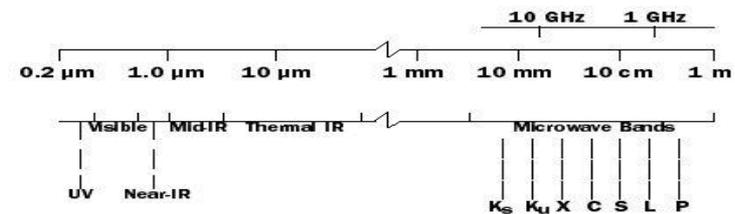
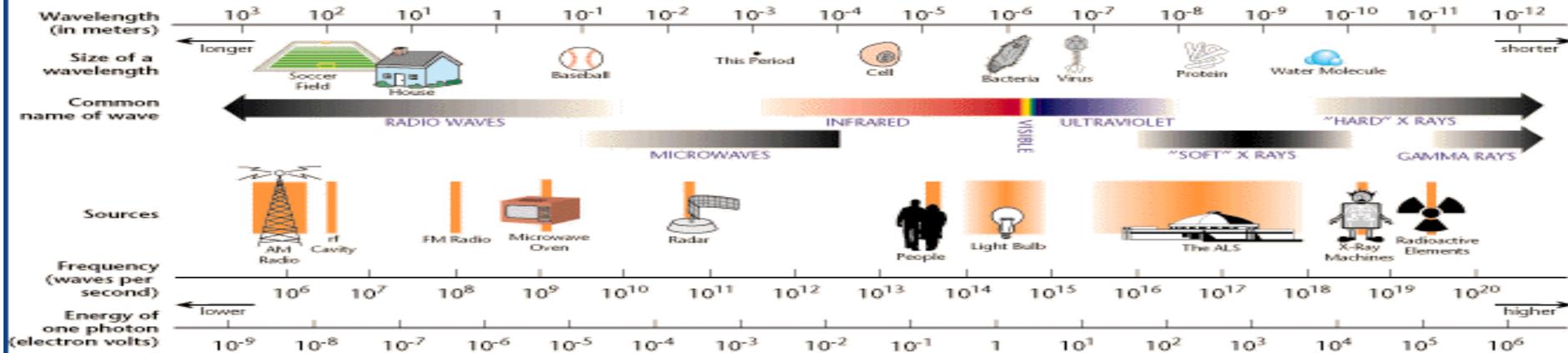


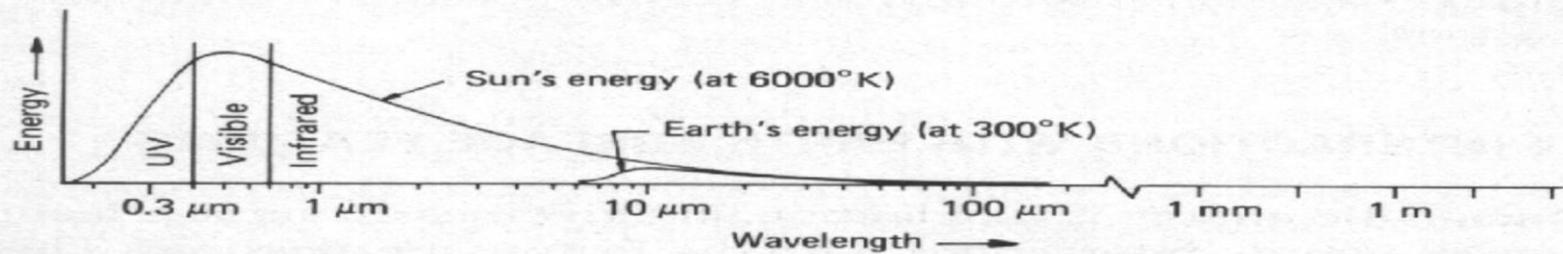
Sol (λ =raios-X)



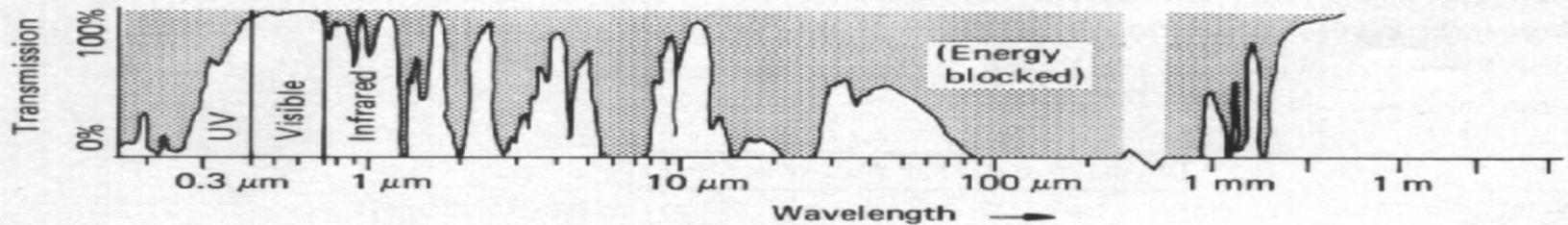
Sol (λ =UV)

THE ELECTROMAGNETIC SPECTRUM

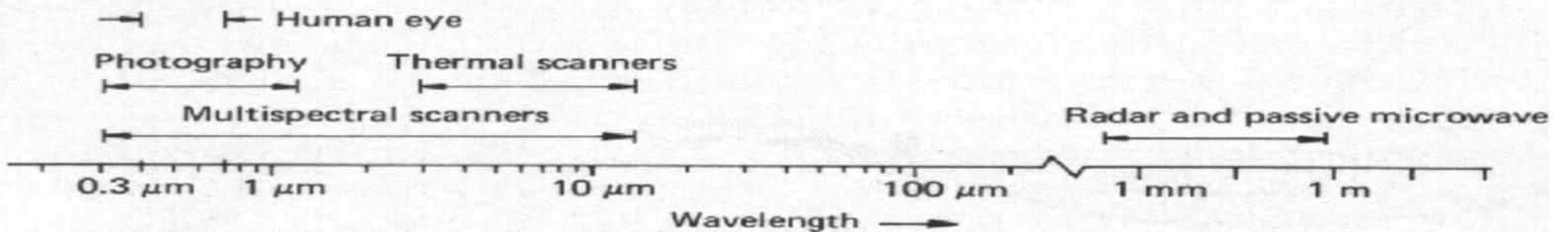


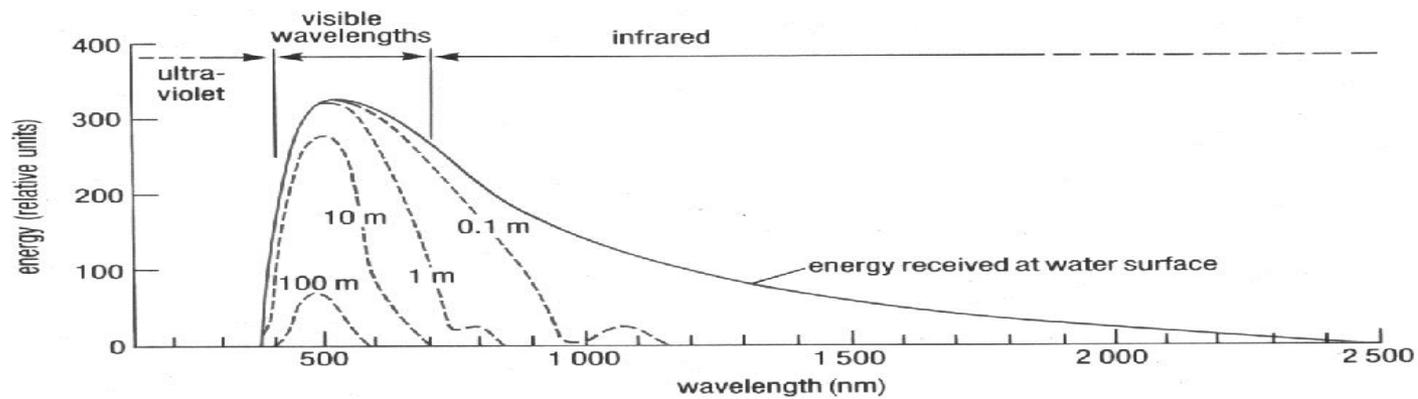
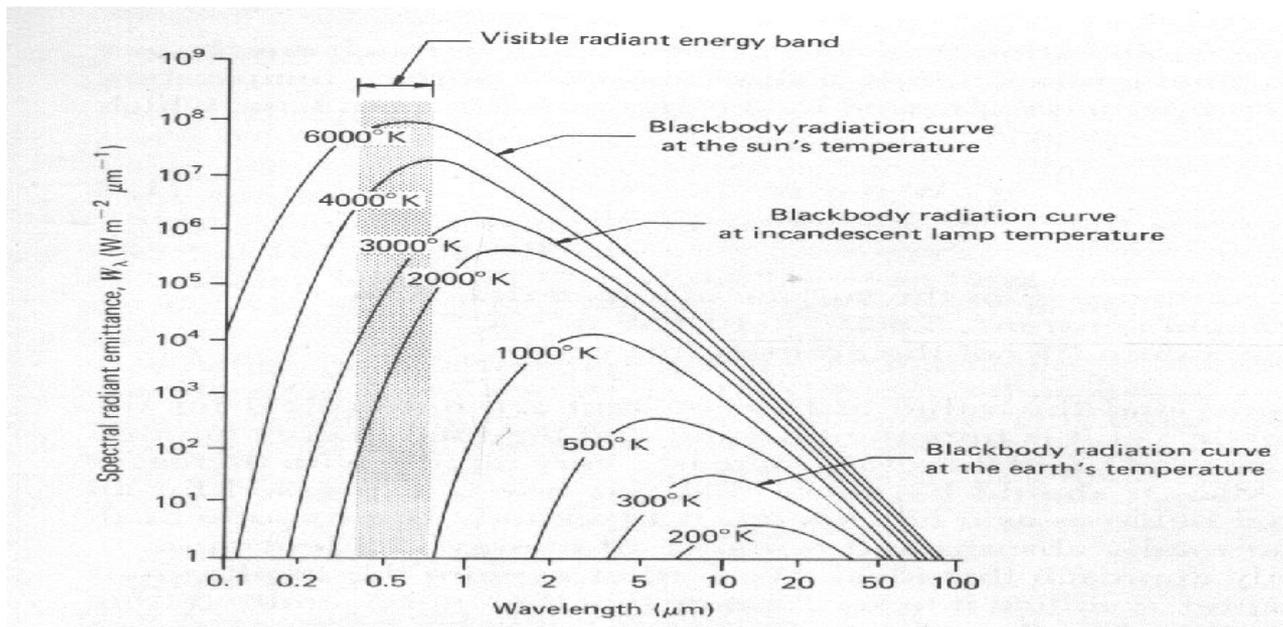


(a) Energy sources

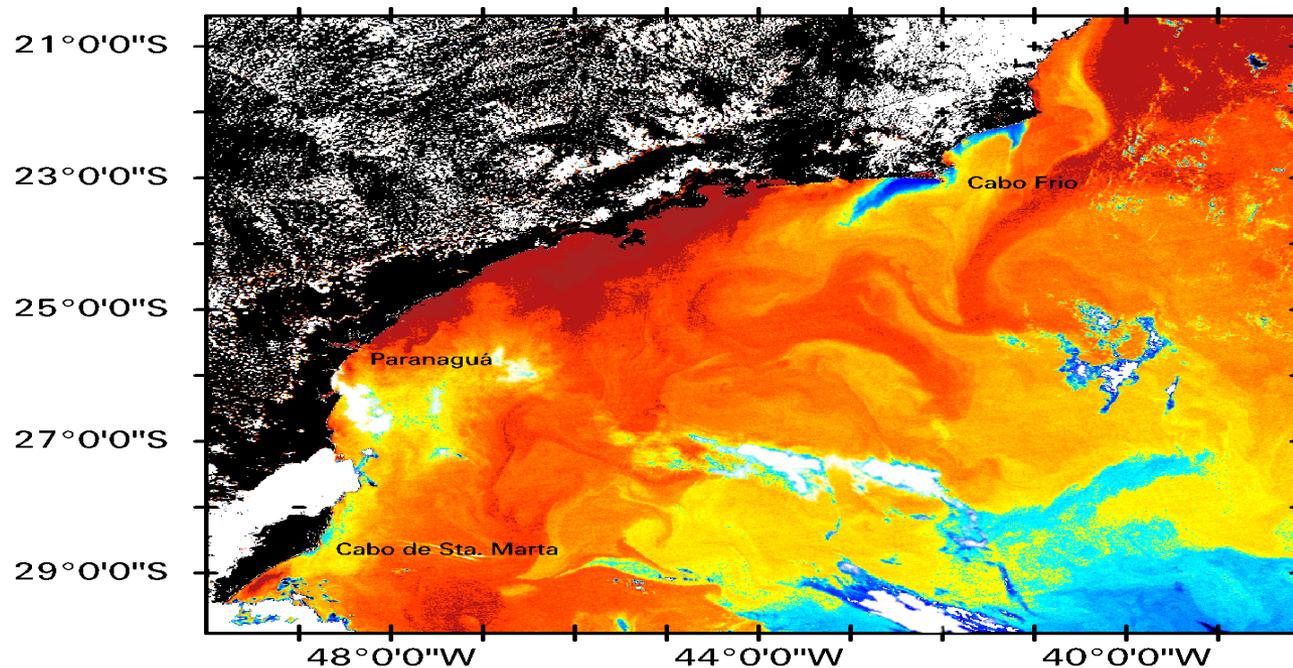


(b) Atmospheric transmittance

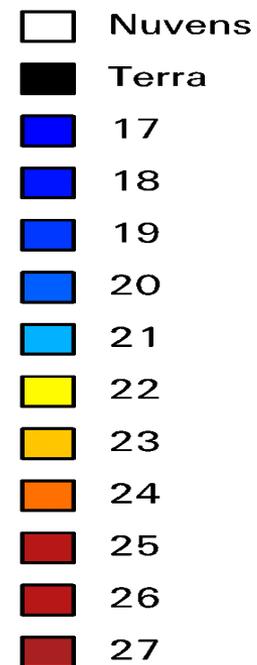




TEMPERATURA SUPERFICIAL DO MAR



TEMPERATURA (°C)



SENSOR/SATÉLITE: AVHRR/NOAA-14

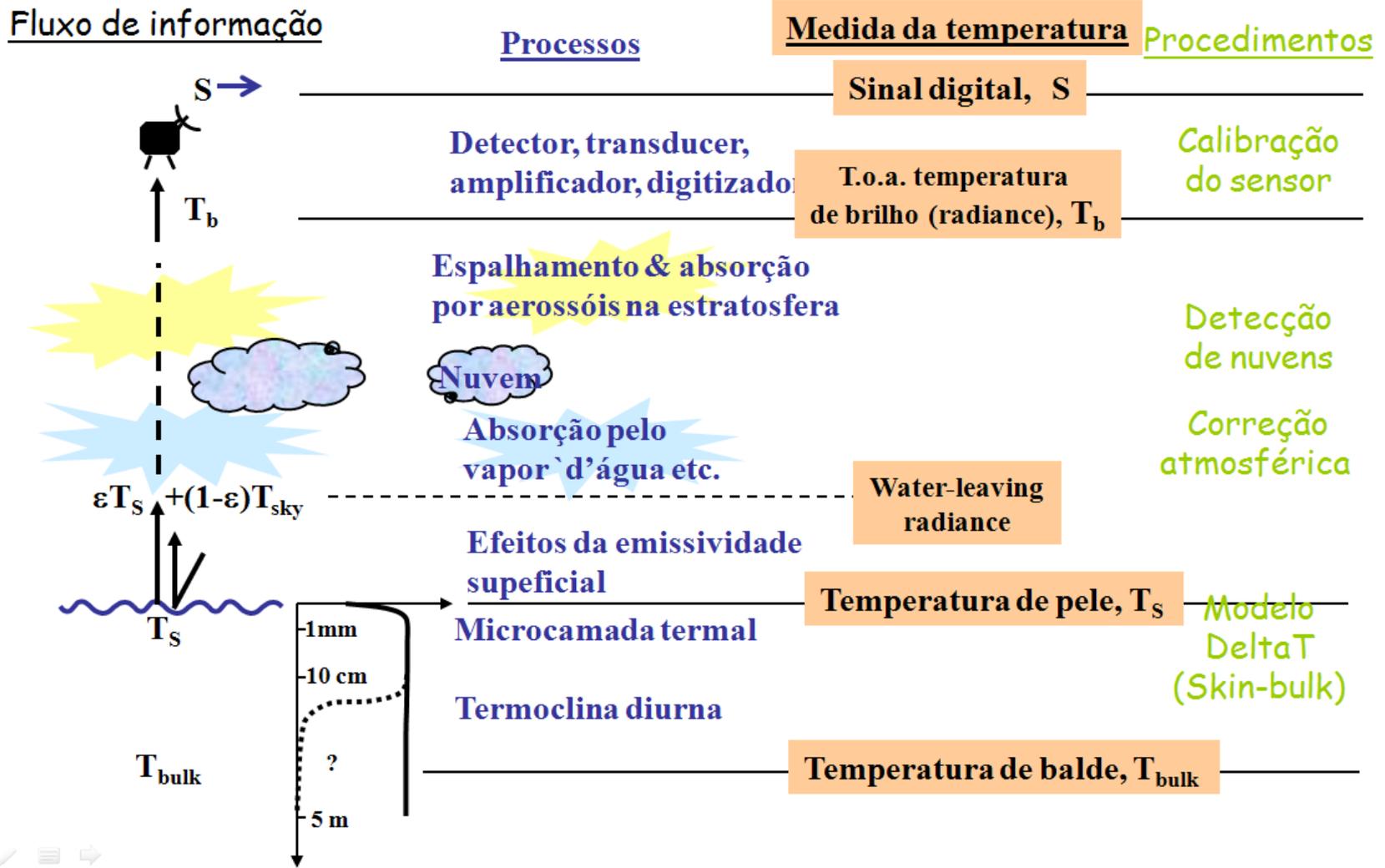
ALGORITMO: DAYTIME NLSST

AQUISIÇÃO: 07/11/95 - 16:54 GMT

REGIÃO CENTRAL DA BACIA DE SANTOS

CONVÊNIO PETROBRÁS/FURG

Processos que afetam a medida de TSM



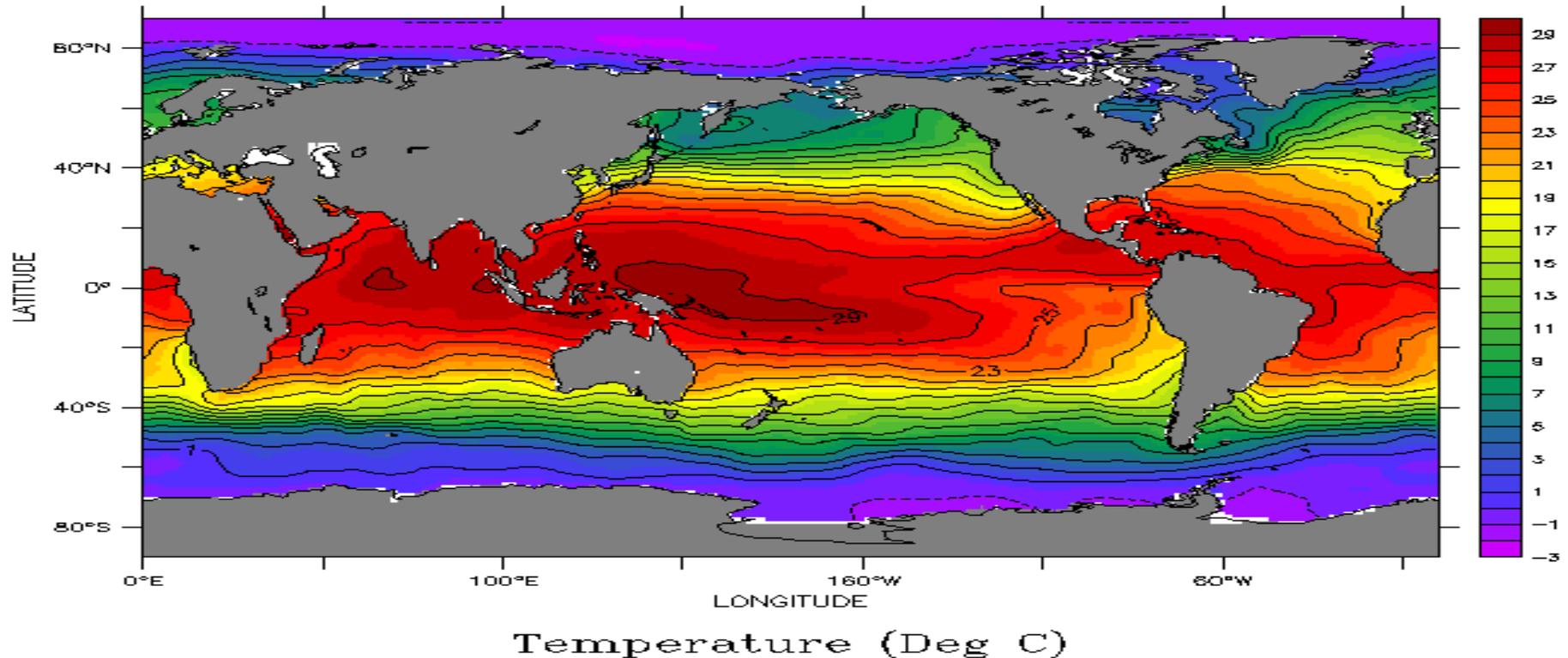
Distribuição horizontal –camada de mistura

NOAA/PMEL TMAP  FERRET Ver 4.0

DEPTH (m) : 0

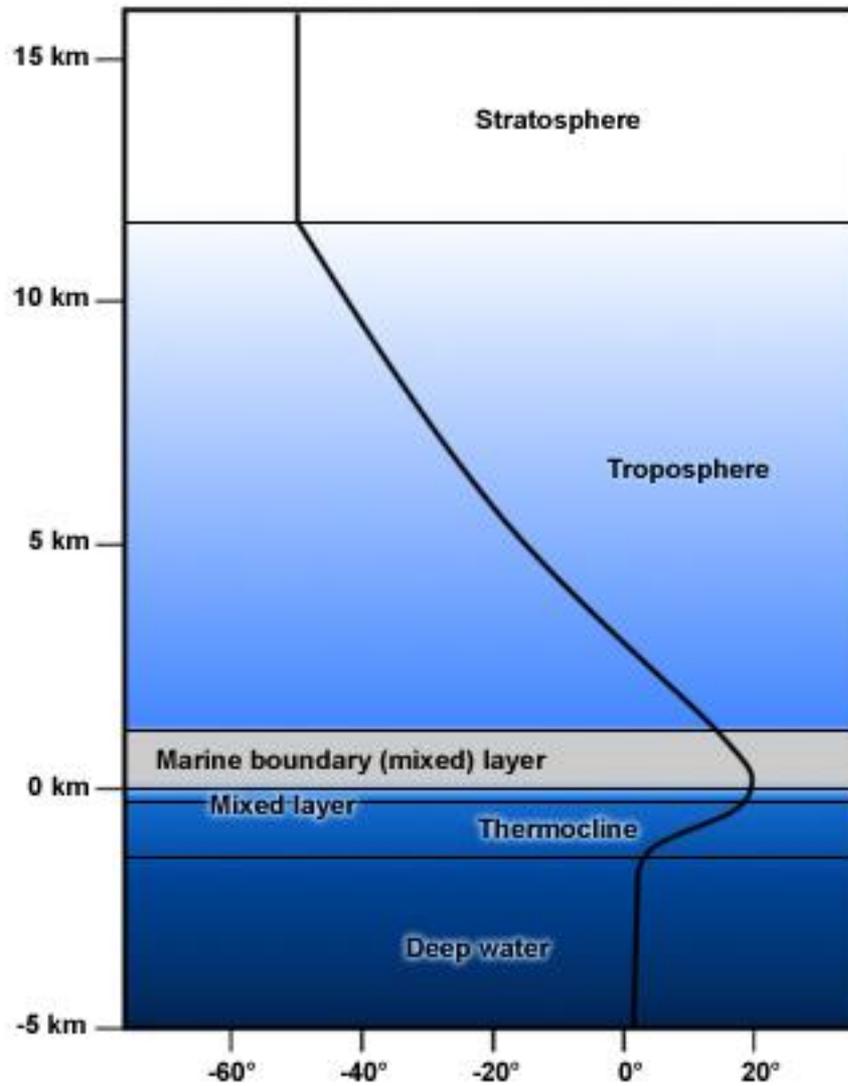
DATA SET: ocean-atlas-annual

World Ocean Atlas 1994 * 1x1 Degree Annual Means



Mean sea-surface temperature calculated from the optimal interpolation technique (Reynolds and Smith, 1995) using shipreports and AVHRR measurements of temperature. Contour interval is 1° C with heavy contours every 5° C. Shaded areas exceed 29° C.

Temperature Profile from Seafloor to Stratosphere



DISTRIBUIÇÃO VERTICAL DA TEMPERATURA ENTRE 60 S E 60 N

~150 m

Camada de mistura – radiação solar, vento, ondas e correntes

Termoclina Permanente

onde a T decresce rapidamente –

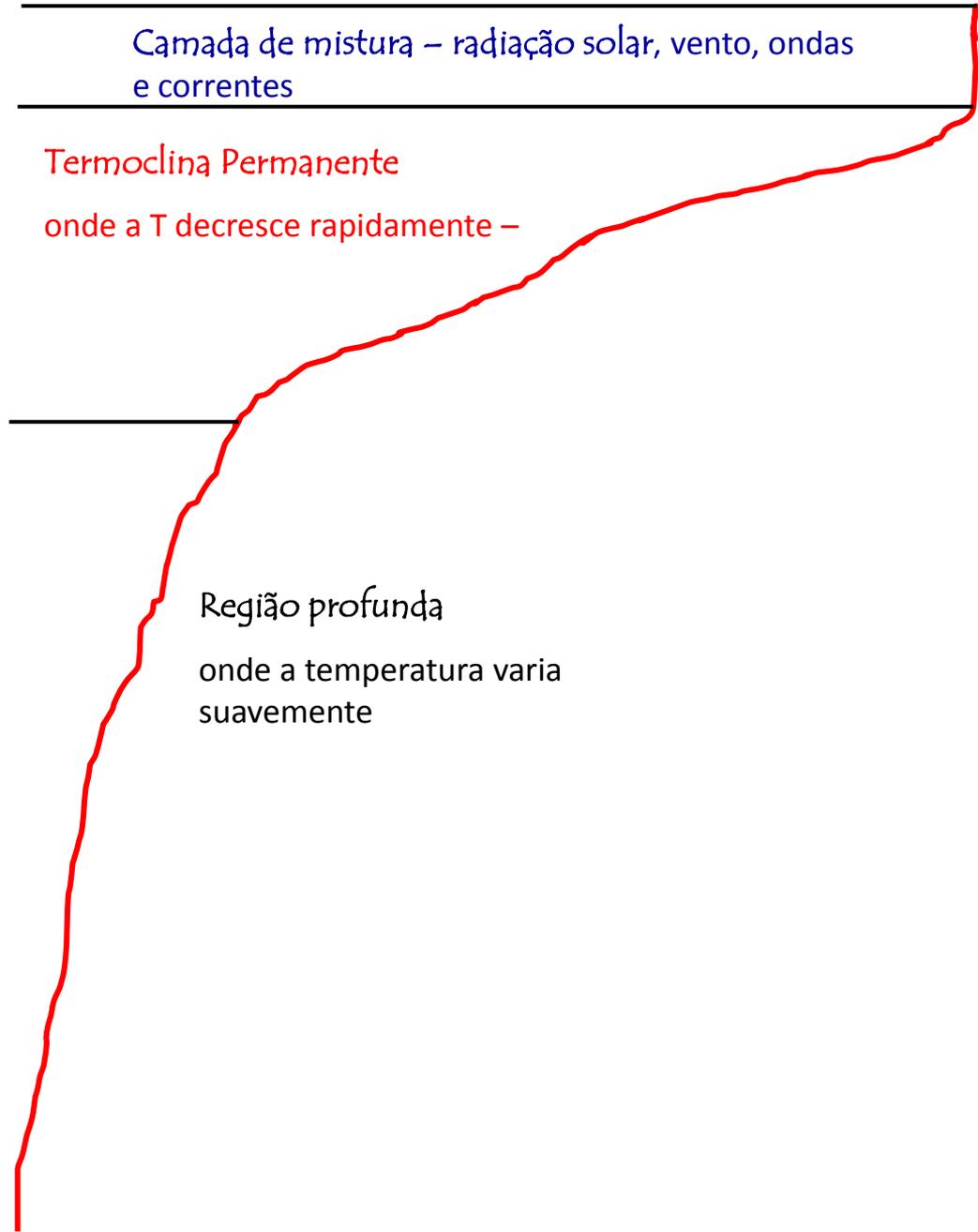
~1000 m

Região profunda

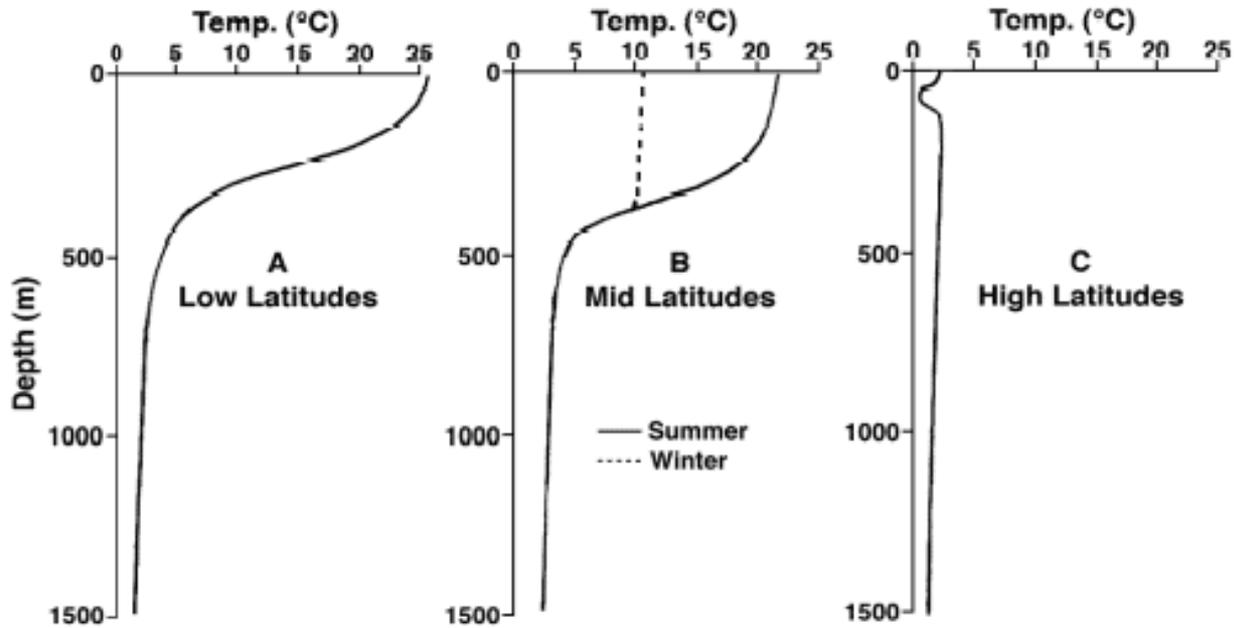
onde a temperatura varia suavemente

Essa estrutura é alterada de acordo com a latitude. Em médias latitudes aparece, também, uma termoclina sazonal.

Em altas latitudes a camada de mistura vai, praticamente, até o fundo do oceano, com perfil com decréscimo suave da T até o fundo



Typical Temperature Profiles



EM FUNÇÃO DA LATITUDE AS ZONAS SE DIVIDEM EM:

BAIXAS LATITUDES Não existem \neq significativas nas estações do ano

T das H₂O_s superficiais são sempre altas

Termoclina Permanente

Camada de mistura de pouca espessura.

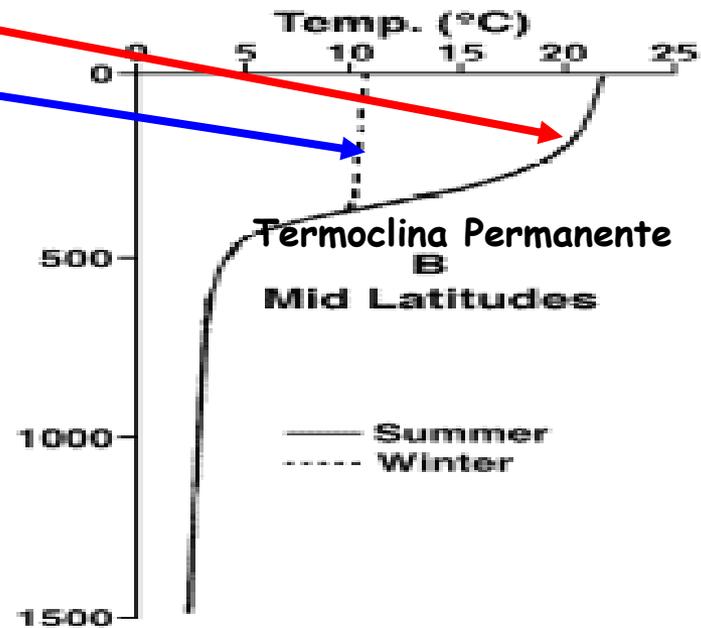
Divergência Equatorial

MÉDIAS LATITUDES

Termoclina Permanente & sazonal

Verão

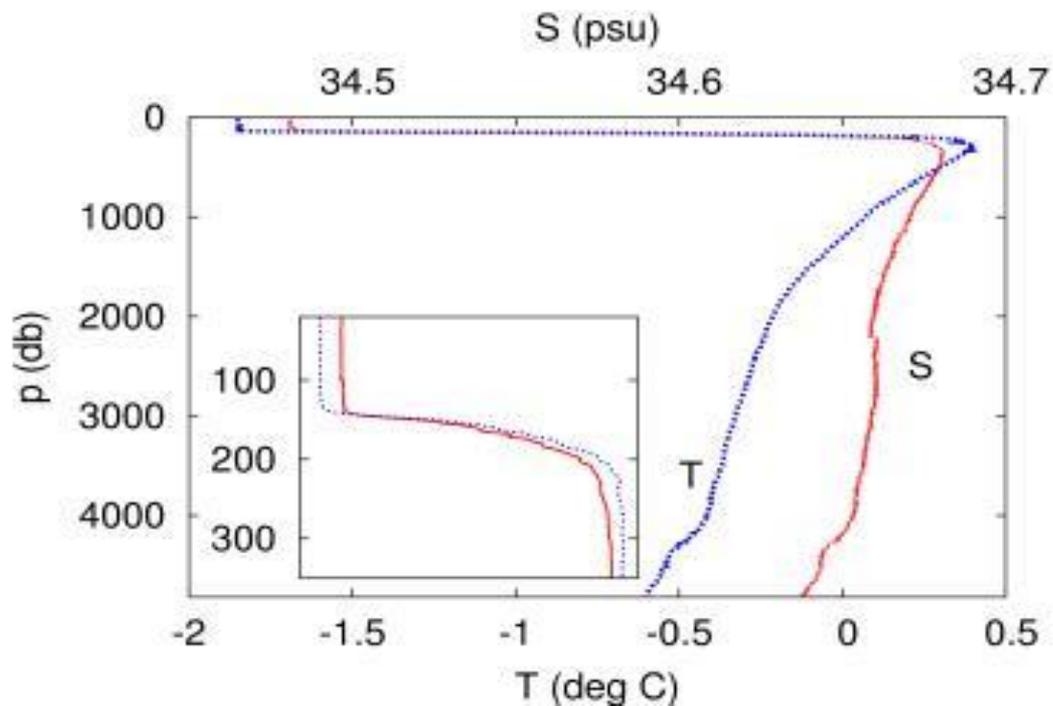
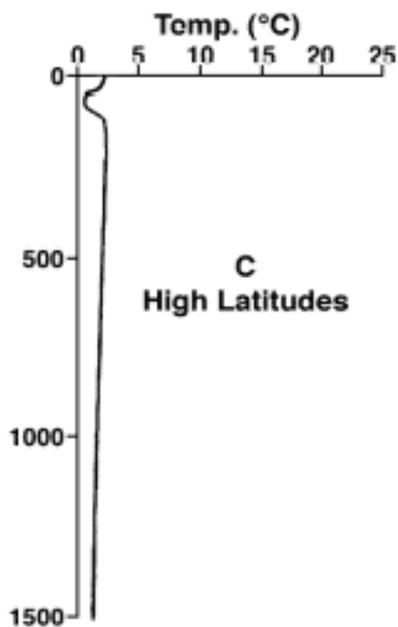
Inverno



ALTAS LATITUDES

Inexistência da Termoclina Permanente

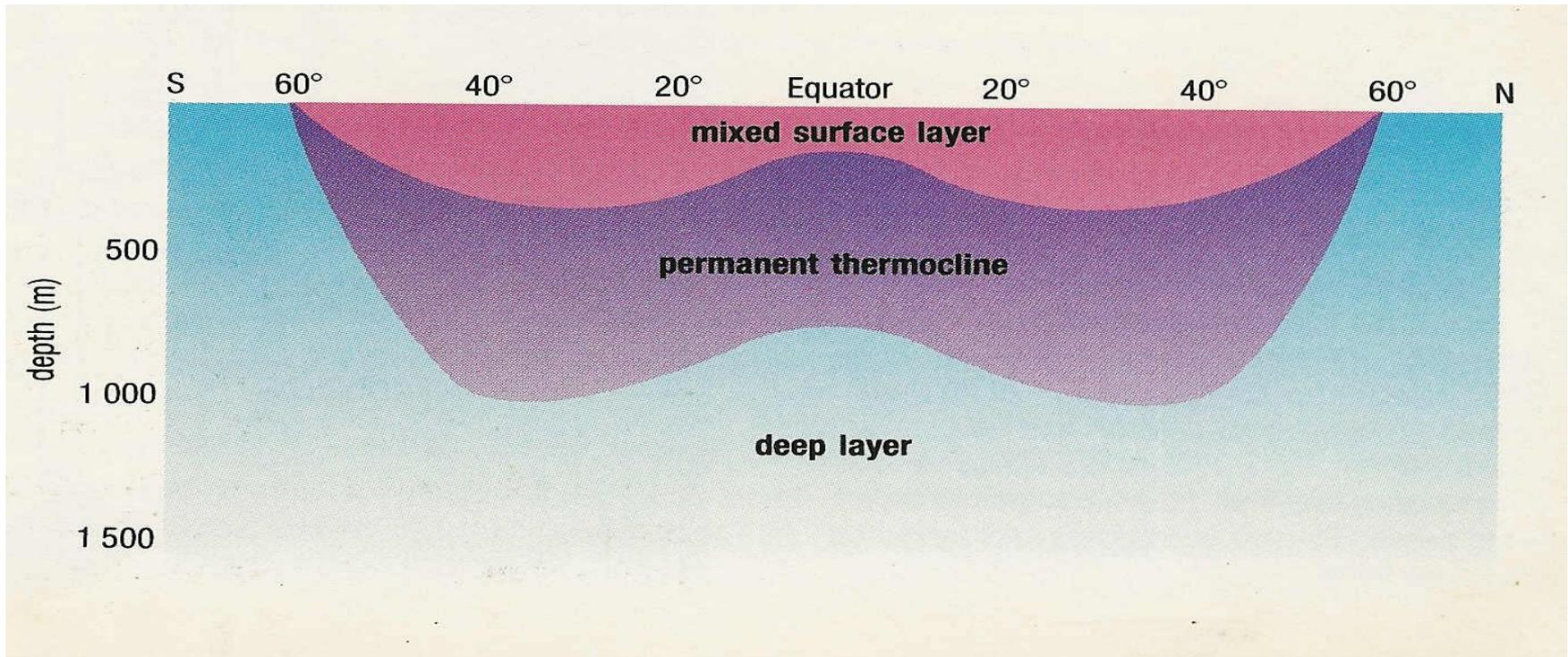
T quase Cte ao longo da coluna d' água

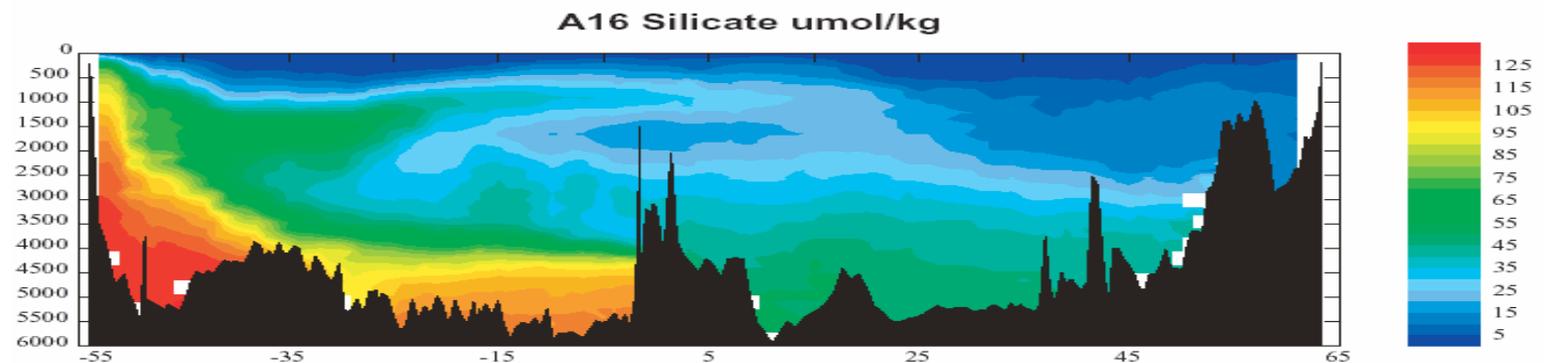
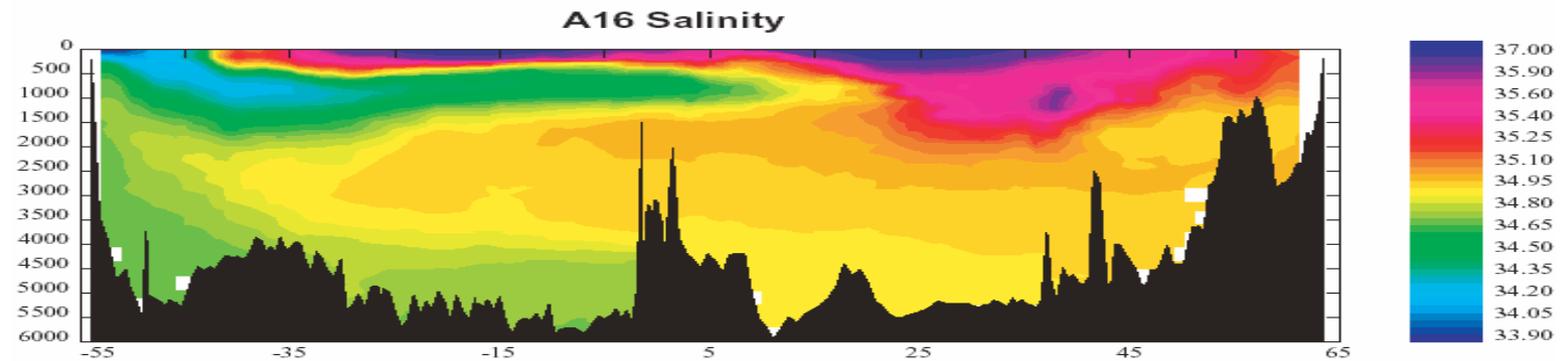
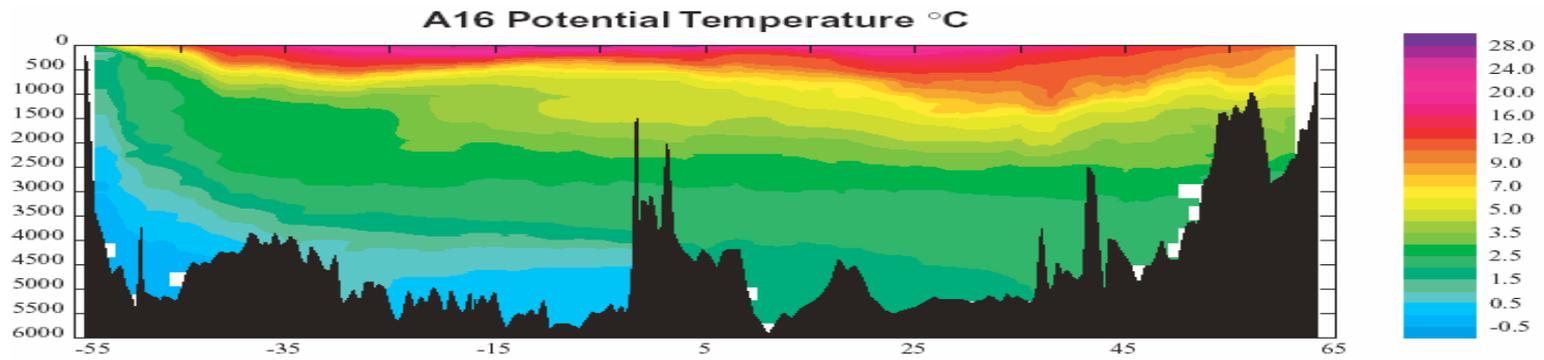


Mar de Weddell

Região Subantártica

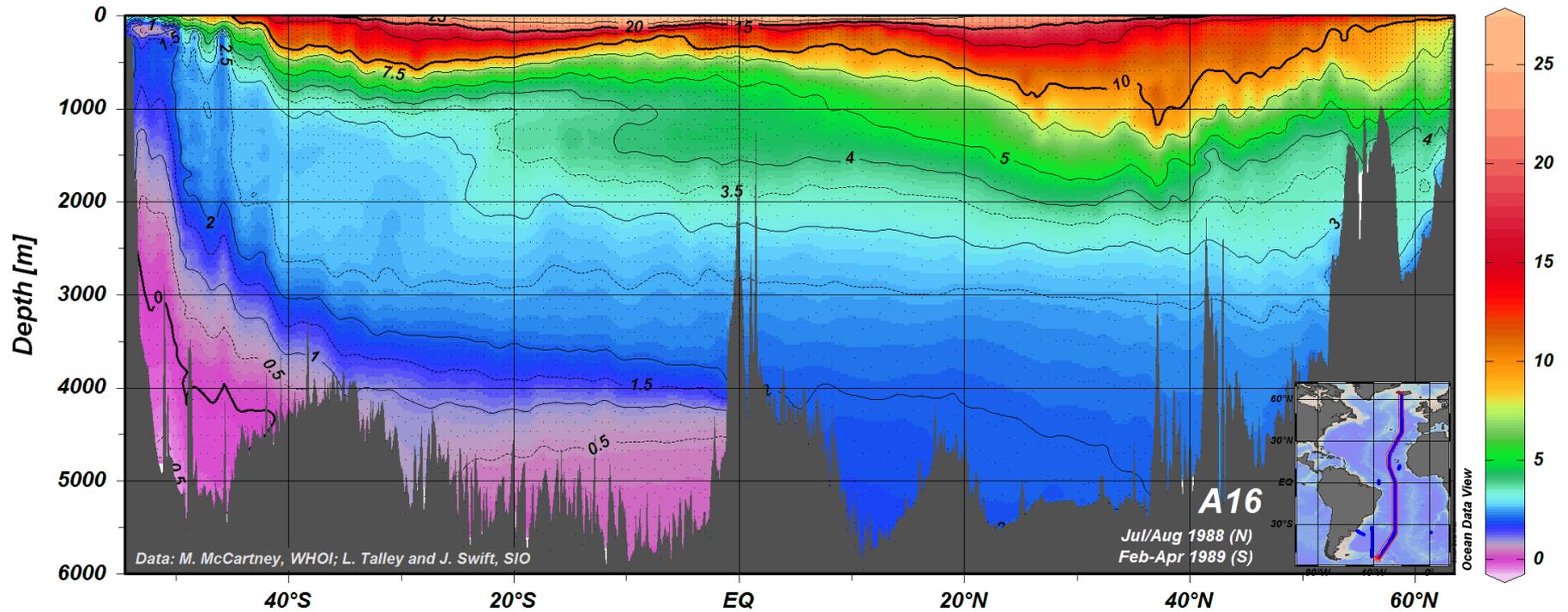
Resumo da distribuição latitudinal, em profundidade, da Temperatura



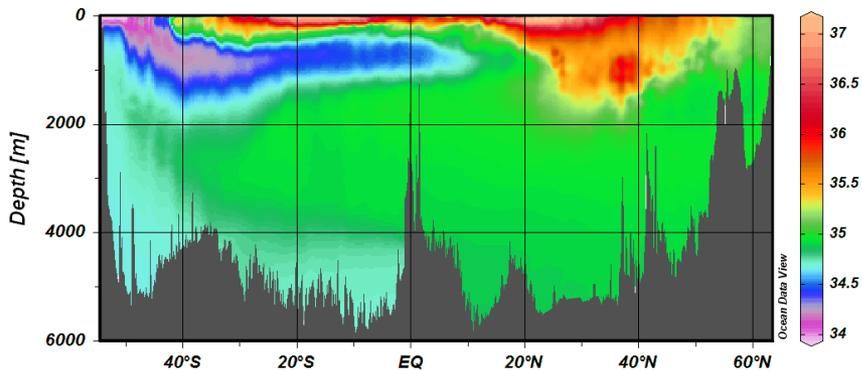


A16 - a WOCE mid Atlantic transect from south (left) to north.
(WOCE IPO)

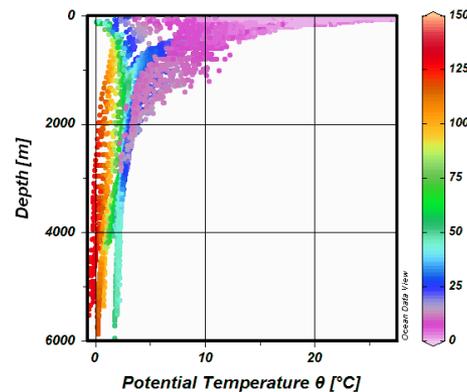
Potential Temperature θ [$^{\circ}\text{C}$]



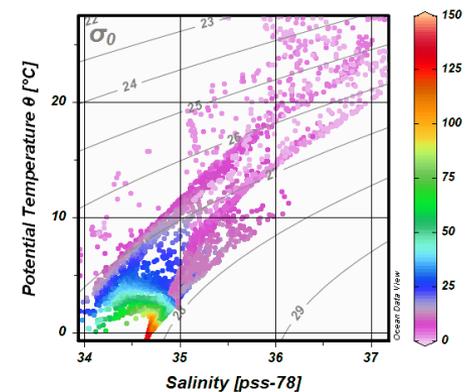
Salinity [pss-78]



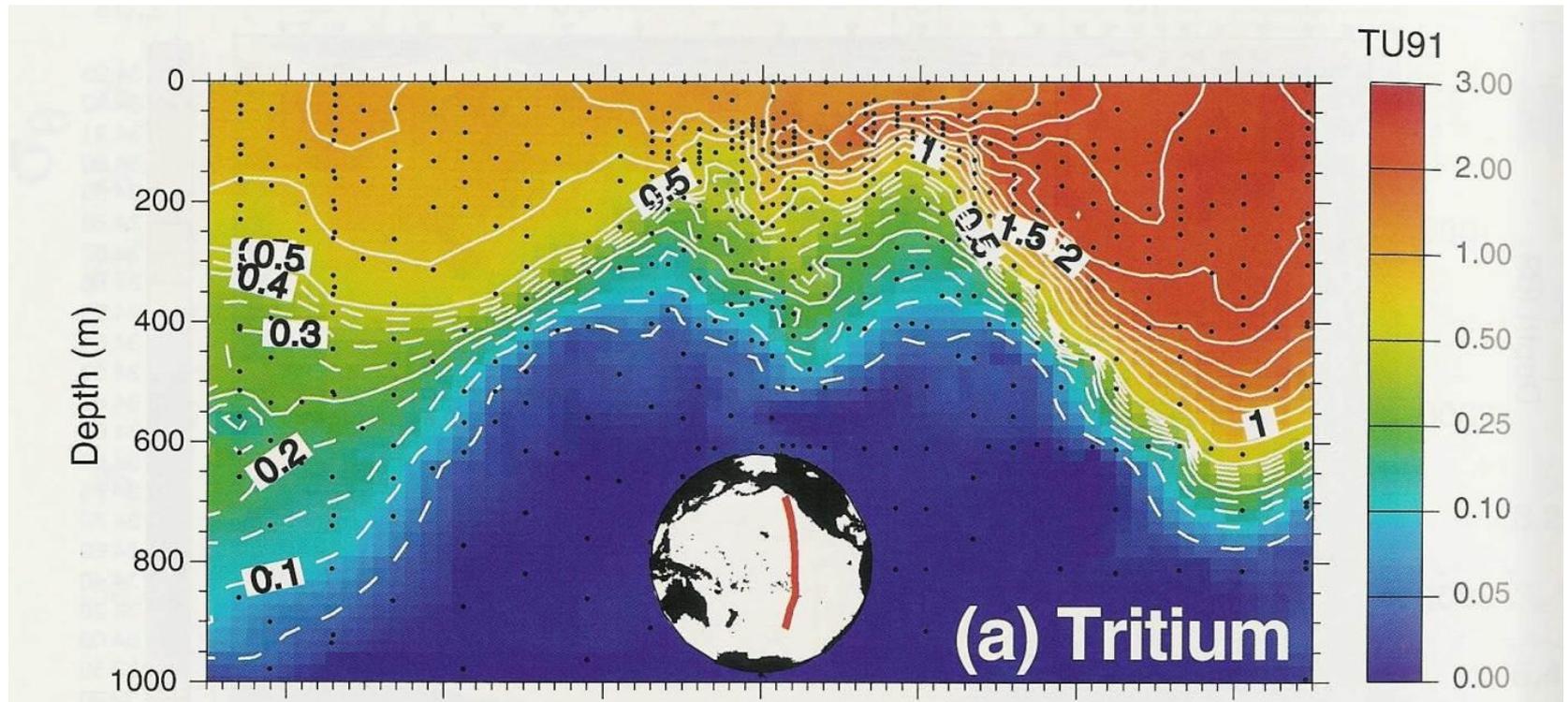
Silicate [$\mu\text{mol/kg}$]



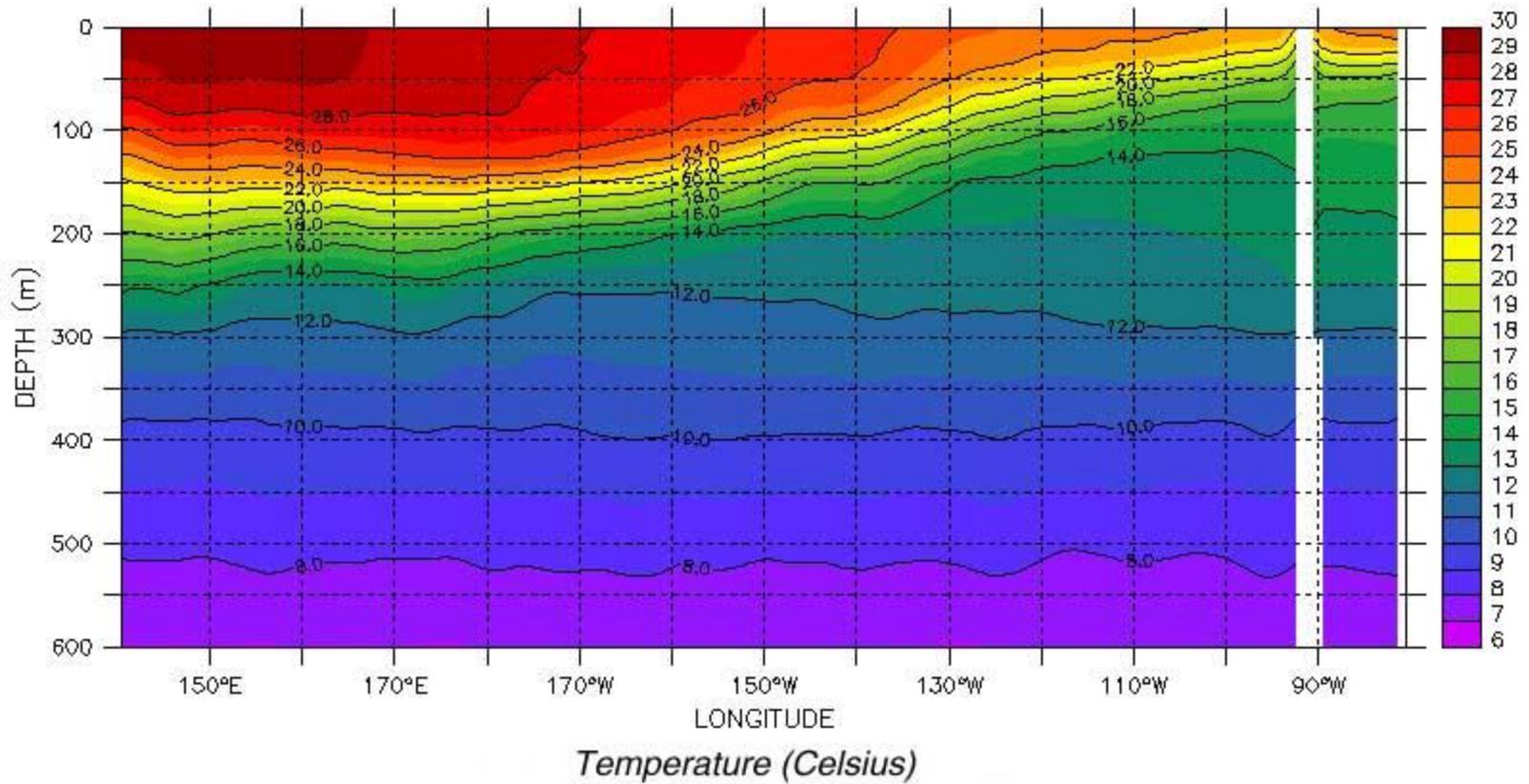
Silicate [$\mu\text{mol/kg}$]



Distribuição de Trítio no Pacífico



Secção transversal no Pacífico Equatorial: termoclina mais rasa na margem leste do oceano



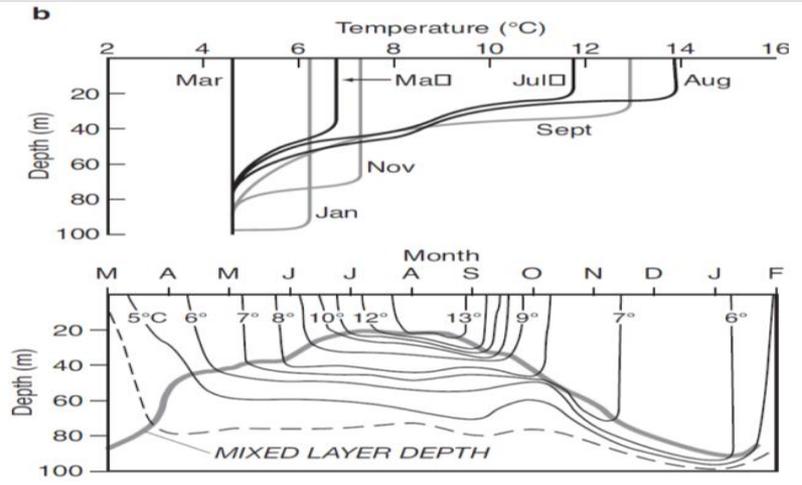
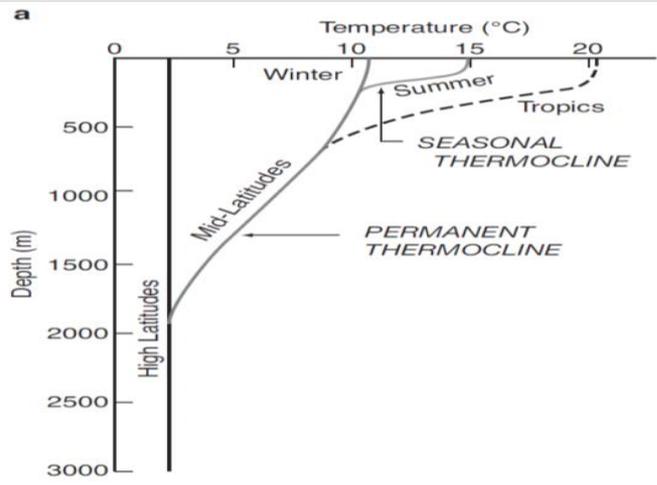
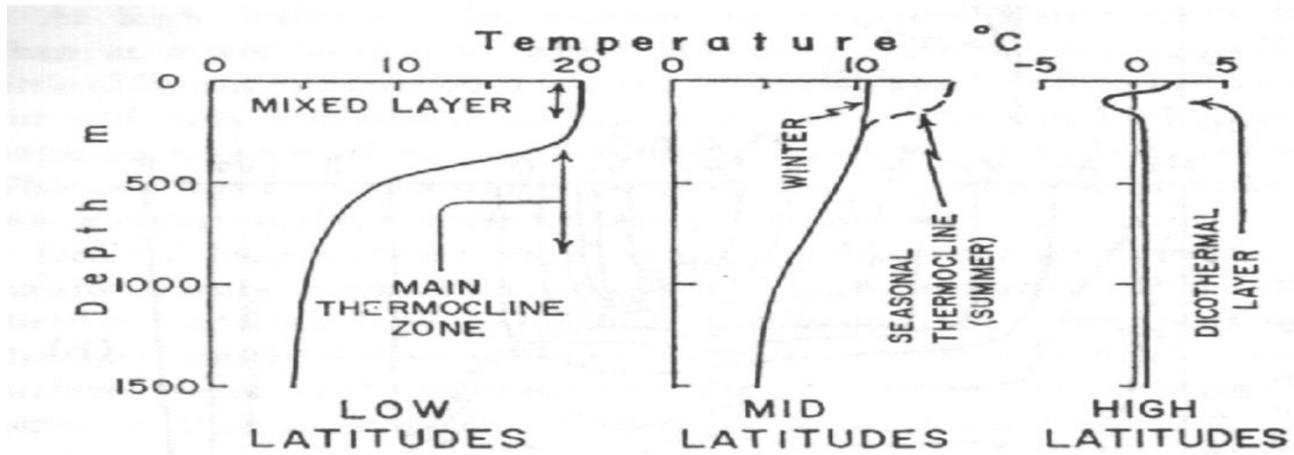
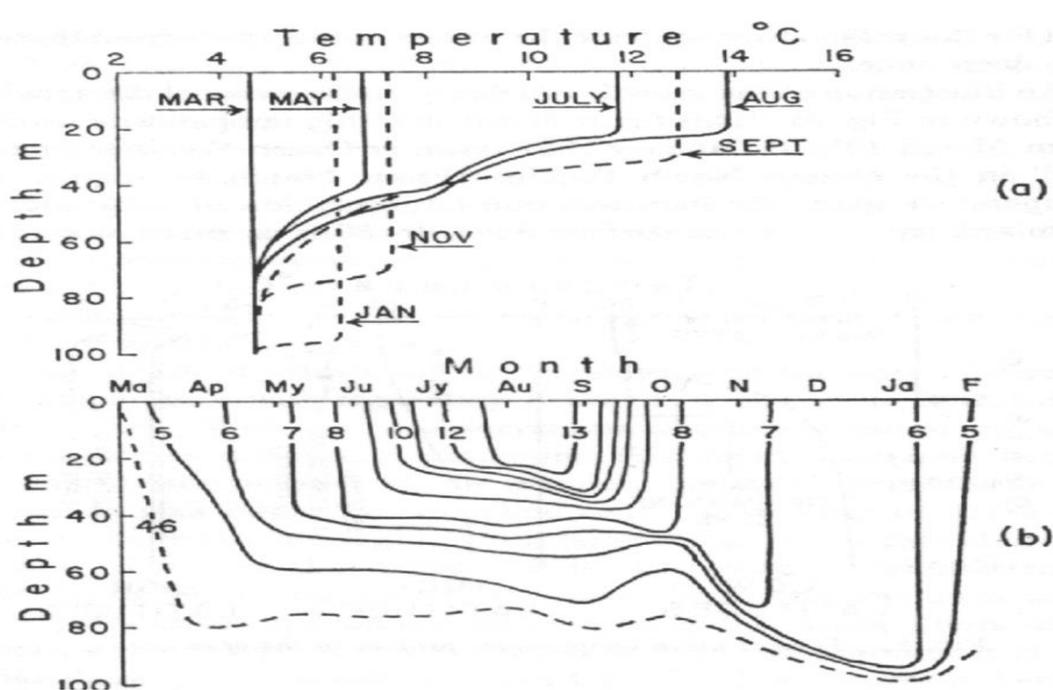


Figure 1.6: Schematic illustration of the vertical variation of temperature through the seasonal cycle (Figure from Sarmiento and Gruber, 2006).

Perfis Verticais de Temperatura



Formação da Termoclina Sazonal
Médias Latitudes



TEMPERATURA POTENCIAL?

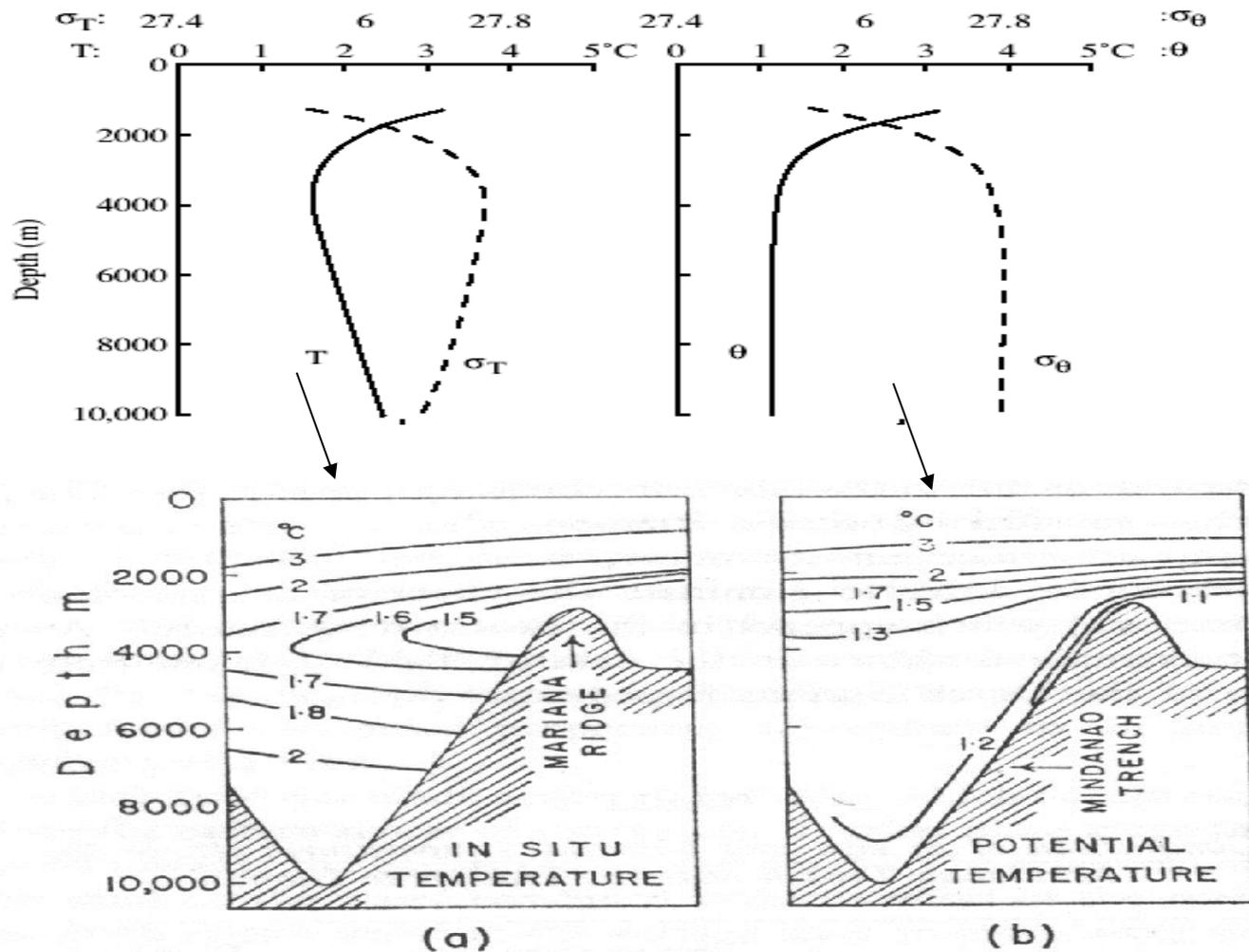
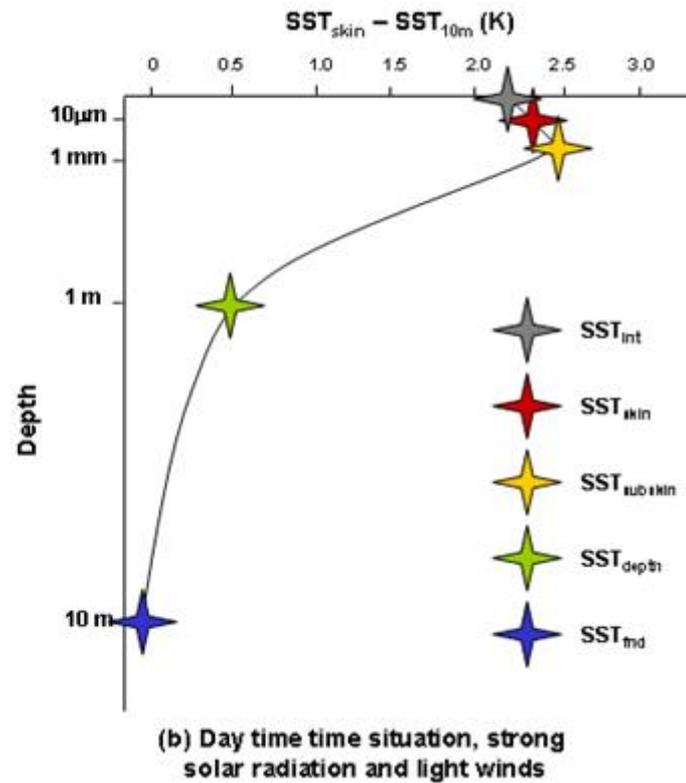
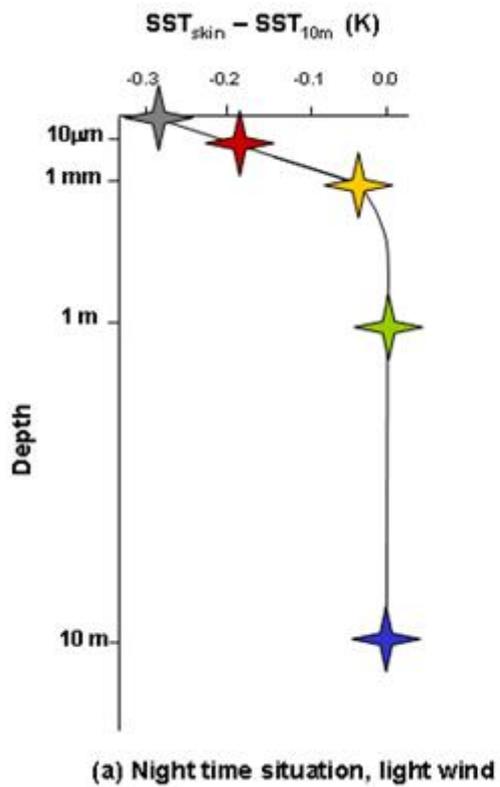
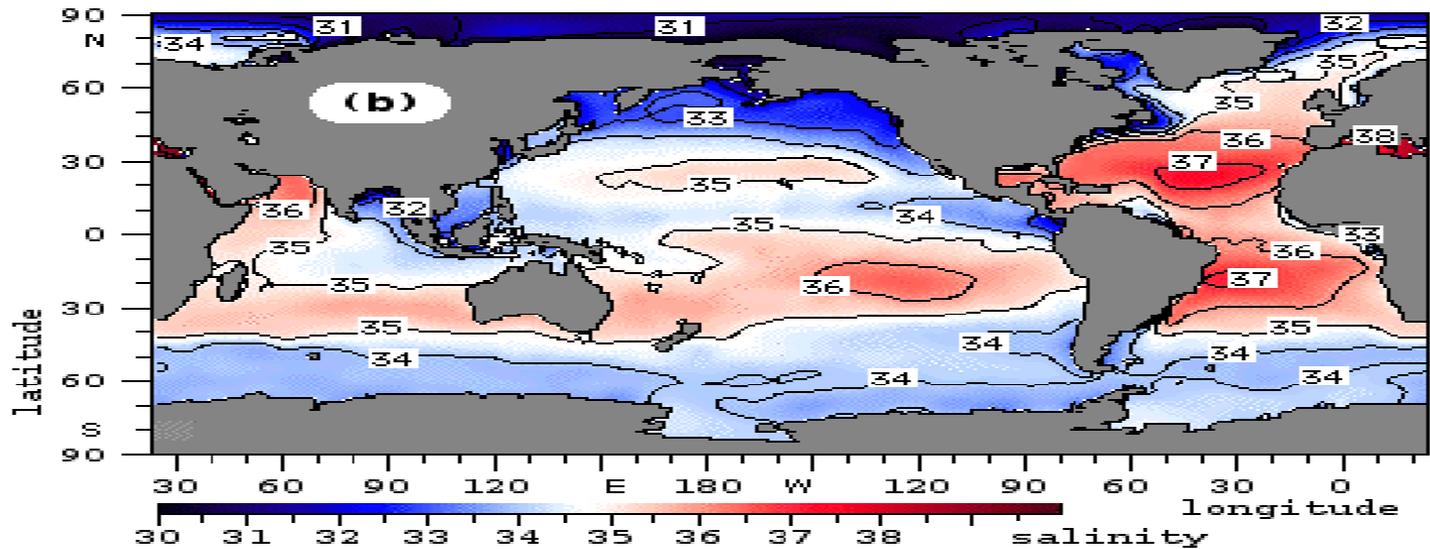
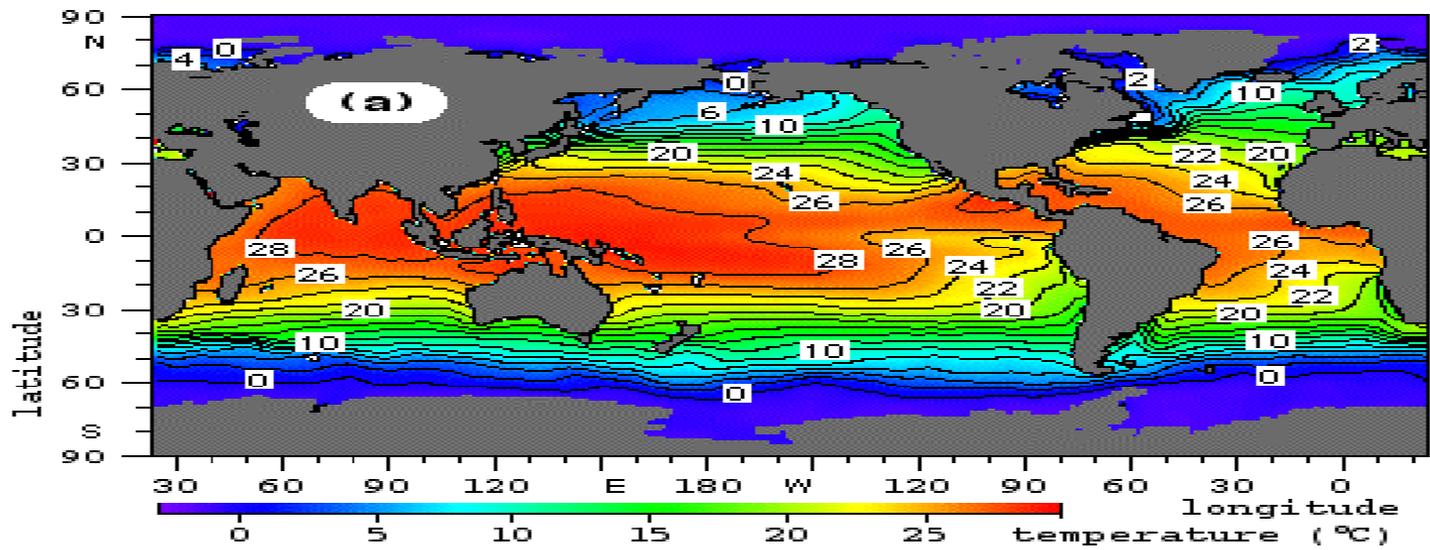


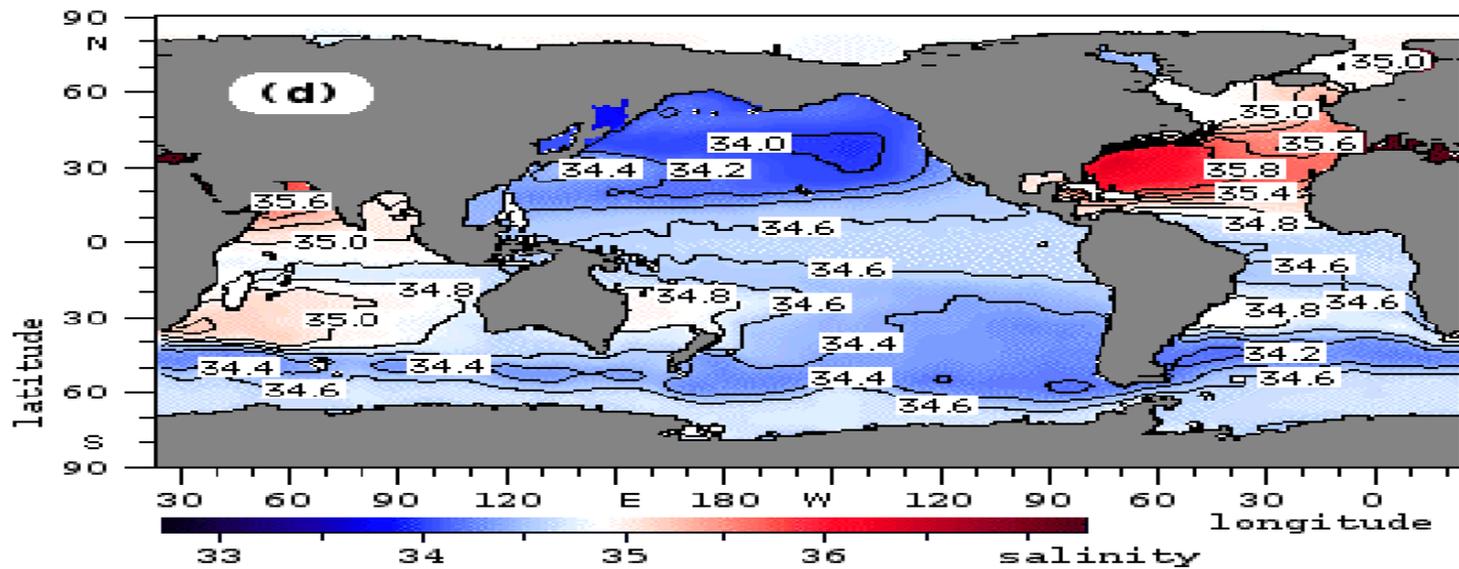
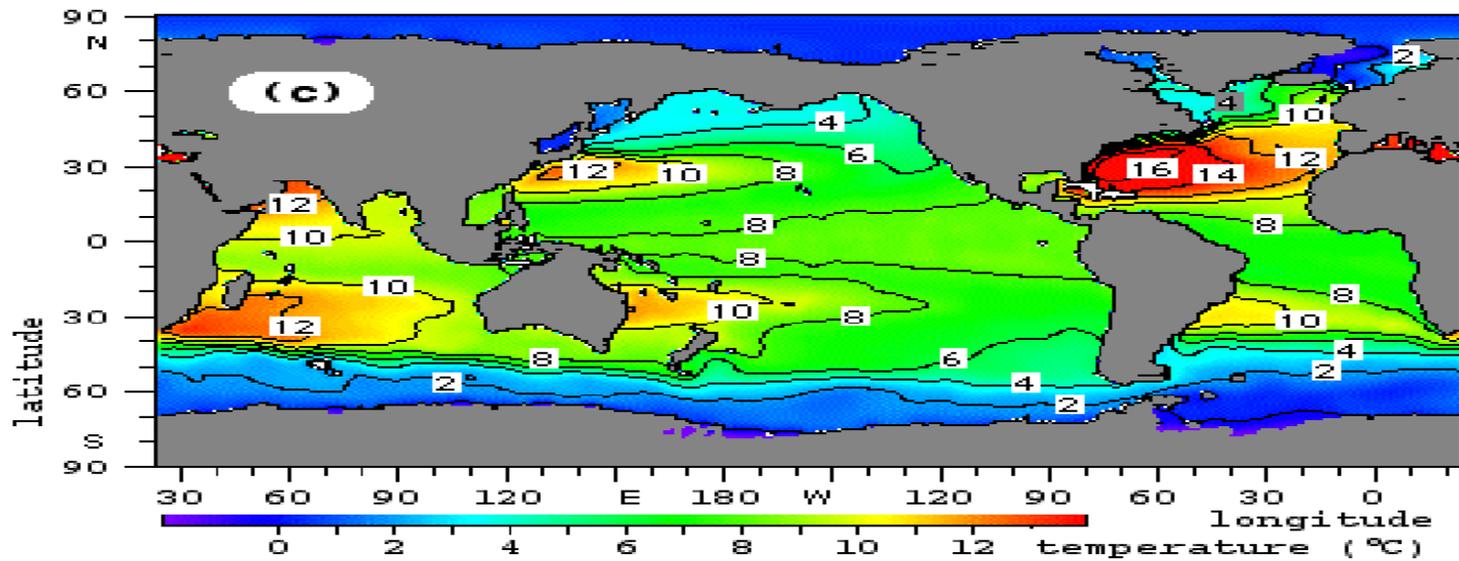
FIG. 10. In situ and potential temperature distributions in the Mindanao Trench (after van Riel).

É a T corrigida a partir da T *in situ* principalmente para grandes profundidades onde existe o efeito da pressão hidrostática.

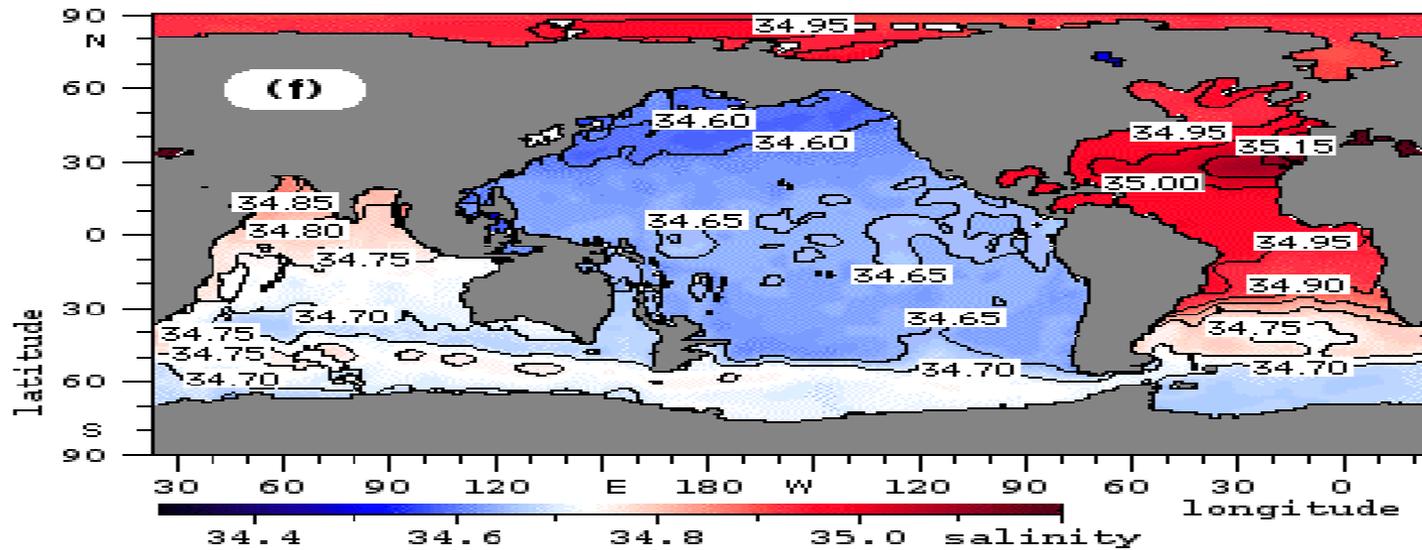
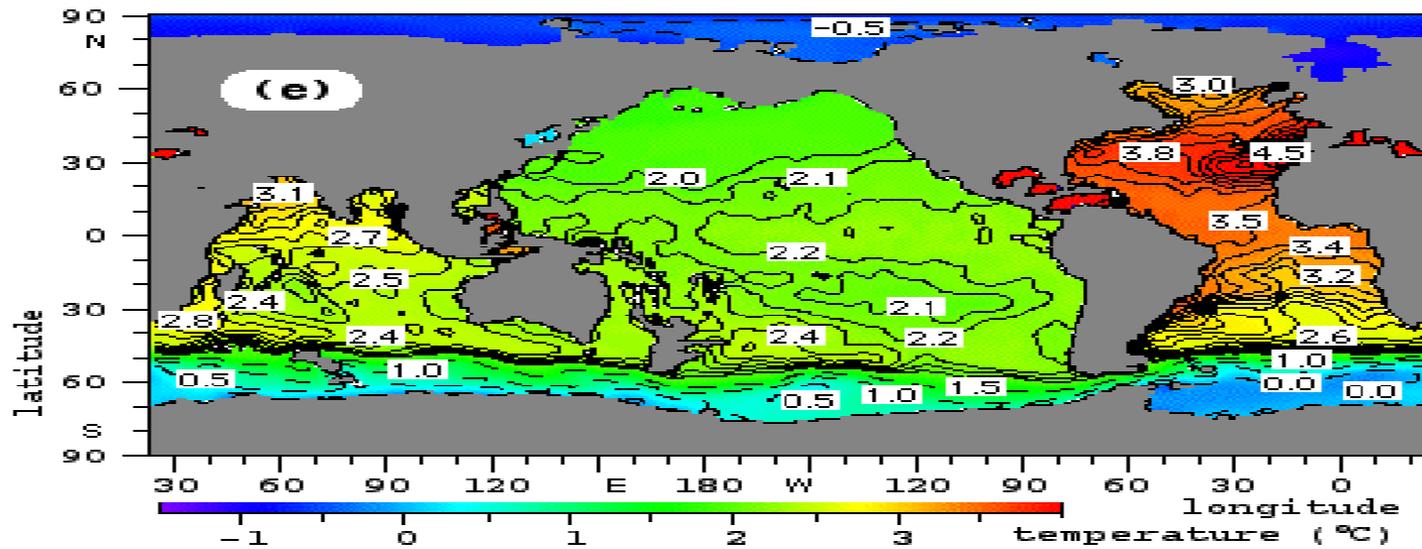




Z = 0 m

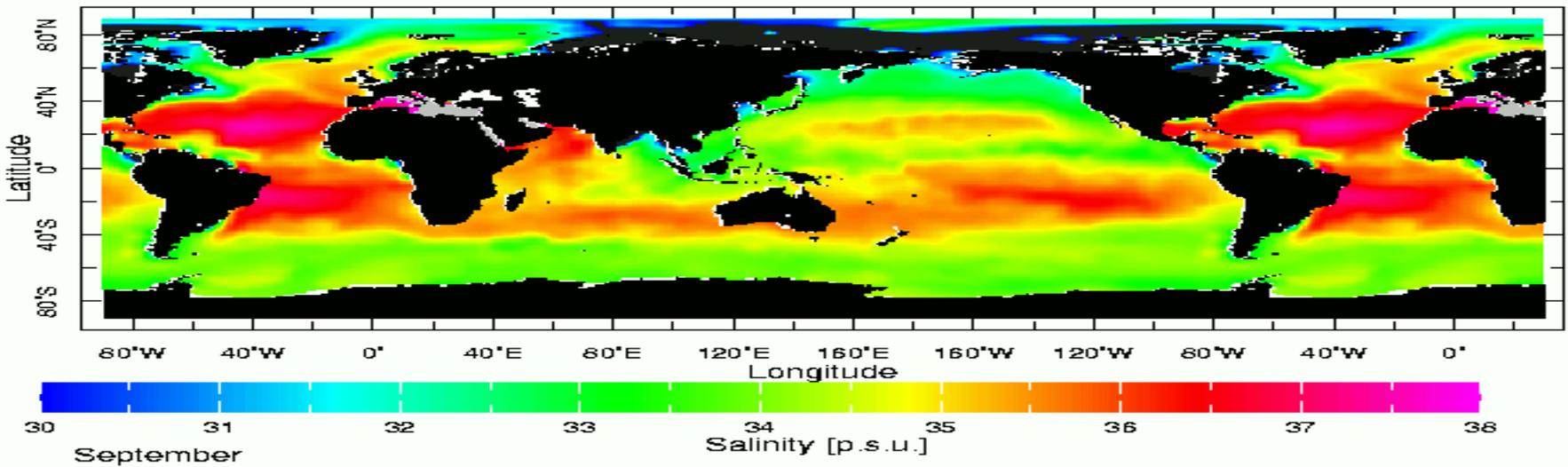
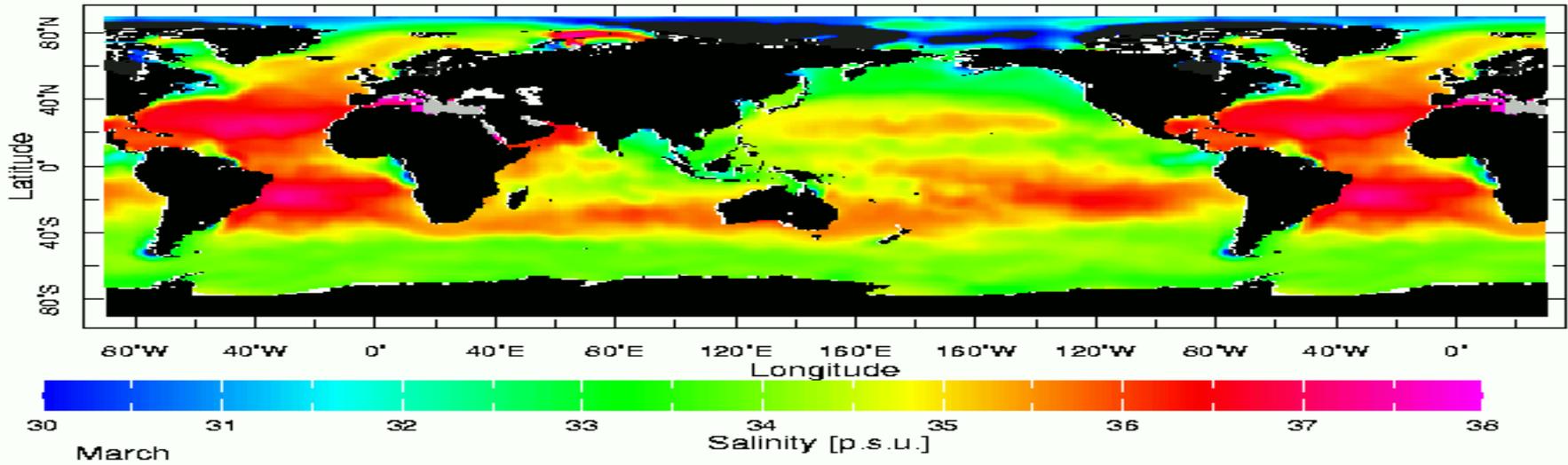


Z = 500 m



Z = 2000 m

Salinidade da Água do Mar



SALINIDADE????

É definida como o peso em gramas de material sólido dissolvido contido em 1 Kg de água do mar

qdo todos os carbonatos foram convertidos a óxidos, os brometos e iodetos em cloretos e todo material orgânico completamente oxidado.

MÉTODOS DE DETERMINAÇÃO

Evaporação [mat orgânica]

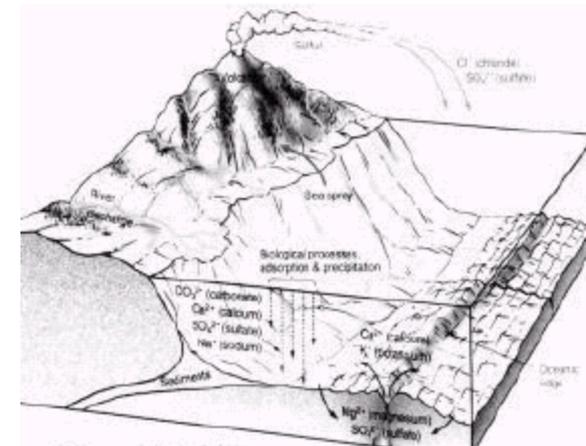
Titulação do íon cloro presente na água (clorinidade)

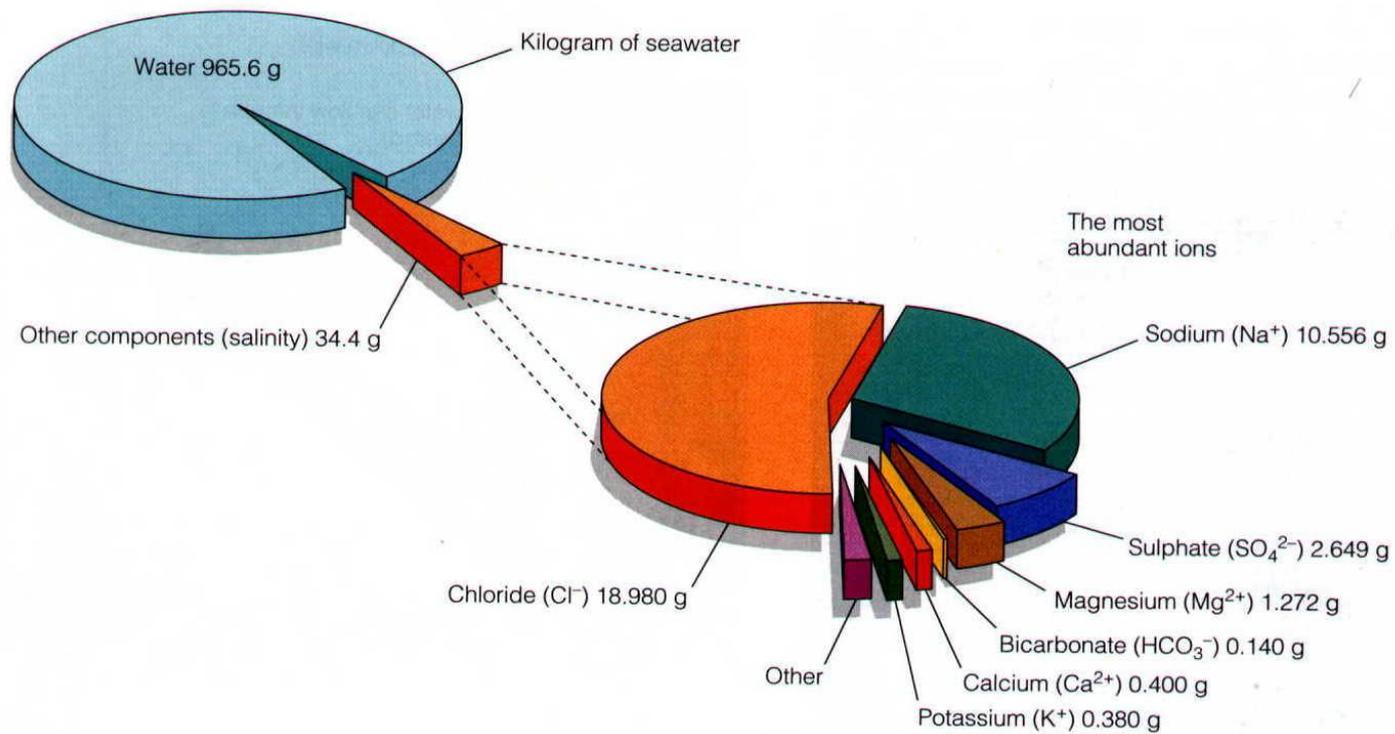
Atualmente condutividade elétrica

Salinidade = 1,80655 Cl.

adicionados pelas erupções vulcânicas e descarga dos rios ao longo do tempo geológico.

1. Cl⁻ com 55% do total de material dissolvido,
2. Na⁺ com 30,6%, sendo que o sal NaCl soma 86%,
3. SO₄ - com 7,7%,
4. Mg⁺⁺ com 3,7%,
5. K⁺ com 1,1%.





Como determinar a salinidade ?

CLORINIDADE

$$S = 1.806\,55\,Cl$$

where *chlorinity* Cl is defined as “the mass of silver required to precipitate completely the halogens in 0.328 523 4 kg of the sea-water sample.”

$$S = -0.089\,96 + 28.297\,29\,R_{15} + 12.808\,32\,R_{15}^2 - 10.678\,69\,R_{15}^3 + 5.986\,24\,R_{15}^4 - 1.323\,11\,R_{15}^5 \quad (6.3a)$$

$$R_{15} = C(S, 15, 0)/C(35, 15, 0) \quad (6.3b)$$

where $C(S, 15, 0)$ is the conductivity of the sea-water sample at 15°C and atmospheric pressure, having a salinity S derived from (6.4), and $C(35, 15, 0)$ is the conductivity of standard “Copenhagen” sea water. Millero (1996) points out that (6.3) is not a new definition of salinity, it merely gives chlorinity as a function of conductivity of seawater relative to standard seawater.

Practical Salinity Scale of 1978 By the early 1970s, accurate conductivity meters could be deployed from ships to measure conductivity at depth. The need to re-evaluate the salinity scale led the Joint Panel to recommend in 1978 that salinity be defined using only conductivity, breaking the link with chlorinity. All water samples with the same conductivity ratio have the same salinity.

The *Practical Salinity Scale of 1978* is now the official definition:

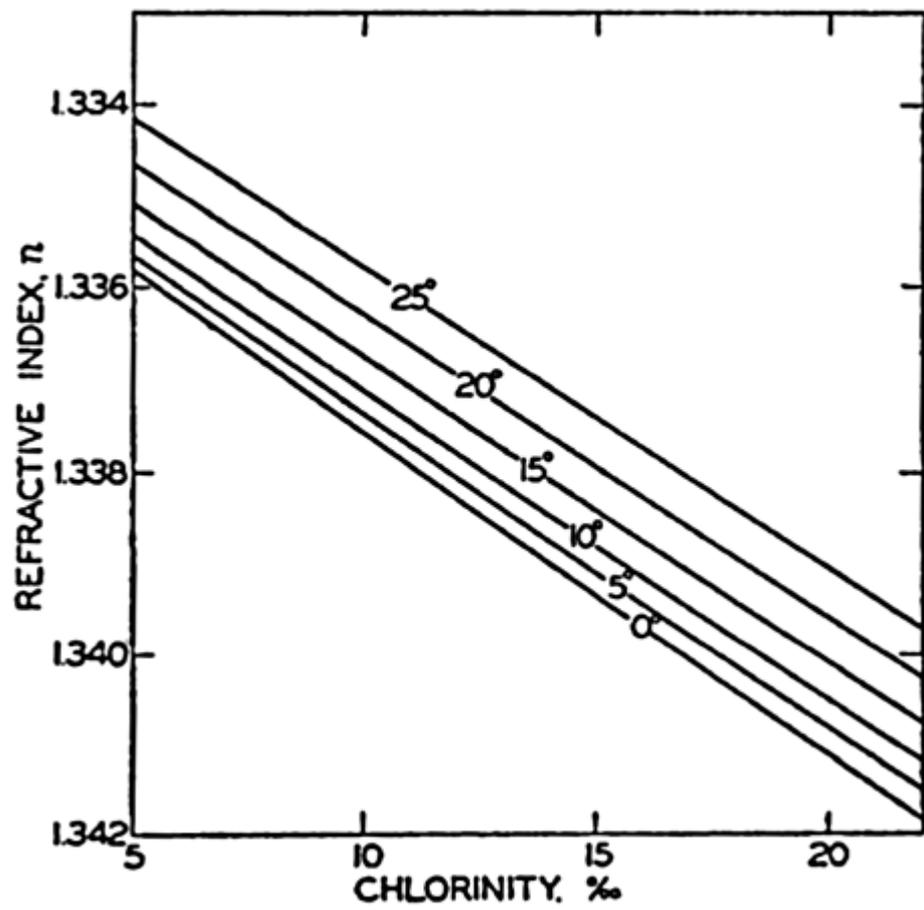
$$S = 0.0080 - 0.1692 R_T^{1/2} + 25.3851 R_T + 14.0941 R_T^{3/2} - 7.0261 R_T^2 + 2.7081 R_T^{5/2} + \Delta S \quad (6.4a)$$

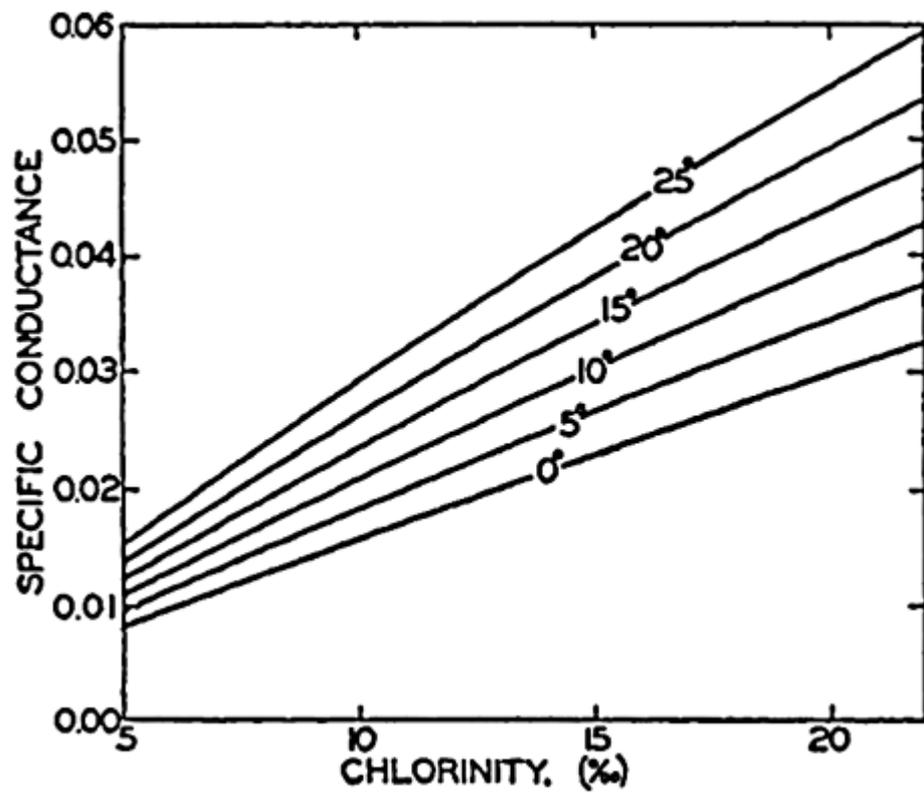
$$R_T = C(S, T, 0) / C(KCl, T, 0) \quad (6.4b)$$

$$\Delta S = \left[\frac{(T - 15)}{1 + 0.0162 (T - 15)} \right] + 0.0005 - 0.0056 R_T^{1/2} - 0.0066 R_T - 0.0375 R_T^{3/2} + 0.636 R_T^2 - 0.0144 R_T^{5/2} \quad (6.4c)$$

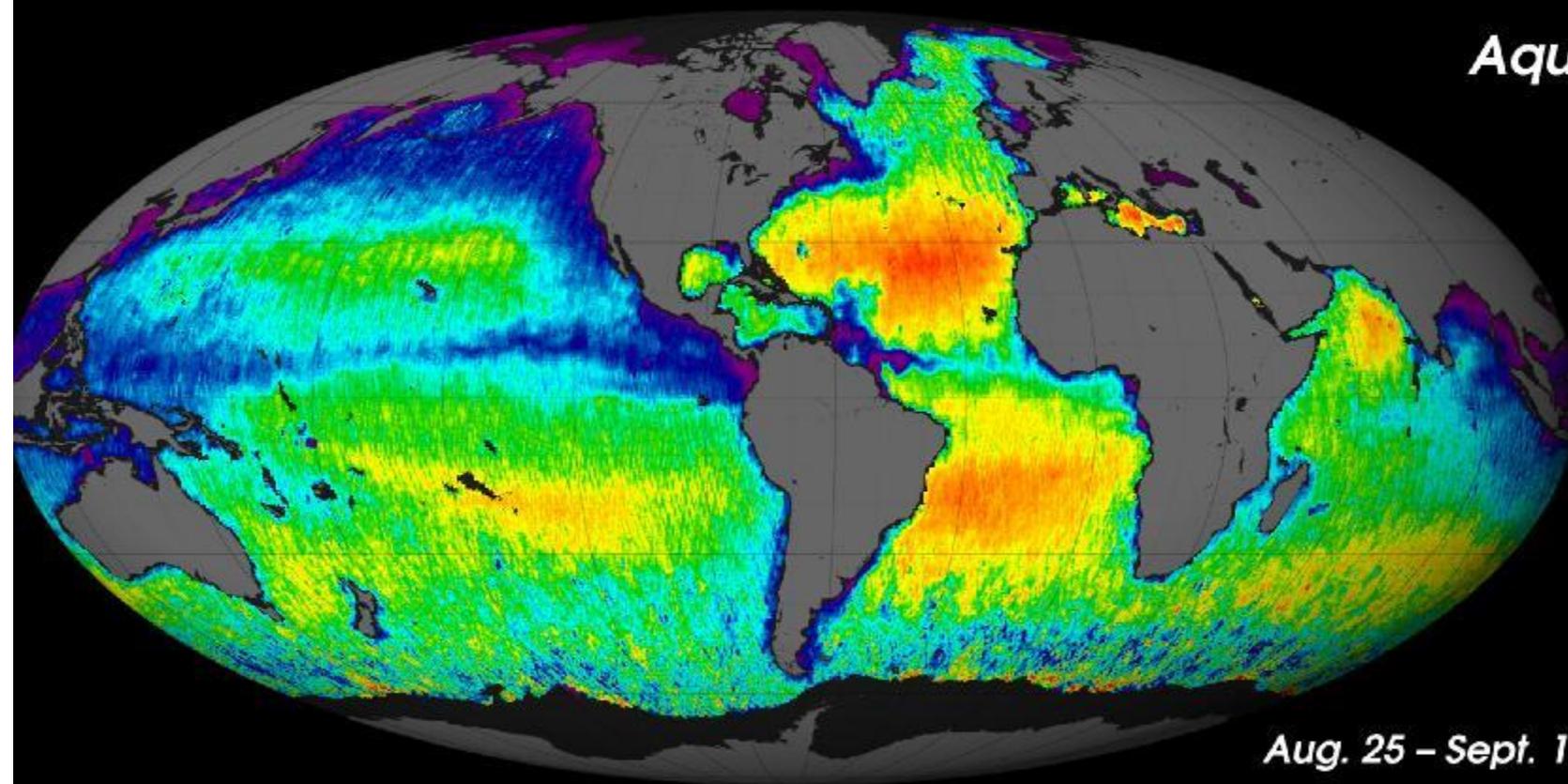
$$2 \leq S \leq 42$$

where $C(S, T, 0)$ is the conductivity of the sea-water sample at temperature T and standard atmospheric pressure, and $C(KCl, T, 0)$ is the conductivity of the standard potassium chloride (KCl) solution at temperature T and standard atmospheric pressure. The standard KCl solution contains a mass of 32.4356 grams of KCl in a mass of 1.000000 kg of solution. An extension of (6.4) gives salinity at any pressure (see Millero 1996: 72).



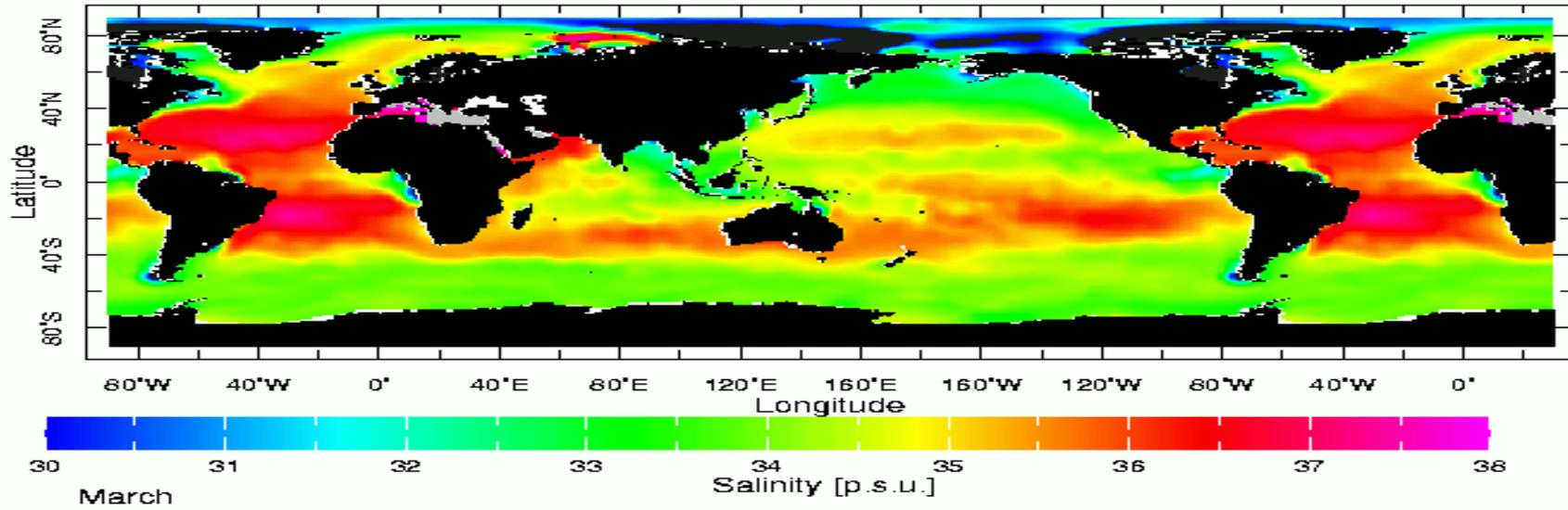


Aquarius

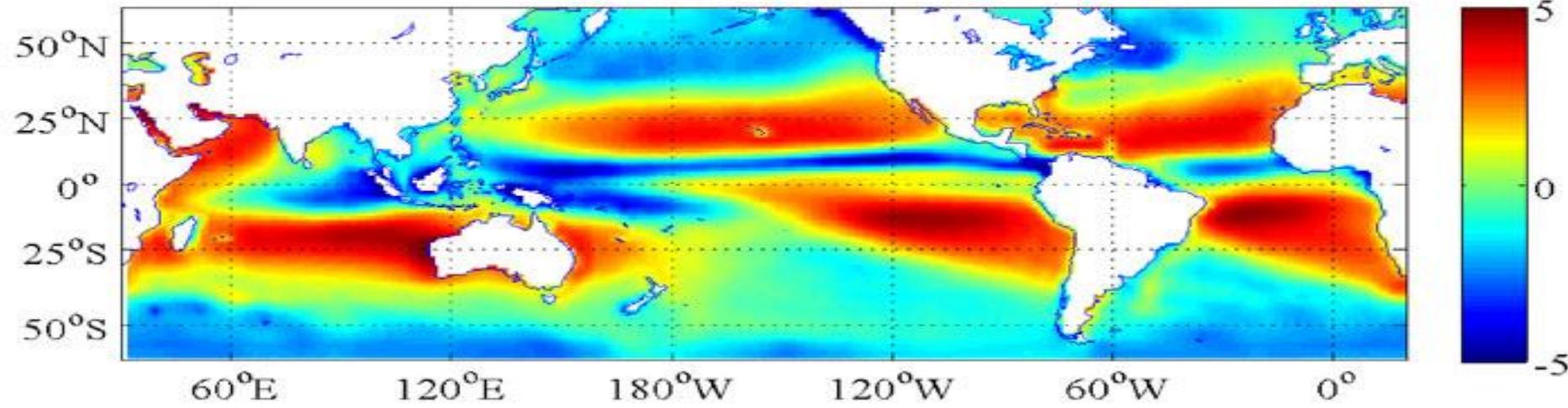


Aug. 25 - Sept. 11, 2011

Variação superficial da salinidade: o que causa isso?



Evaporation Minus Precipitation 1981--2002



Variação da Salinidade

Evaporação, precipitação, aportes continentais, congelamento/degelo

DISTRIBUIÇÃO HORIZONTAL DA SALINIDADE

aproximadamente zonal

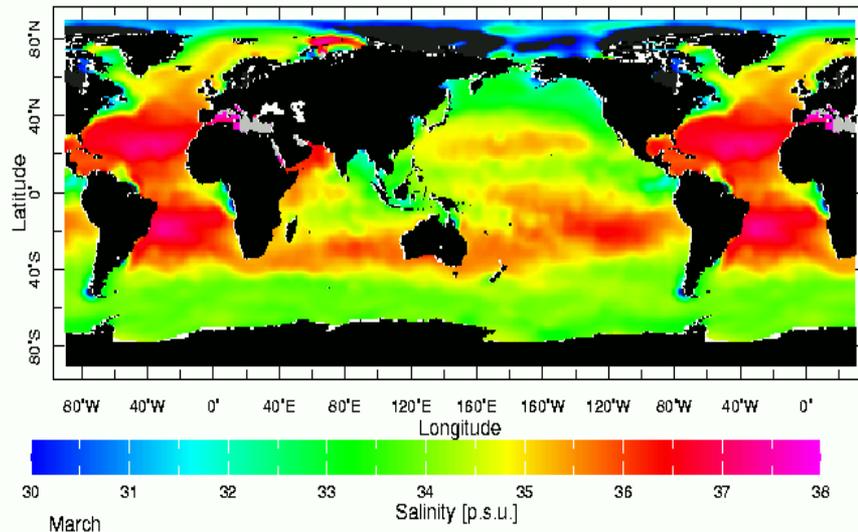
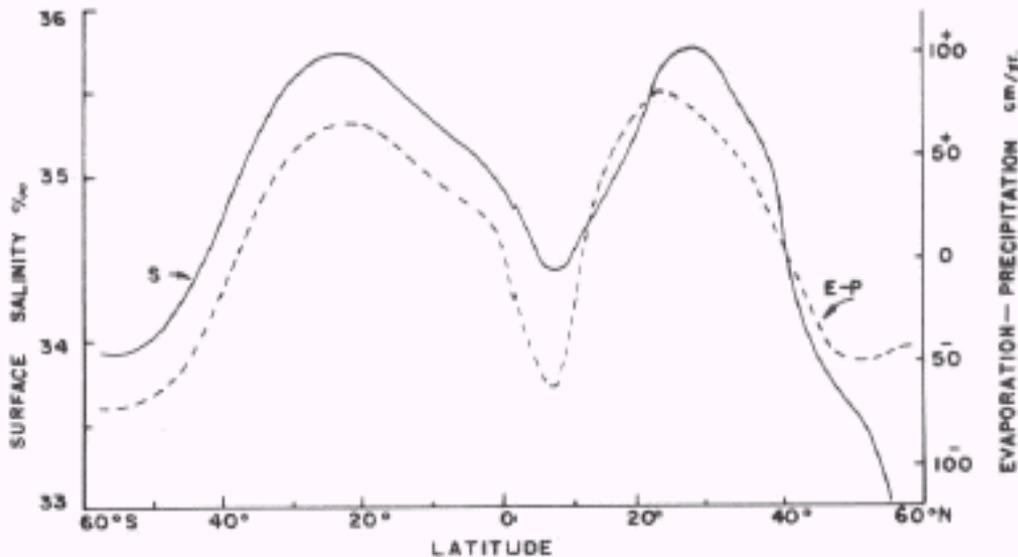
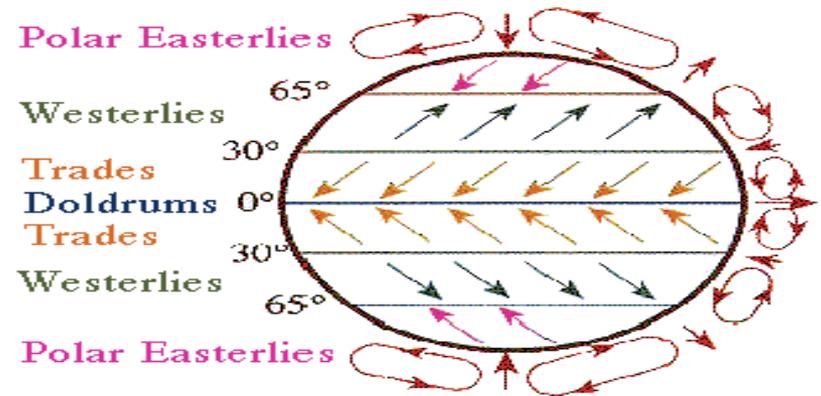
Depende do balanço E-P

Max ~20° E>P

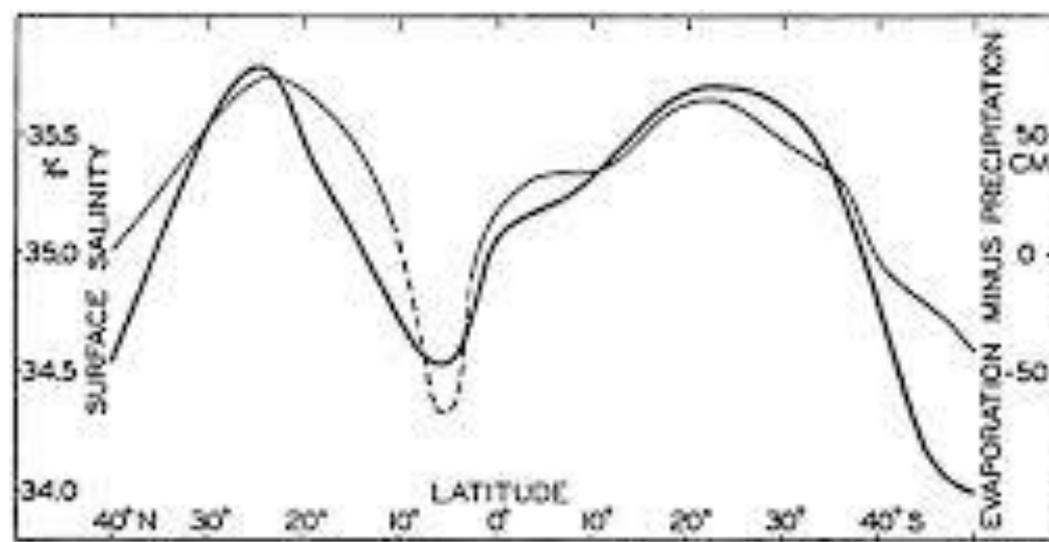
Min Equador P>E

S média sup 33.8 e 37.5

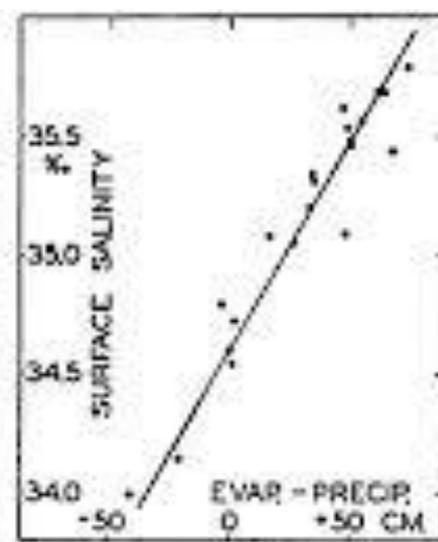
Altas latitudes: congelamento/degelo



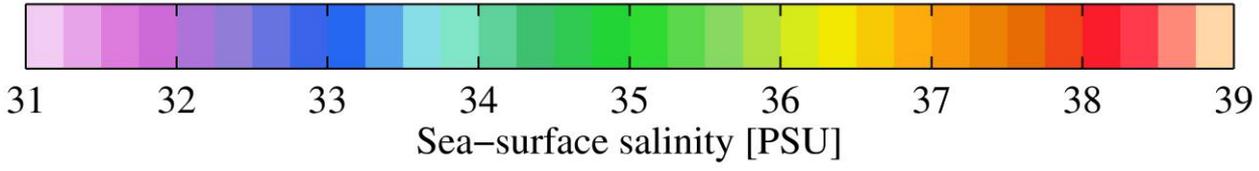
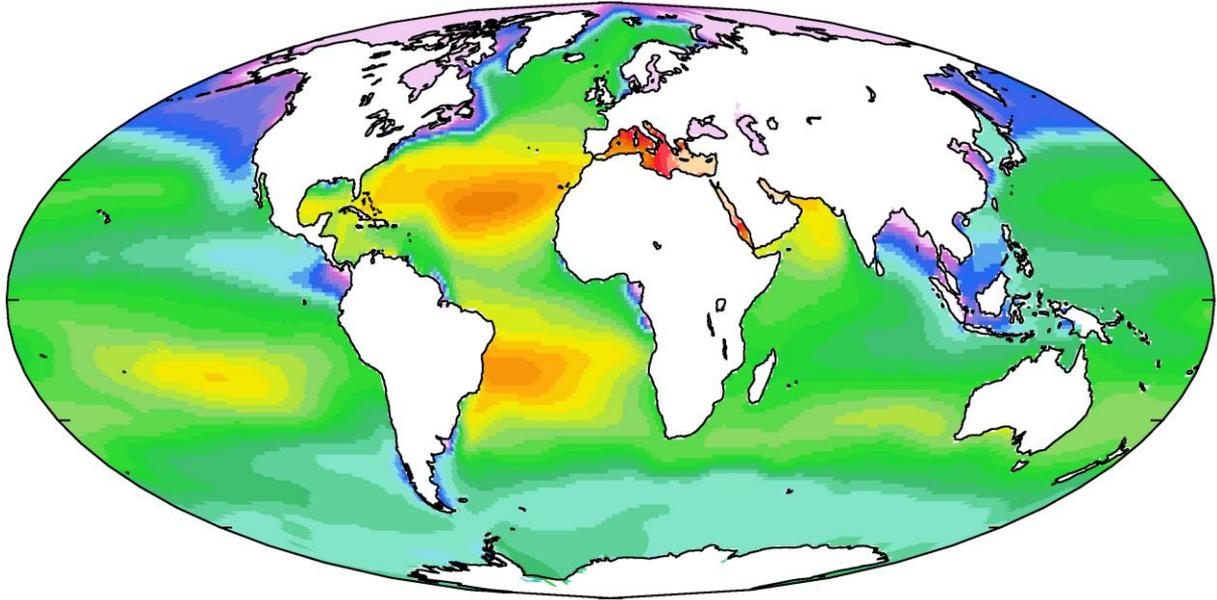
S superficial média para todos os oceanos e diferenças entre evaporação e precipitação (E-P) em função da latitude.



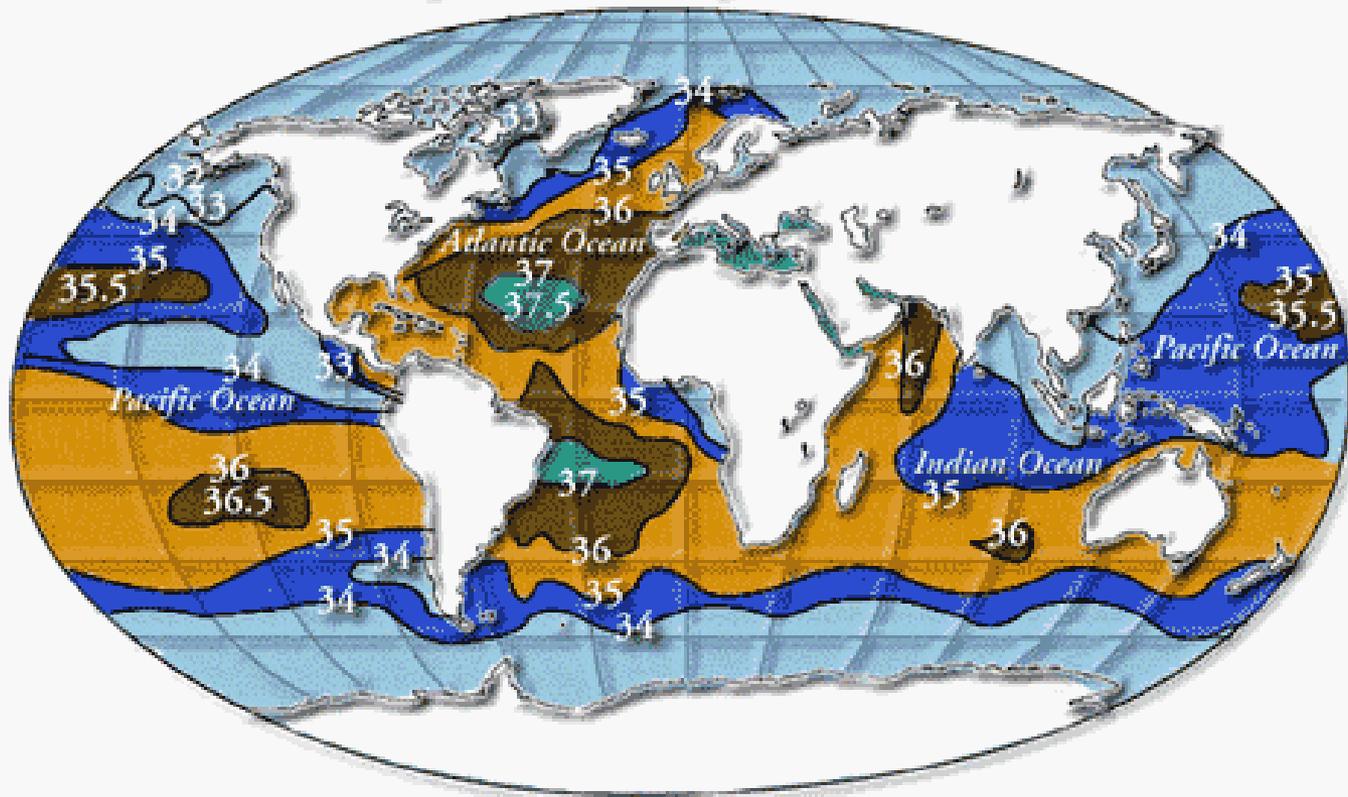
A

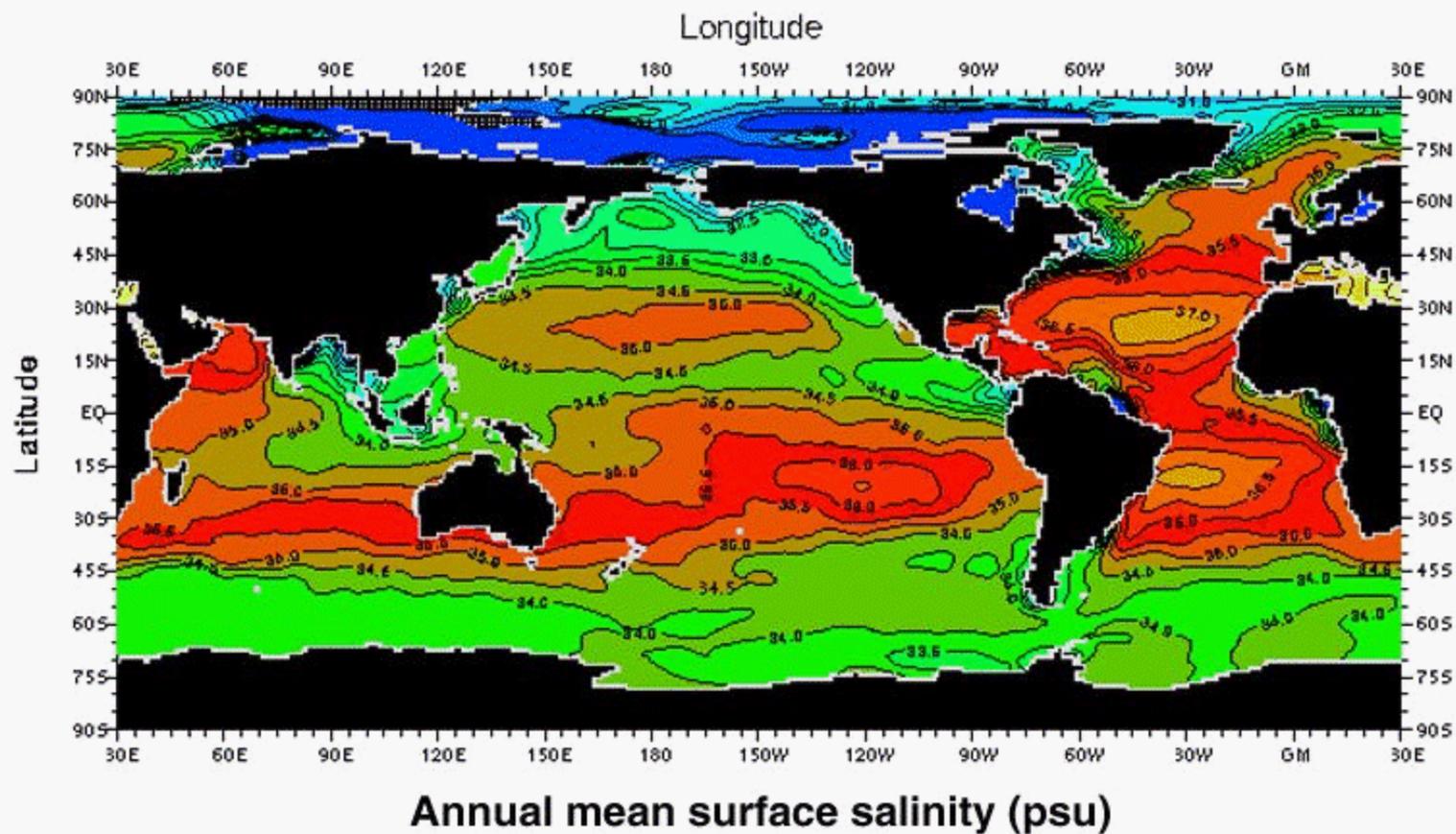


B

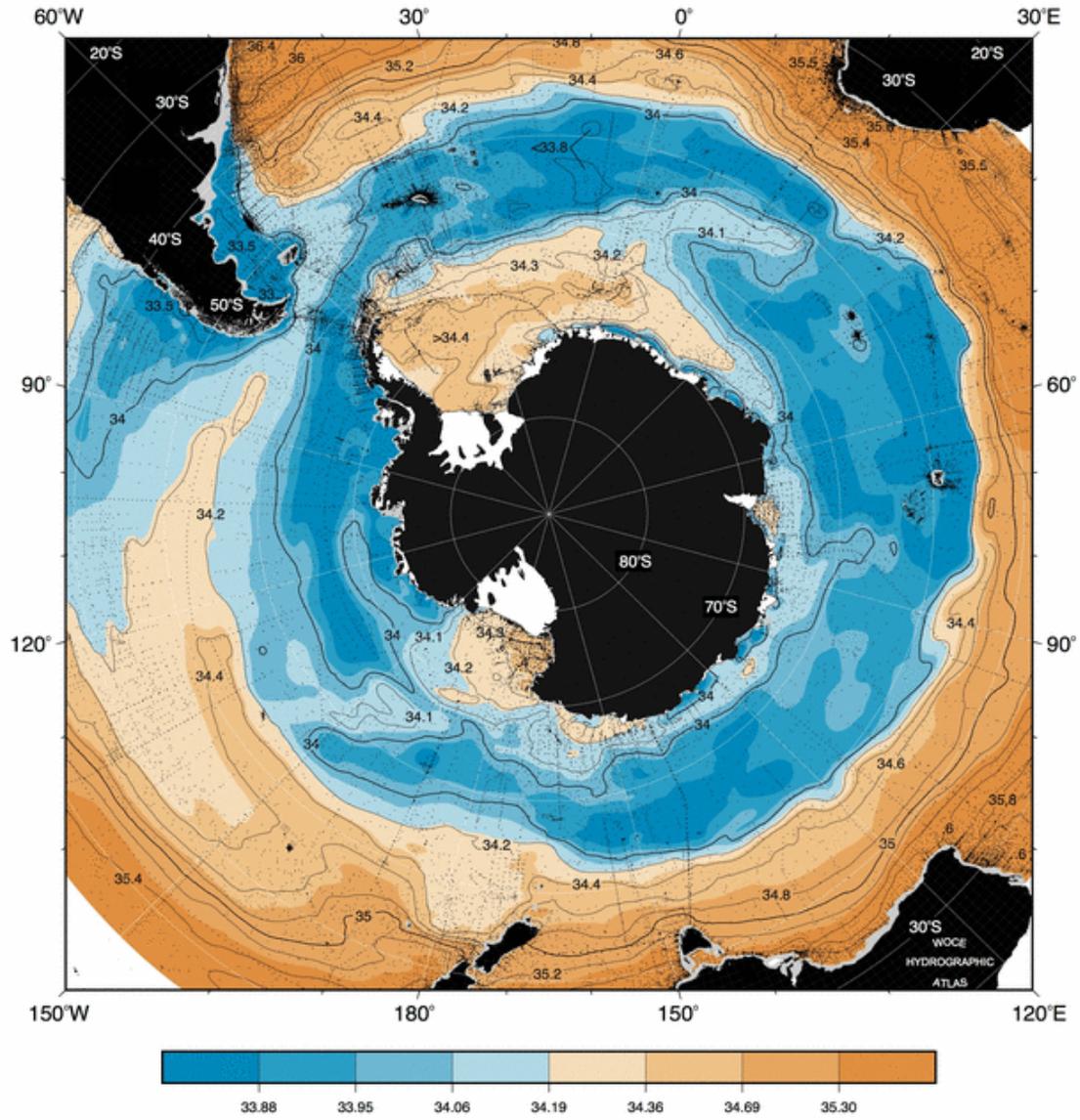


Surface Salinities of the Oceans (‰)



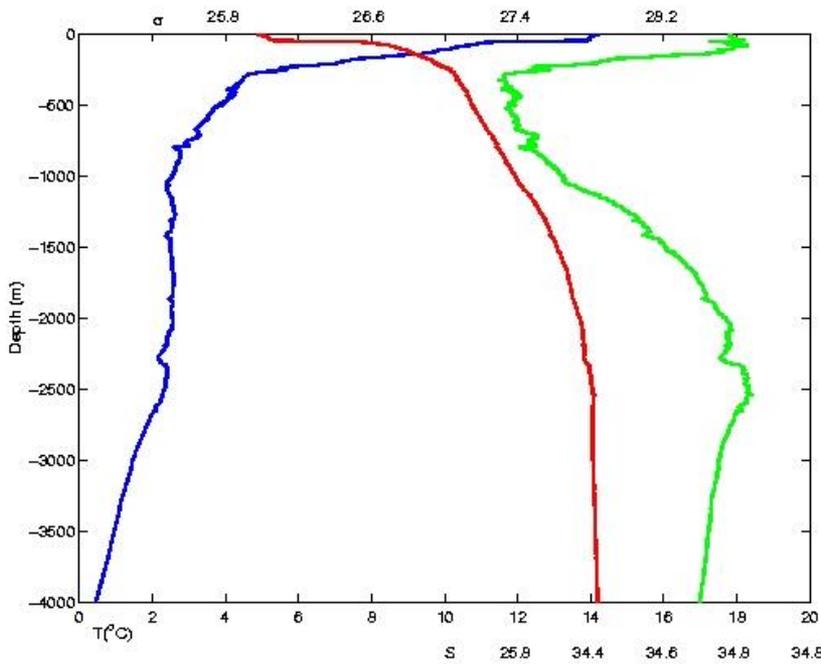
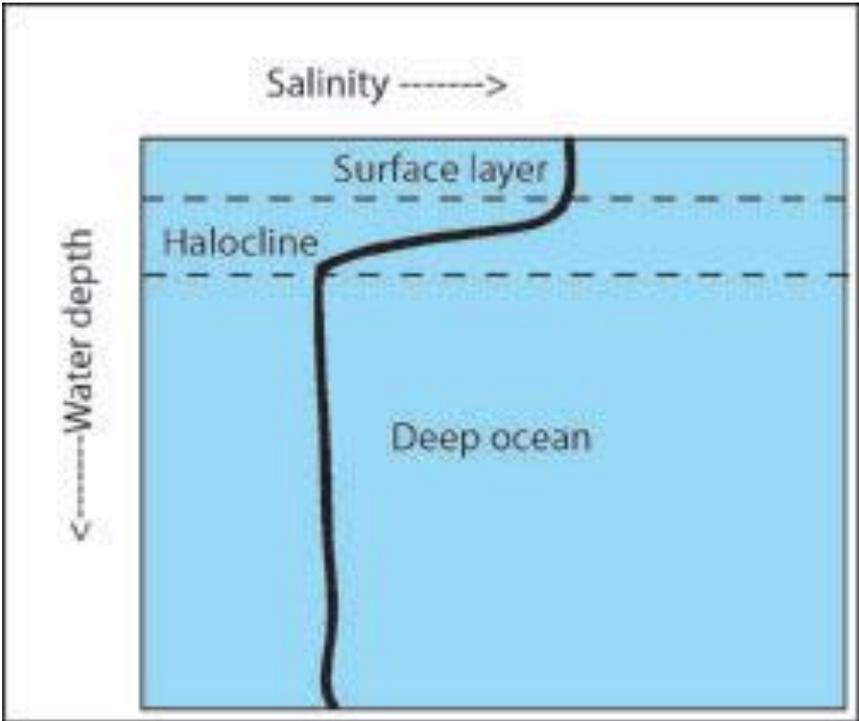


Depth: 50m

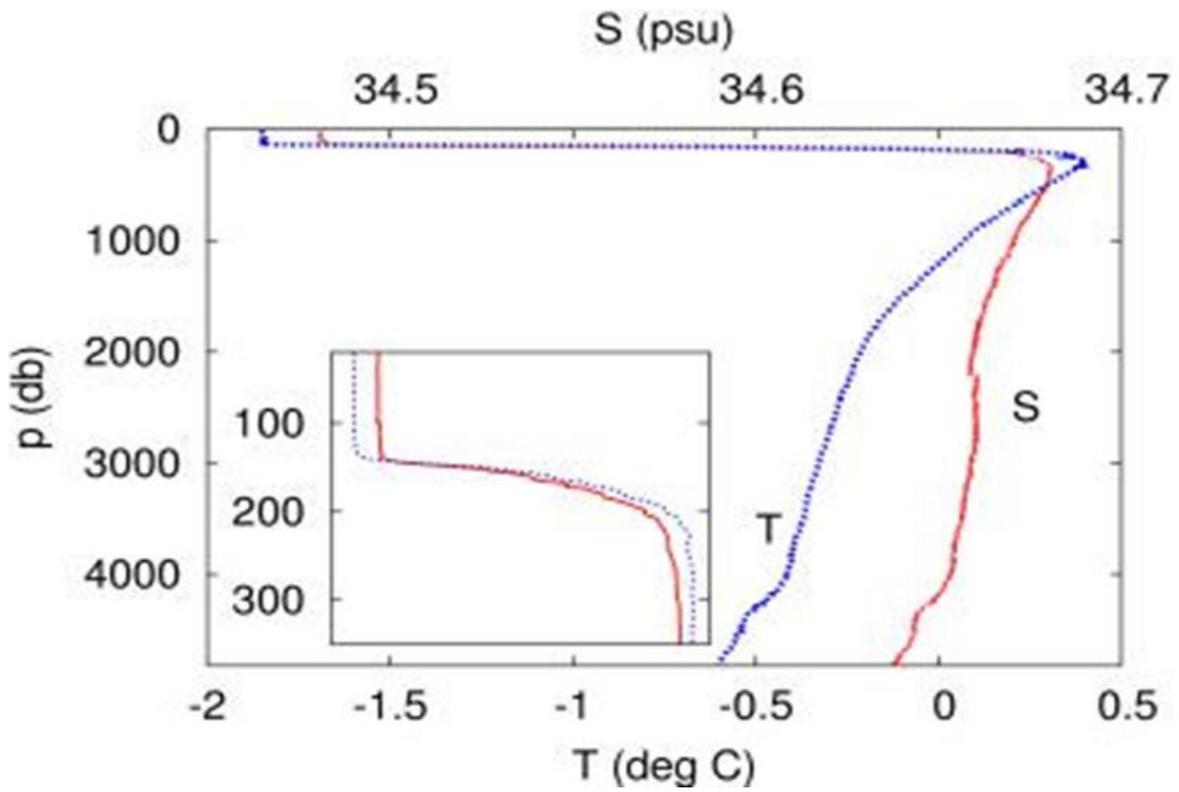


Salinity (PSS78)

DISTRIBUIÇÃO VERTICAL DA SALINIDADE ENTRE 60 S E 60 N



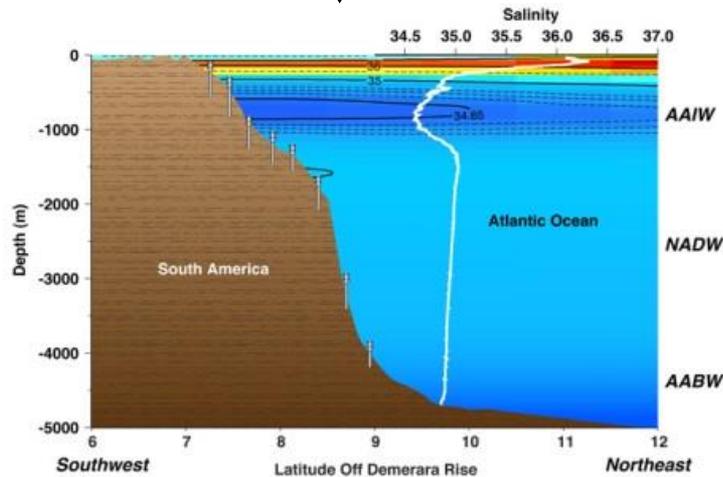
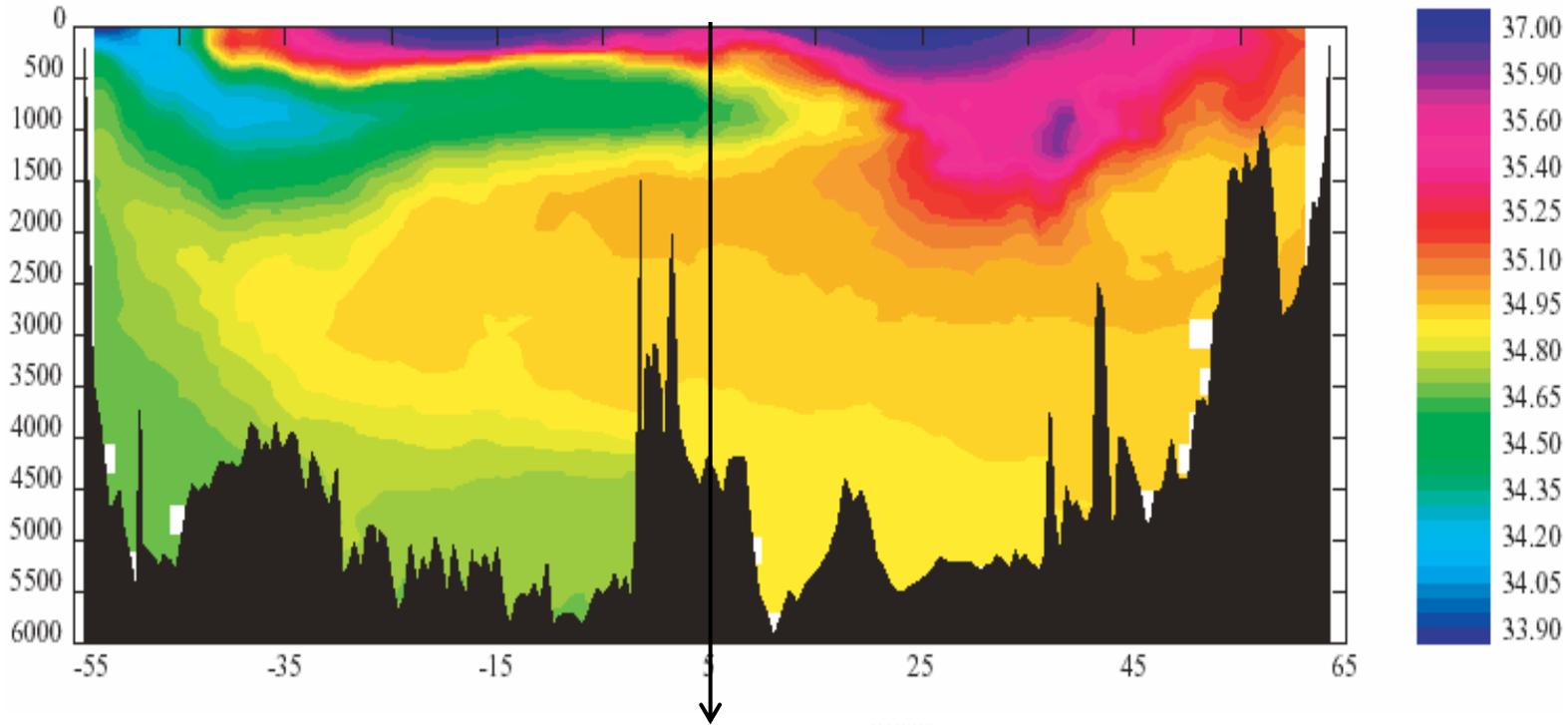
DISTRIBUIÇÃO VERTICAL DA
SALINIDADE EM ALTAS LATITUDES —
MAR DE WEDDELL



DISTRIBUIÇÃO VERTICAL DA SALINIDADE

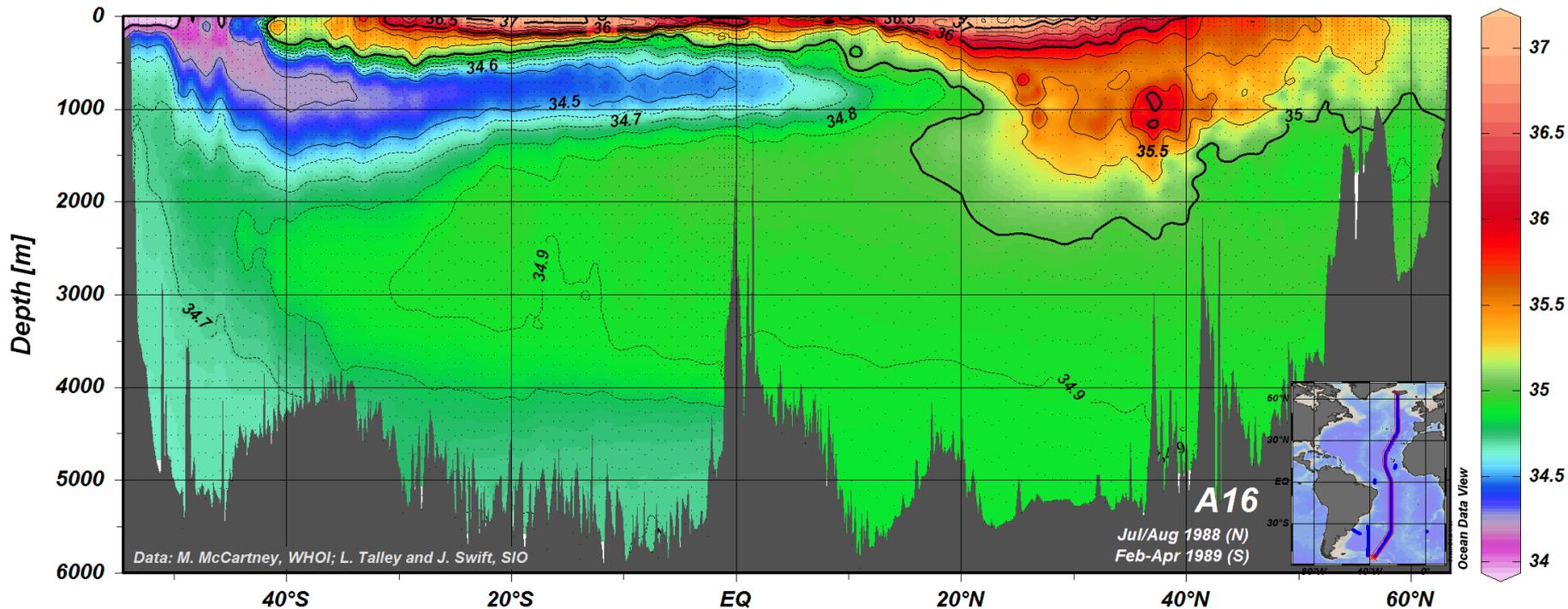
A distribuição vertical da S em baixas e médias lat é caracterizada por um mín de S entre 600 a 1000 m com um aumento a partir dos 2000 m

A16 Salinity

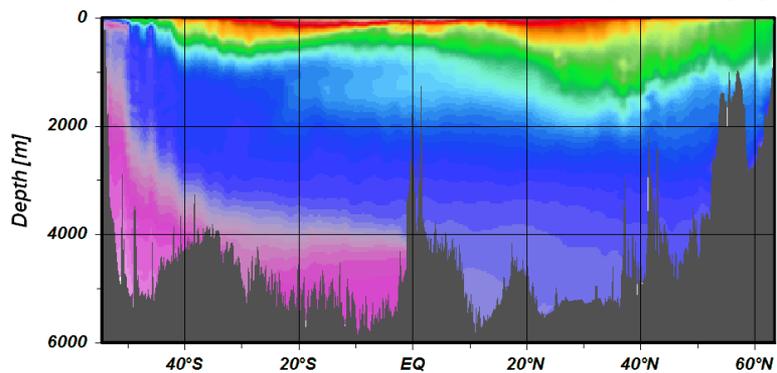


eWOCE

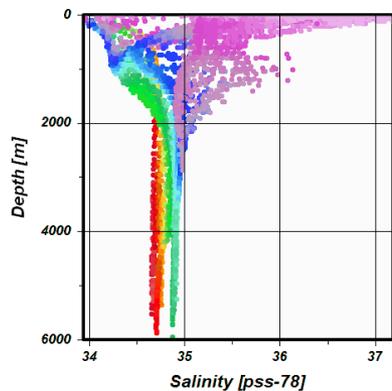
Salinity [pss-78]



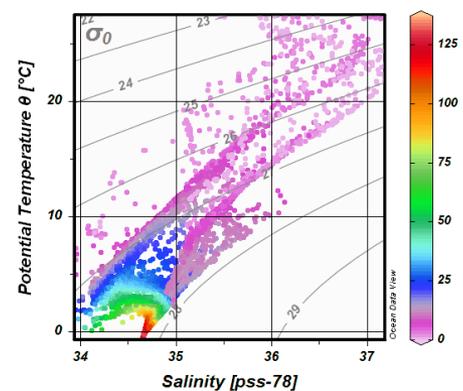
Potential Temperature θ [°C]



Silicate [$\mu\text{mol/kg}$]

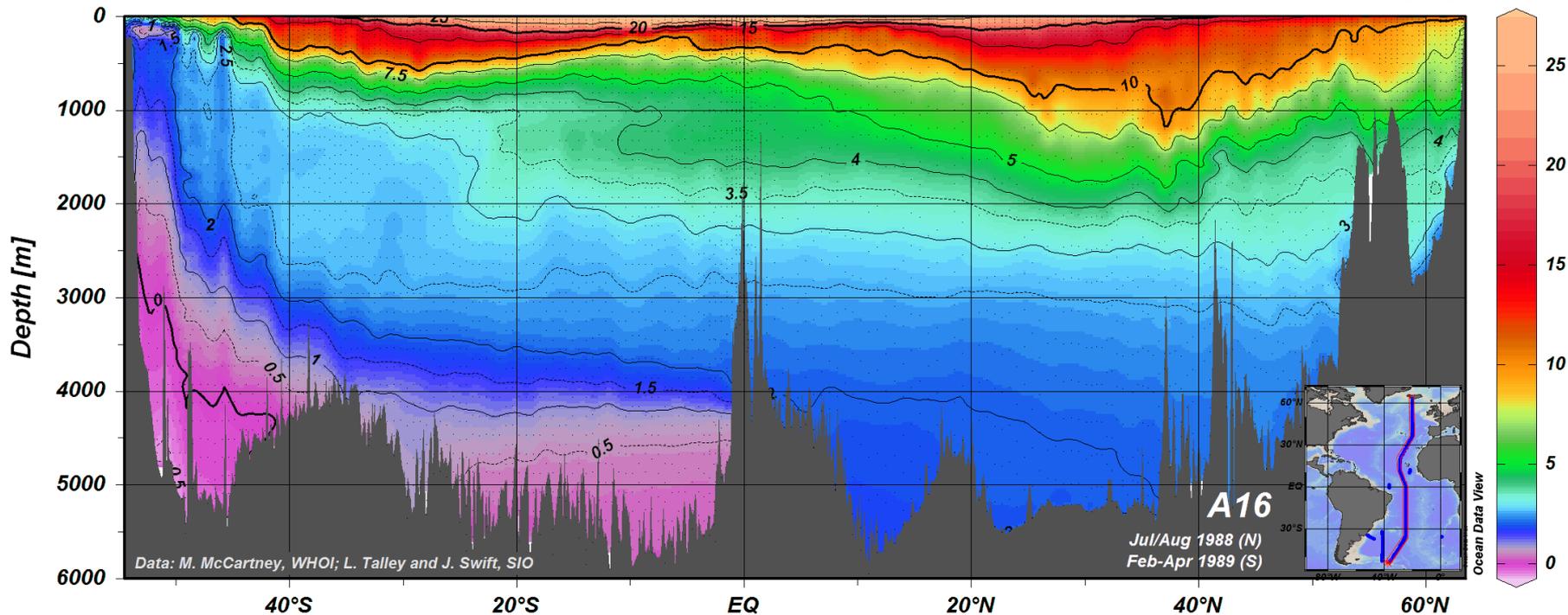


Silicate [$\mu\text{mol/kg}$]

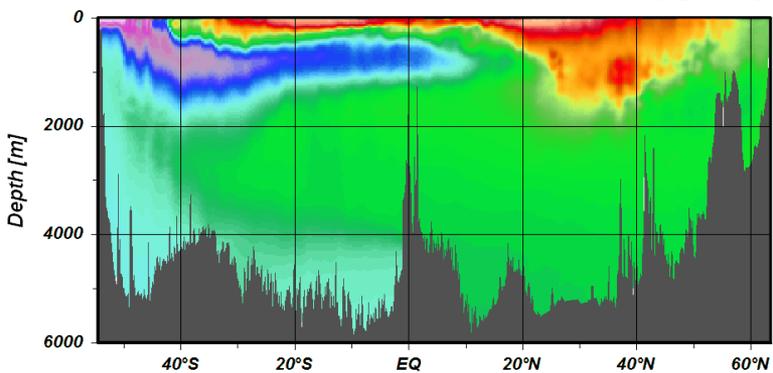


eWOCE

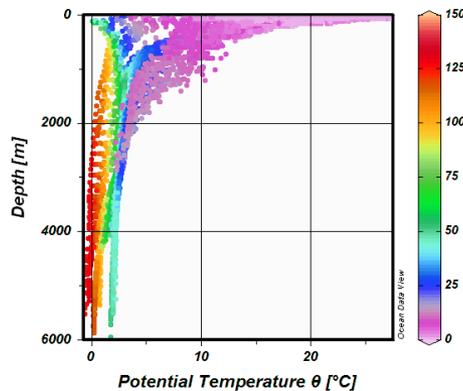
Potential Temperature θ [$^{\circ}\text{C}$]



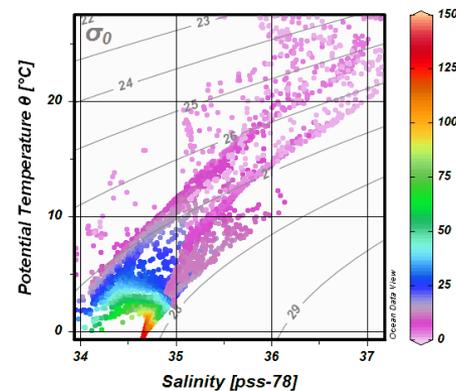
Salinity [pss-78]



Silicate [$\mu\text{mol/kg}$]

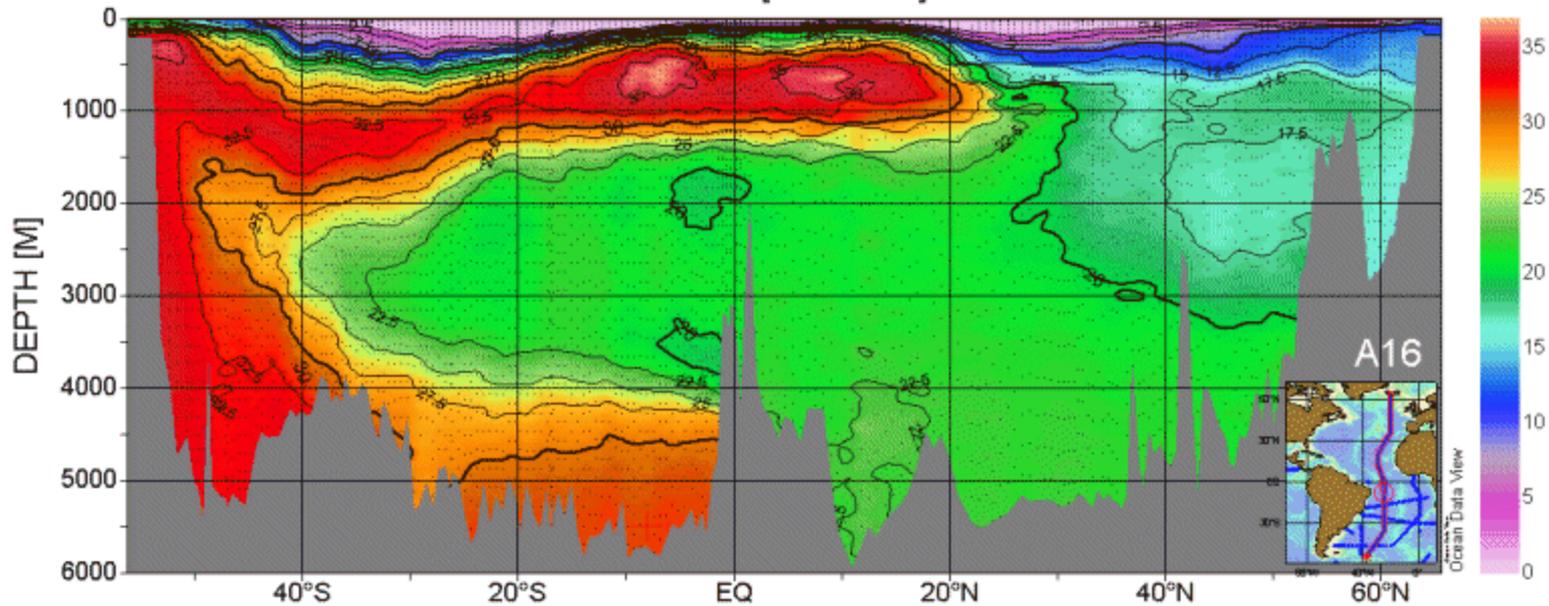


Silicate [$\mu\text{mol/kg}$]

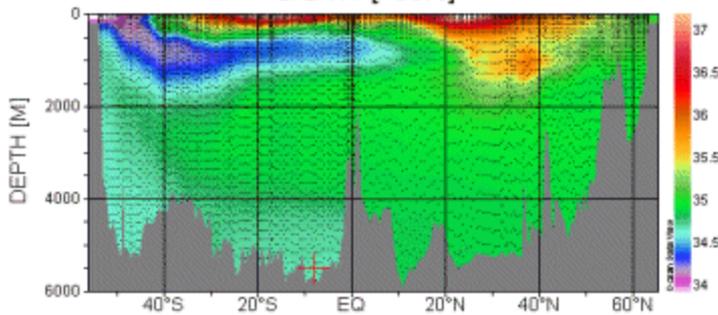


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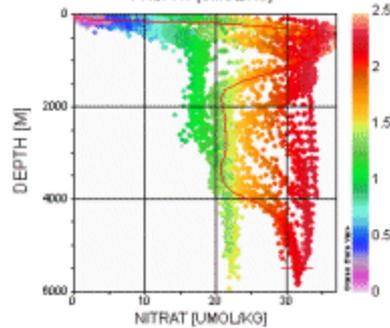
NITRAT [UMOL/KG]



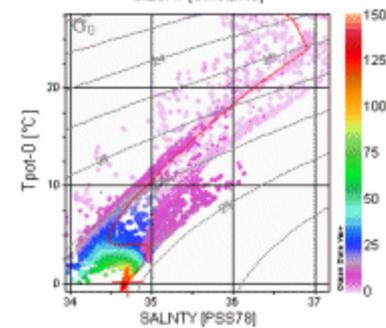
SALNTY [PSS78]



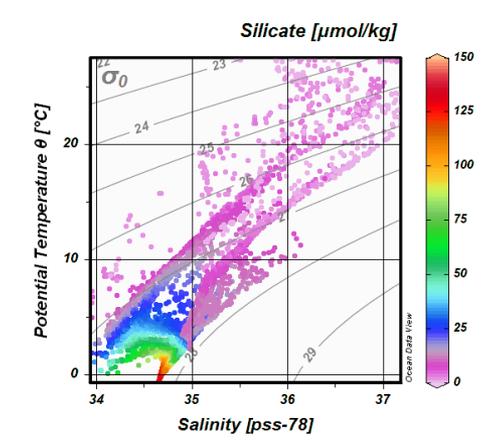
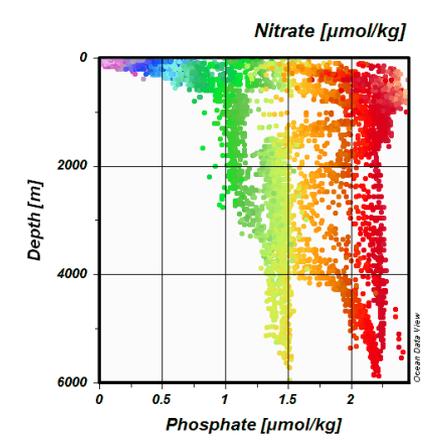
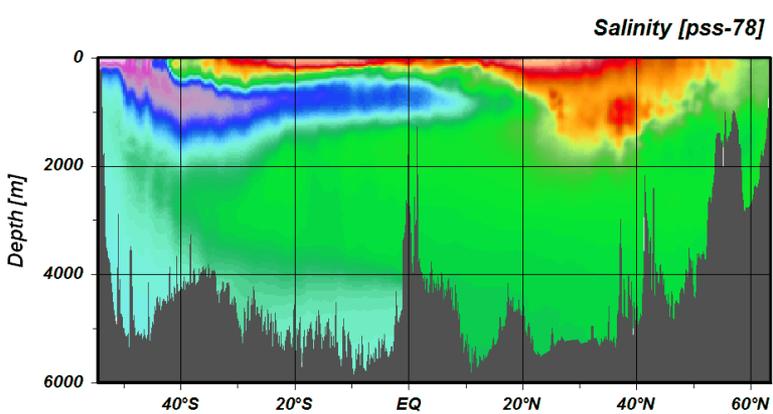
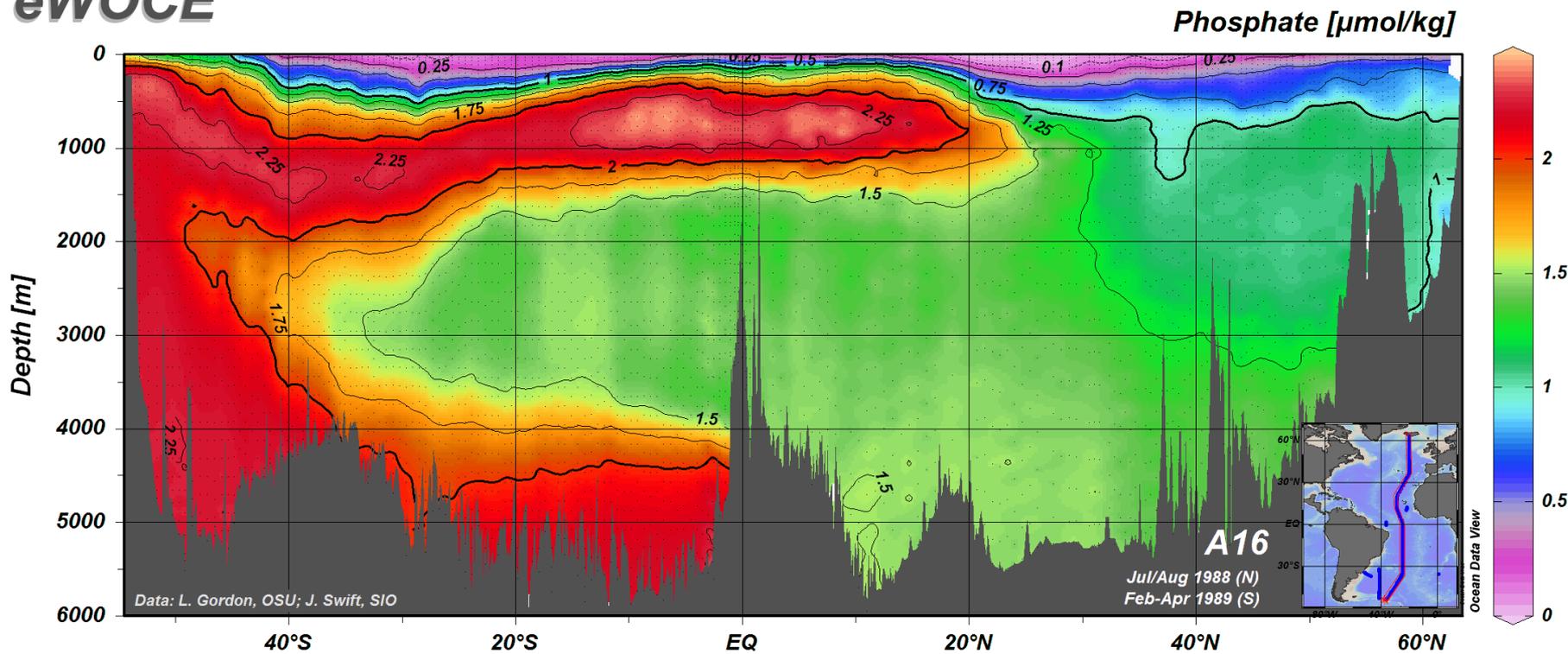
PHSPHT [UMOL/KG]



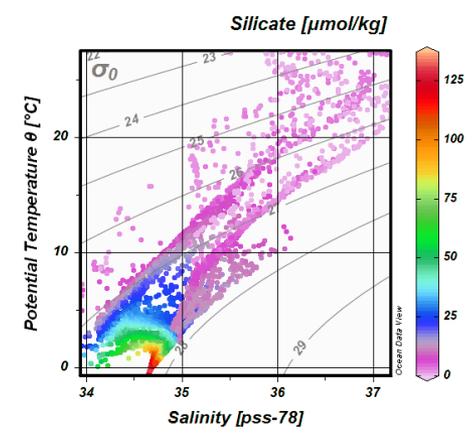
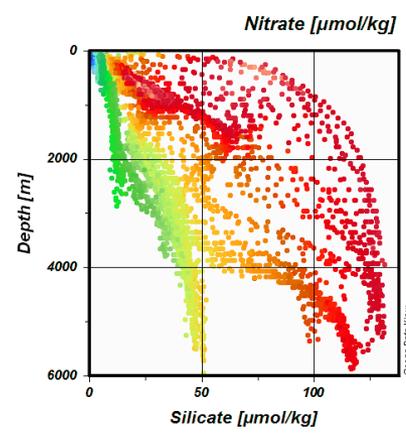
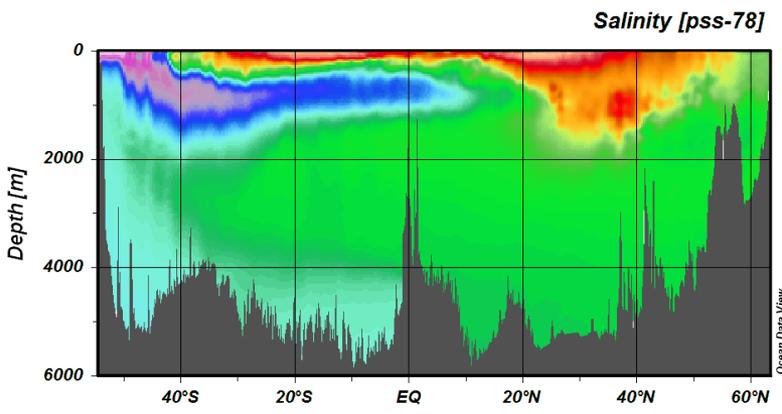
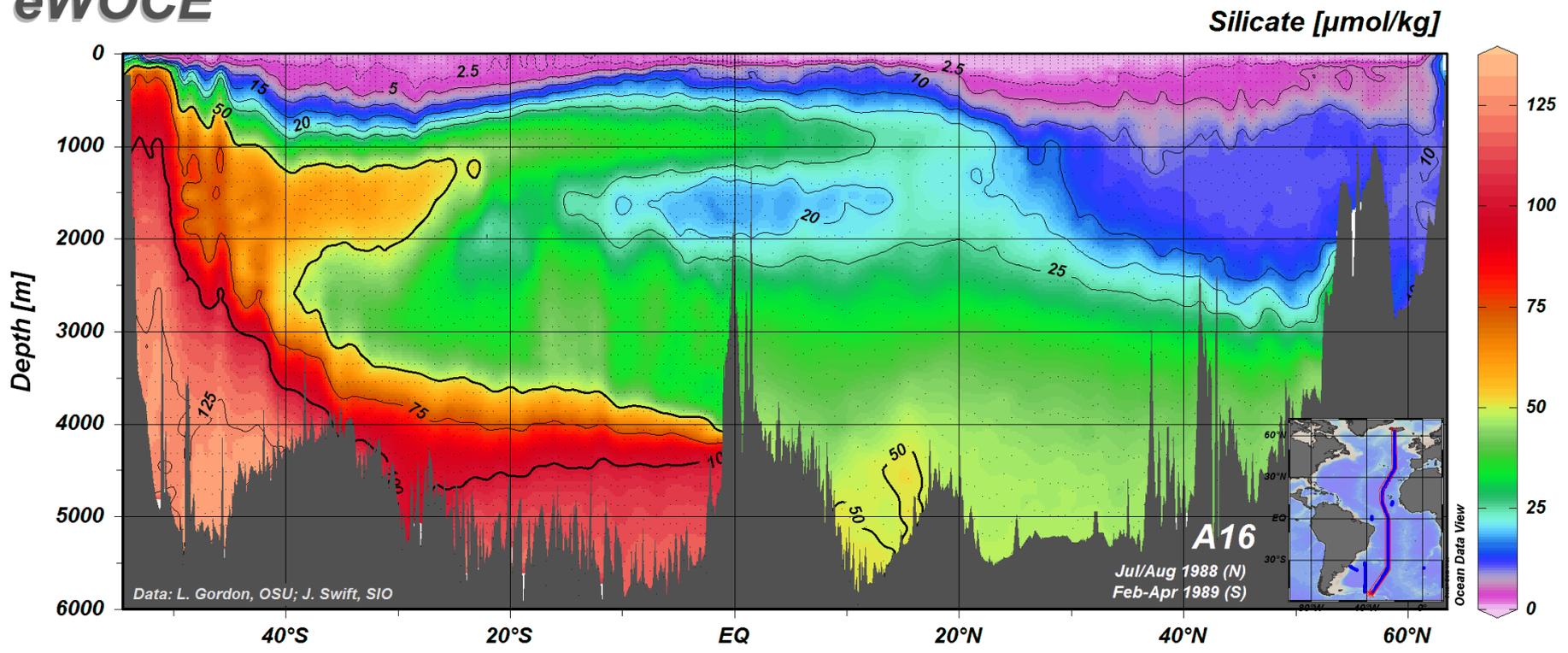
SILCAT [UMOL/KG]



eWOCE

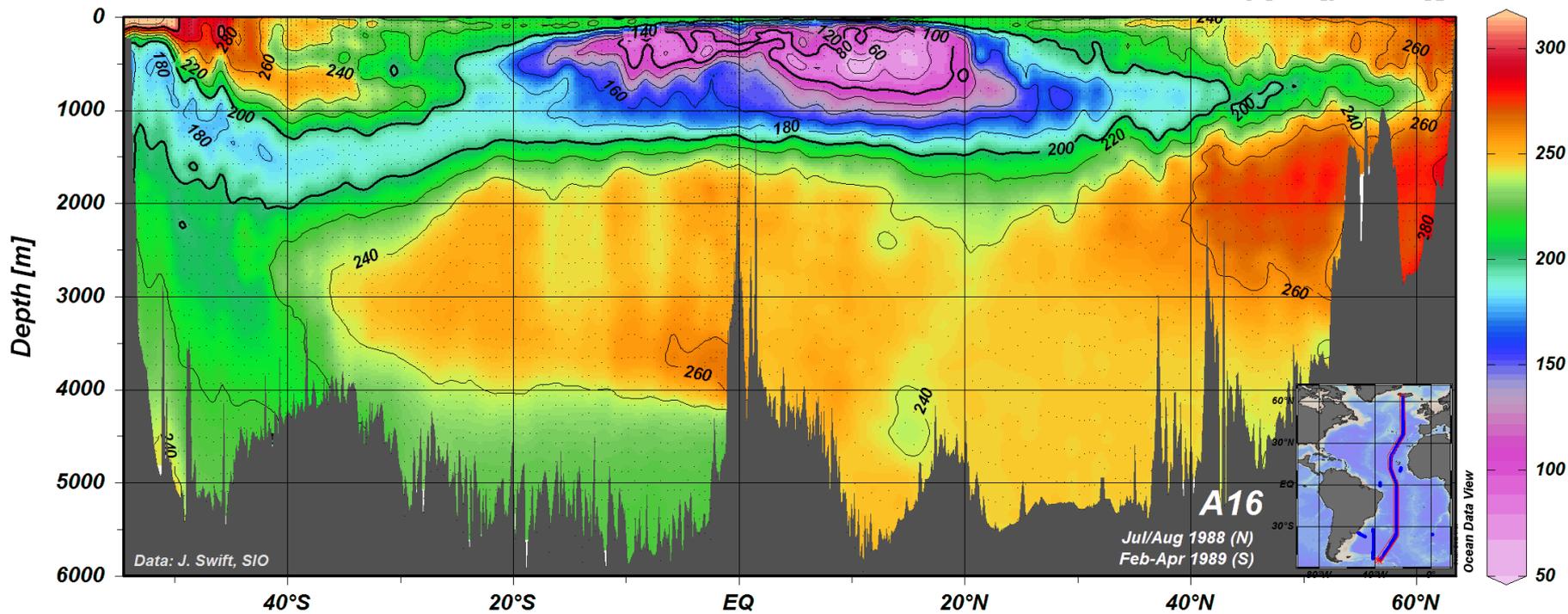


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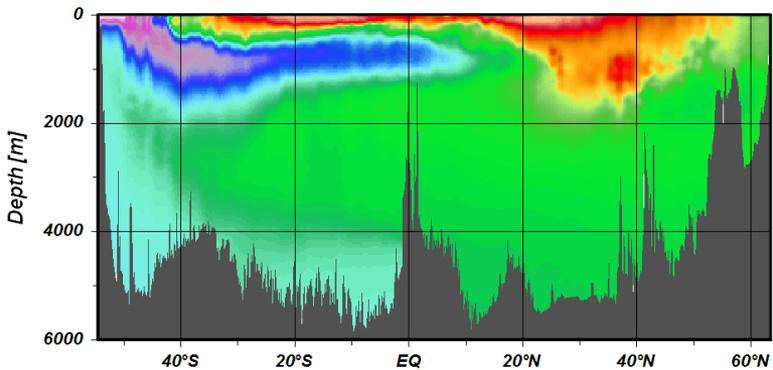


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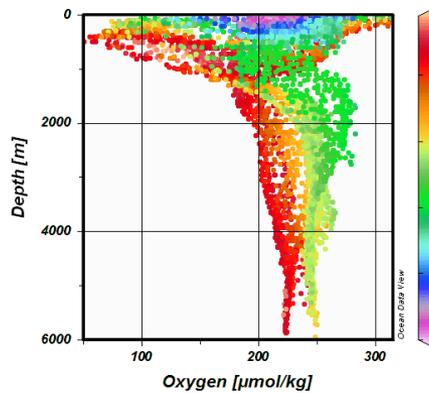
Oxygen [$\mu\text{mol/kg}$]



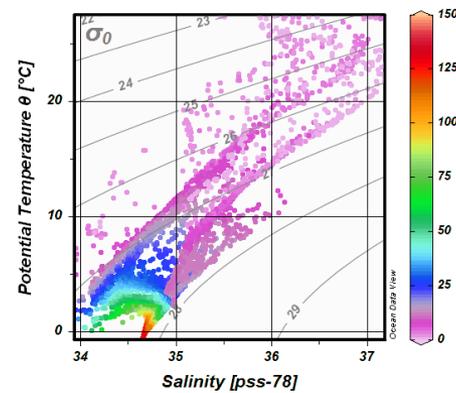
Salinity [pss-78]



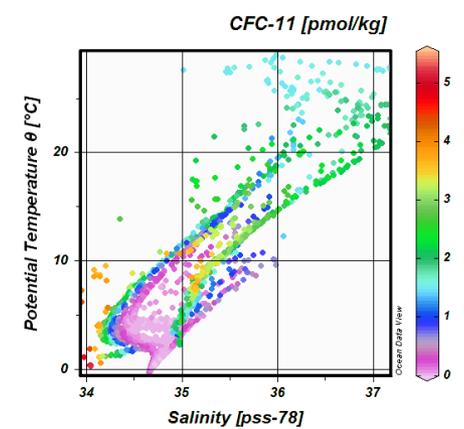
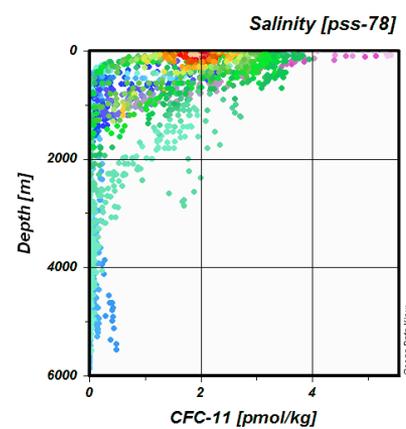
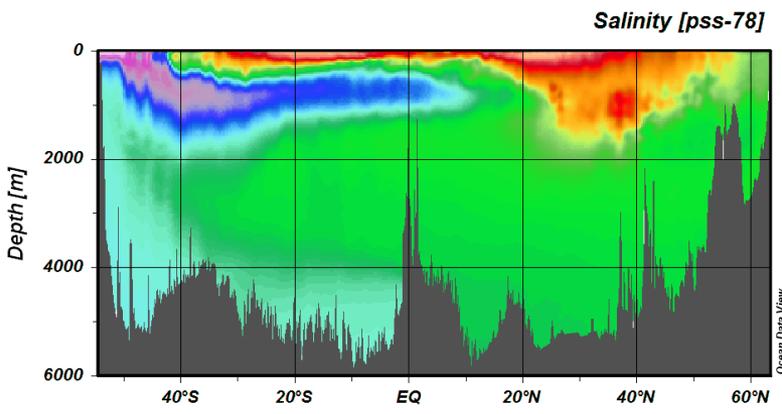
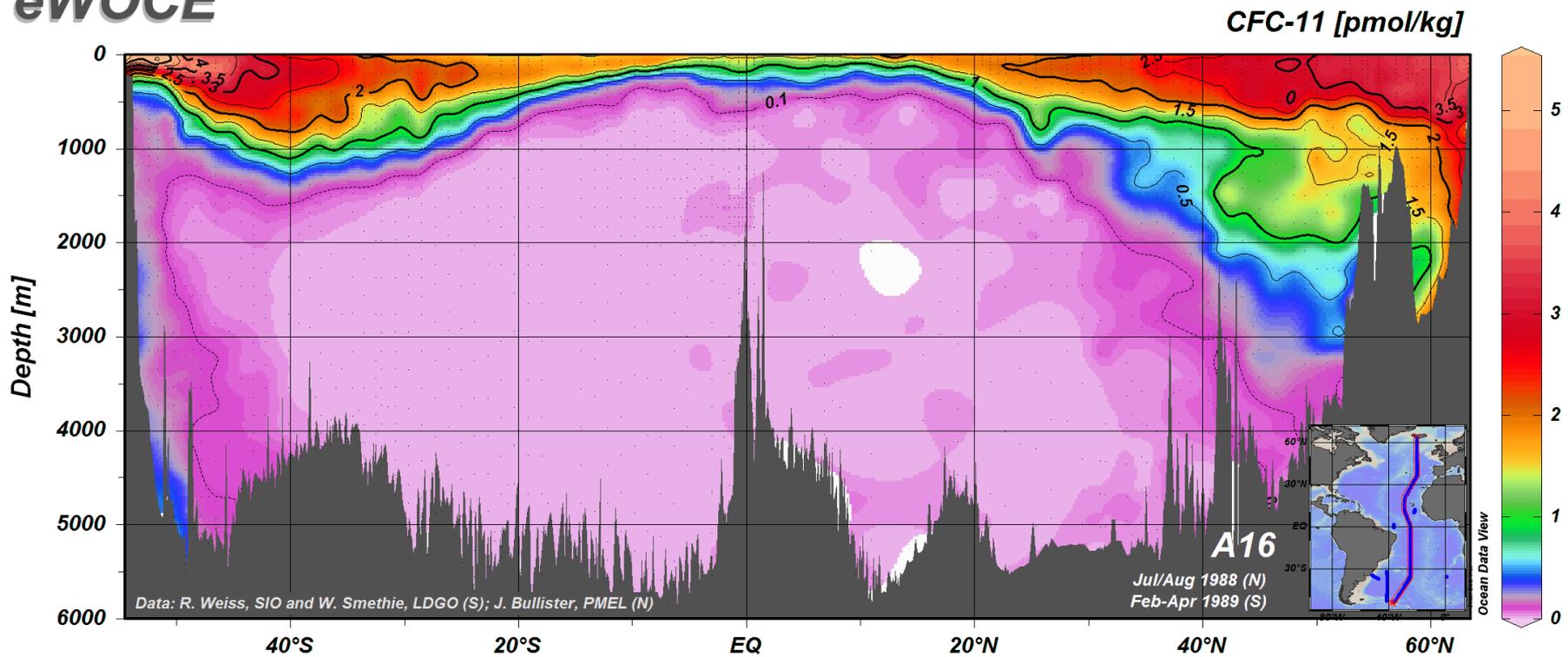
Phosphate [$\mu\text{mol/kg}$]



Silicate [$\mu\text{mol/kg}$]



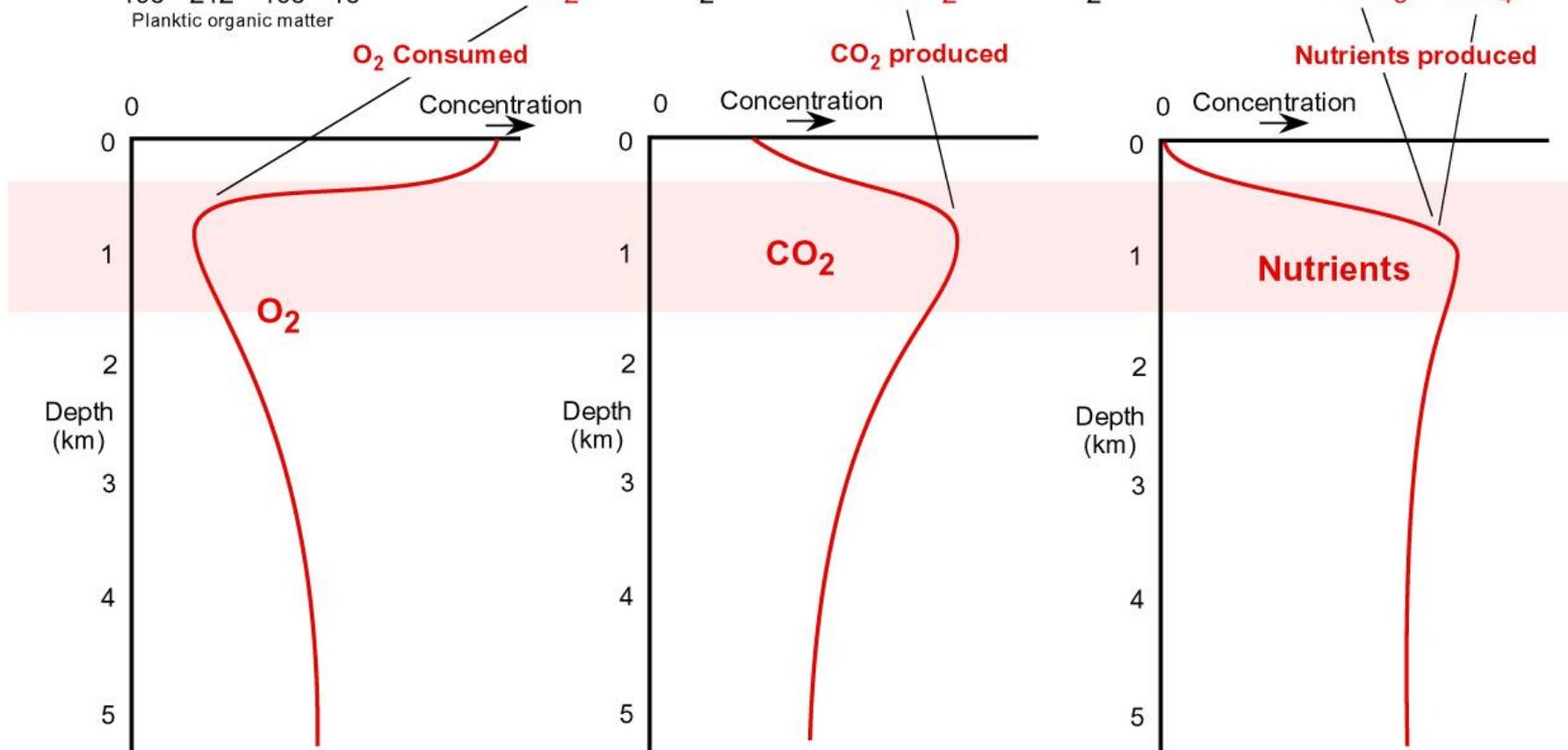
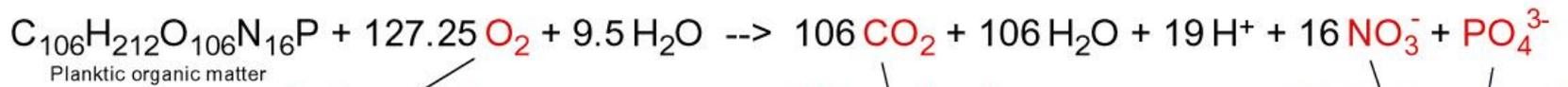
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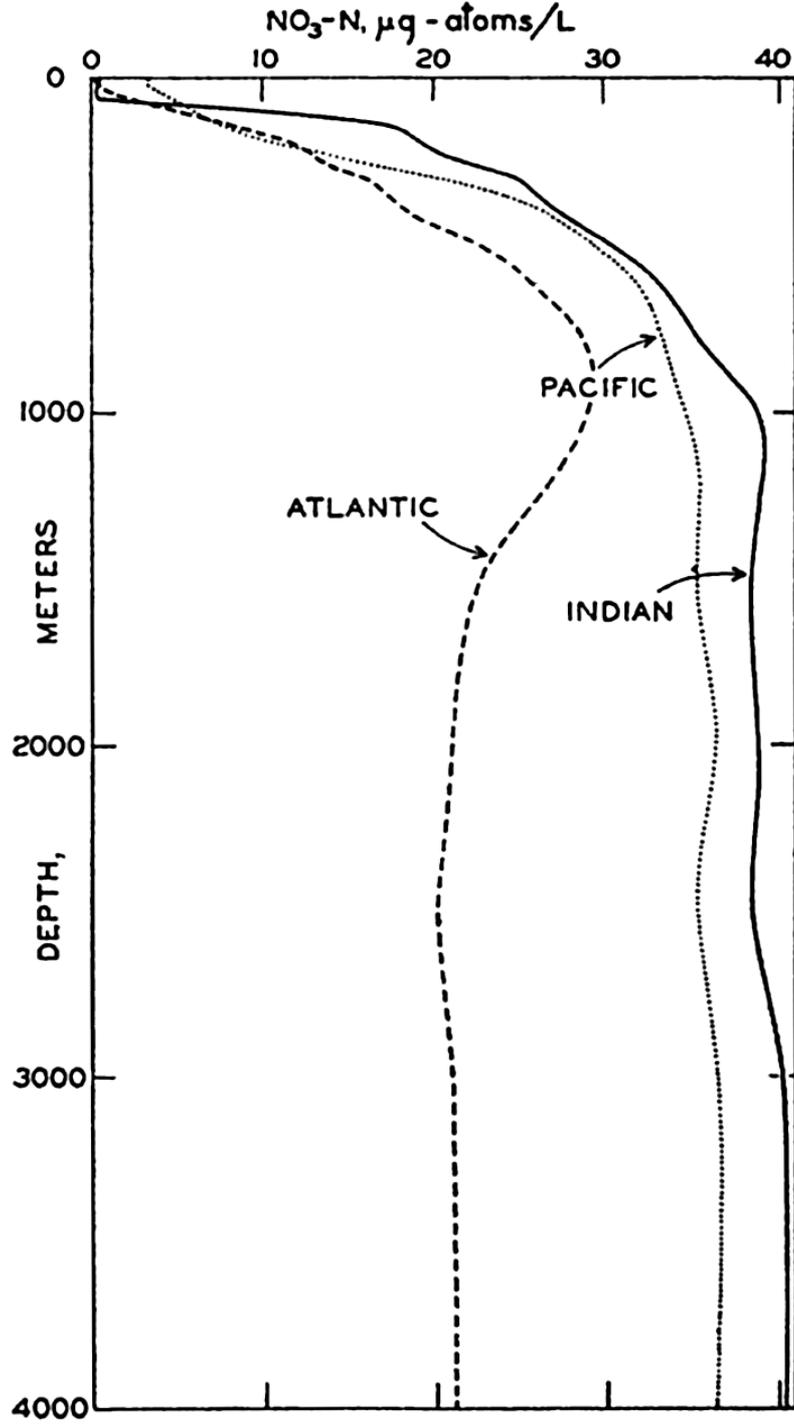


Variation in concentration of solutes in the oceans IV: oxidation of sinking organic particles - a summary

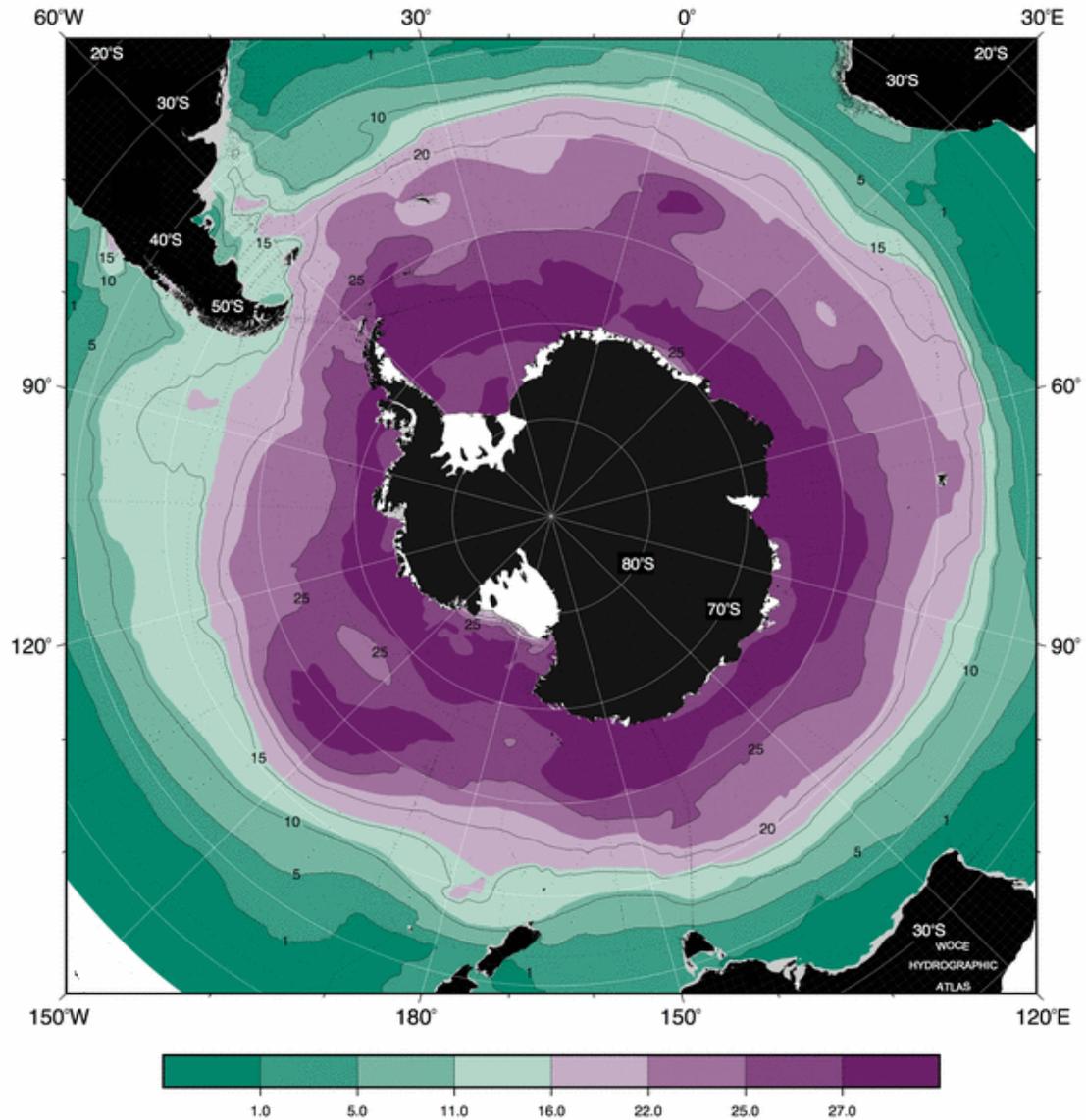
Parts I, II, and III of this series have shown how concentrations of dissolved nutrients, O₂, and CO₂ vary both vertically and laterally through the oceans. This page merely emphasizes the linkages among these distributions, all of which are greatly influenced by oxidation of organic particles sinking from the photic zone through the thermocline and into the

deeper ocean. As one can see below, this oxidation of organic matter consumes O₂ and releases CO₂ and the nutrients NO₃⁻ and PO₄³⁻. In addition, oxidative (or other) decay of organic coatings on siliceous tests in the deep ocean allows dissolution of those siliceous tests, allowing release of a third critical nutrient, SiO_{2(aq)}, in the deep water.



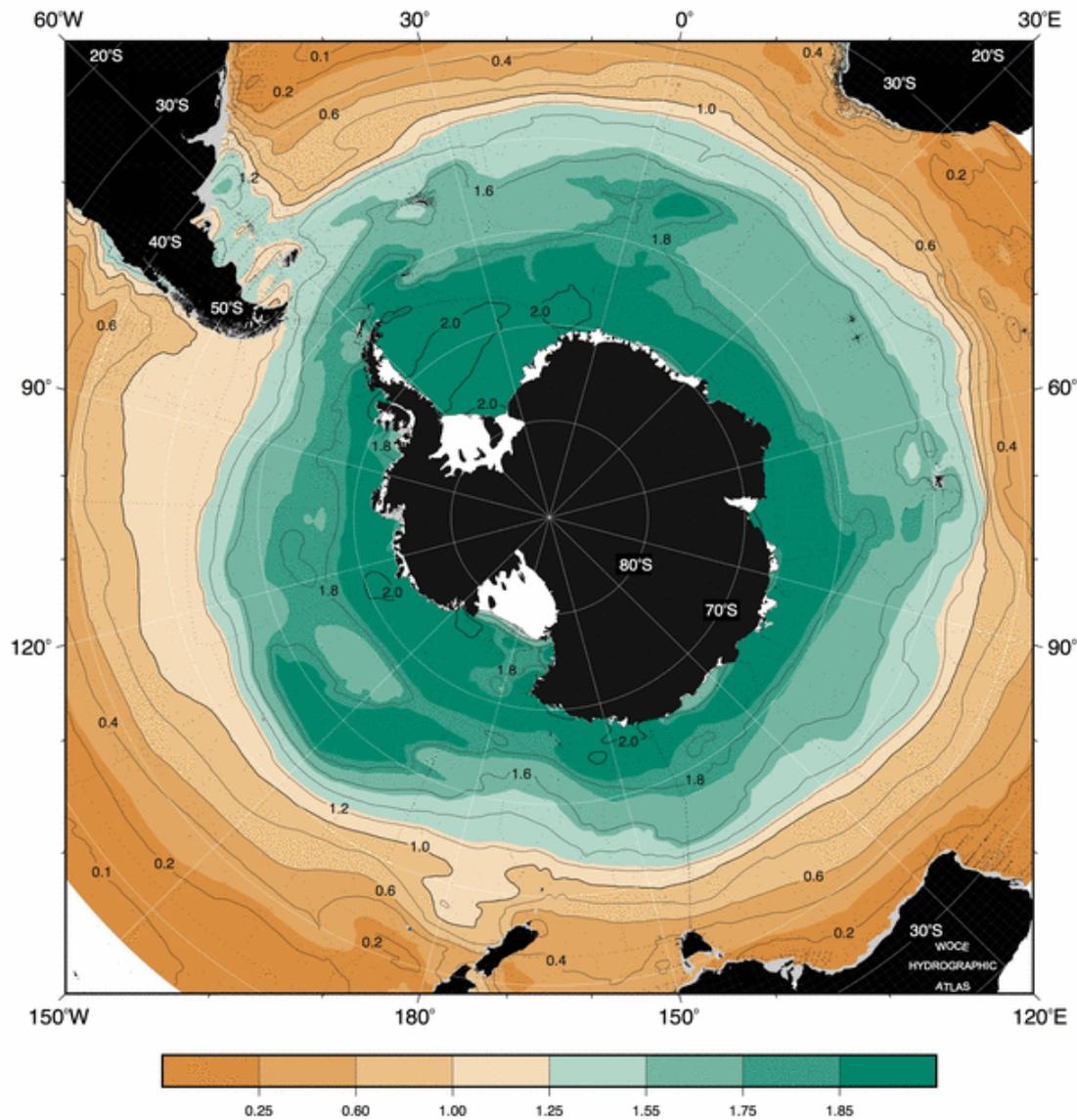


Depth: 50m

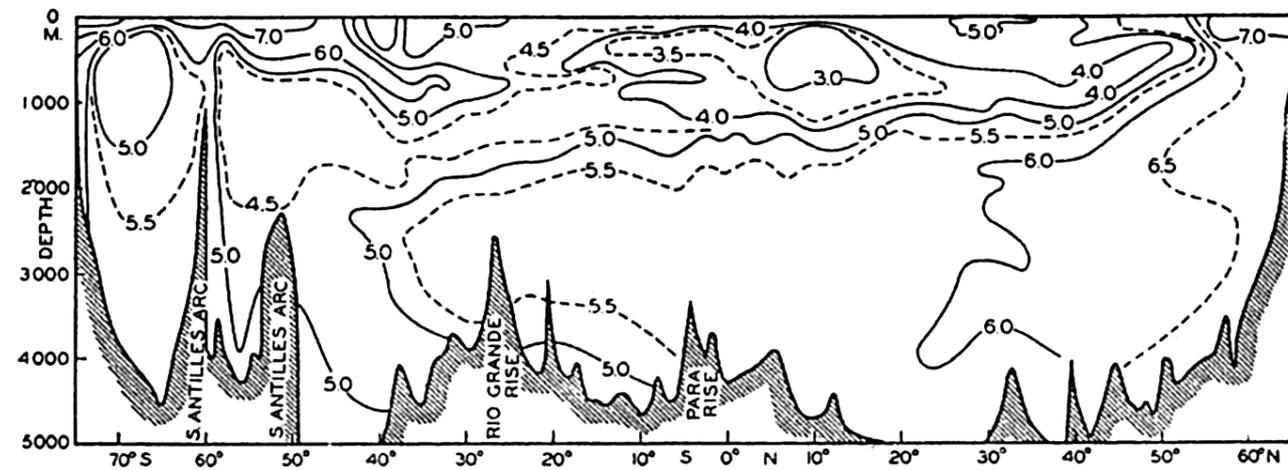
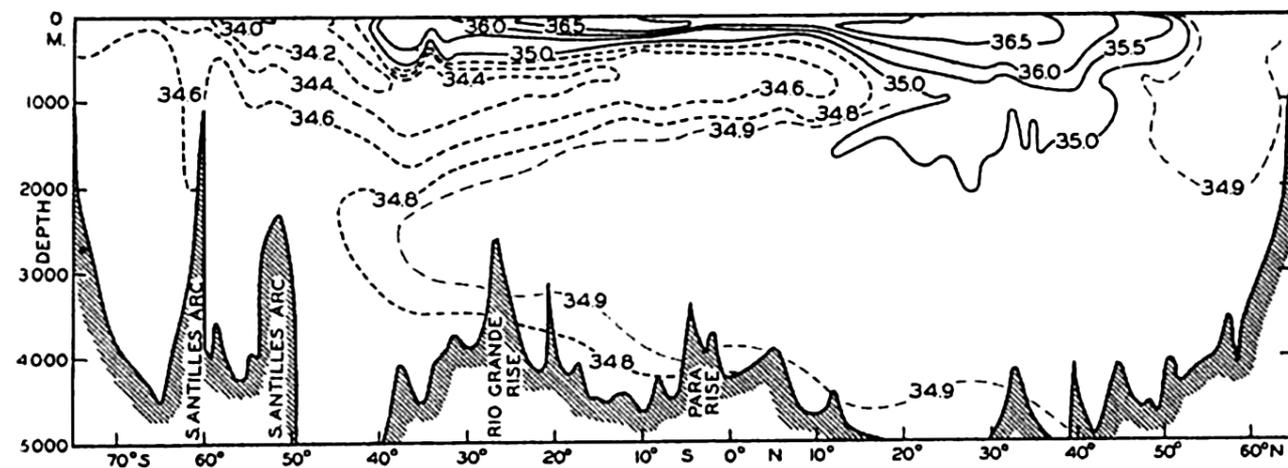
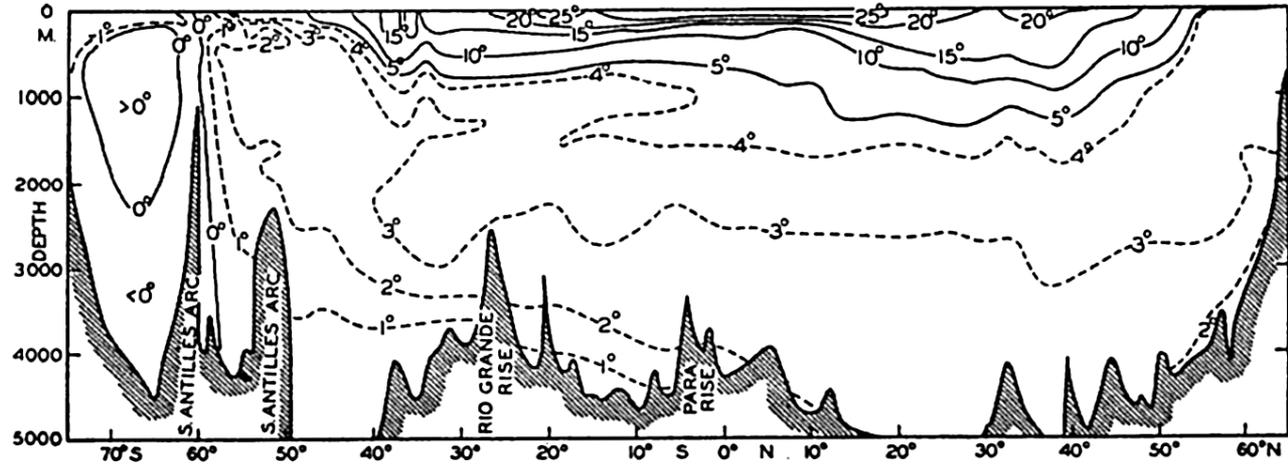


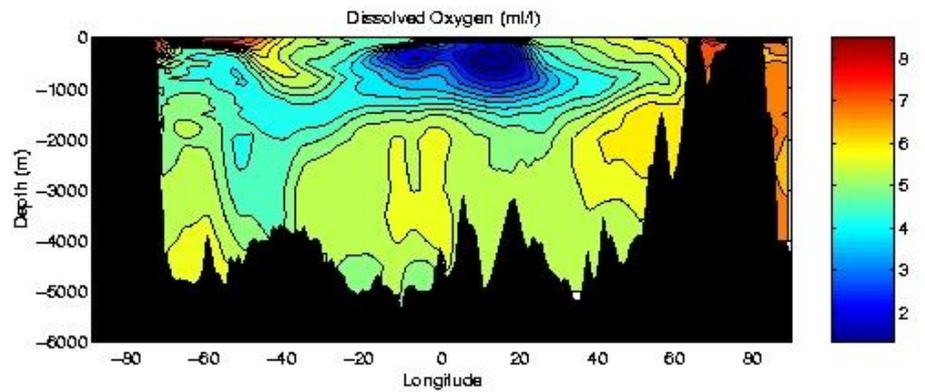
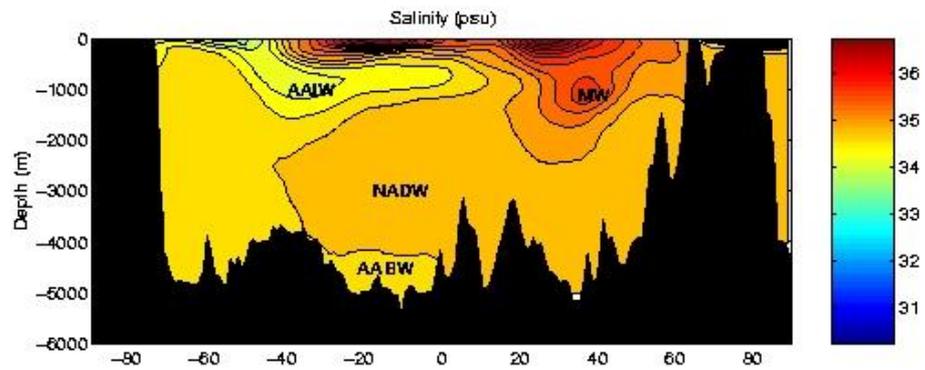
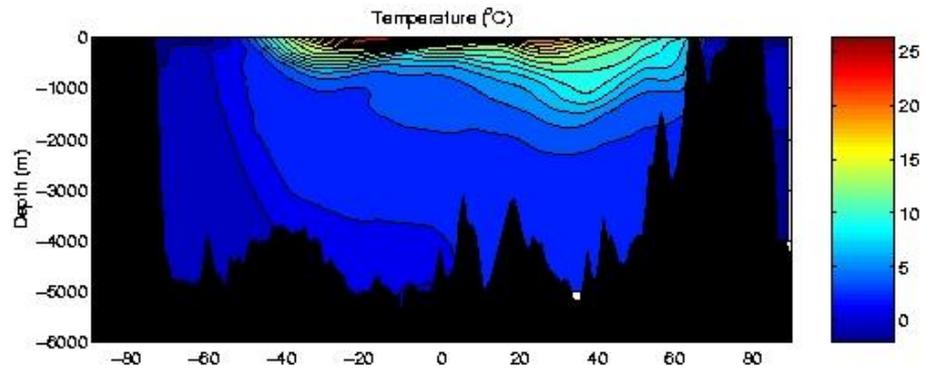
Nitrate ($\mu\text{mol/kg}$)

Depth: 50m



Phosphate ($\mu\text{mol/kg}$)





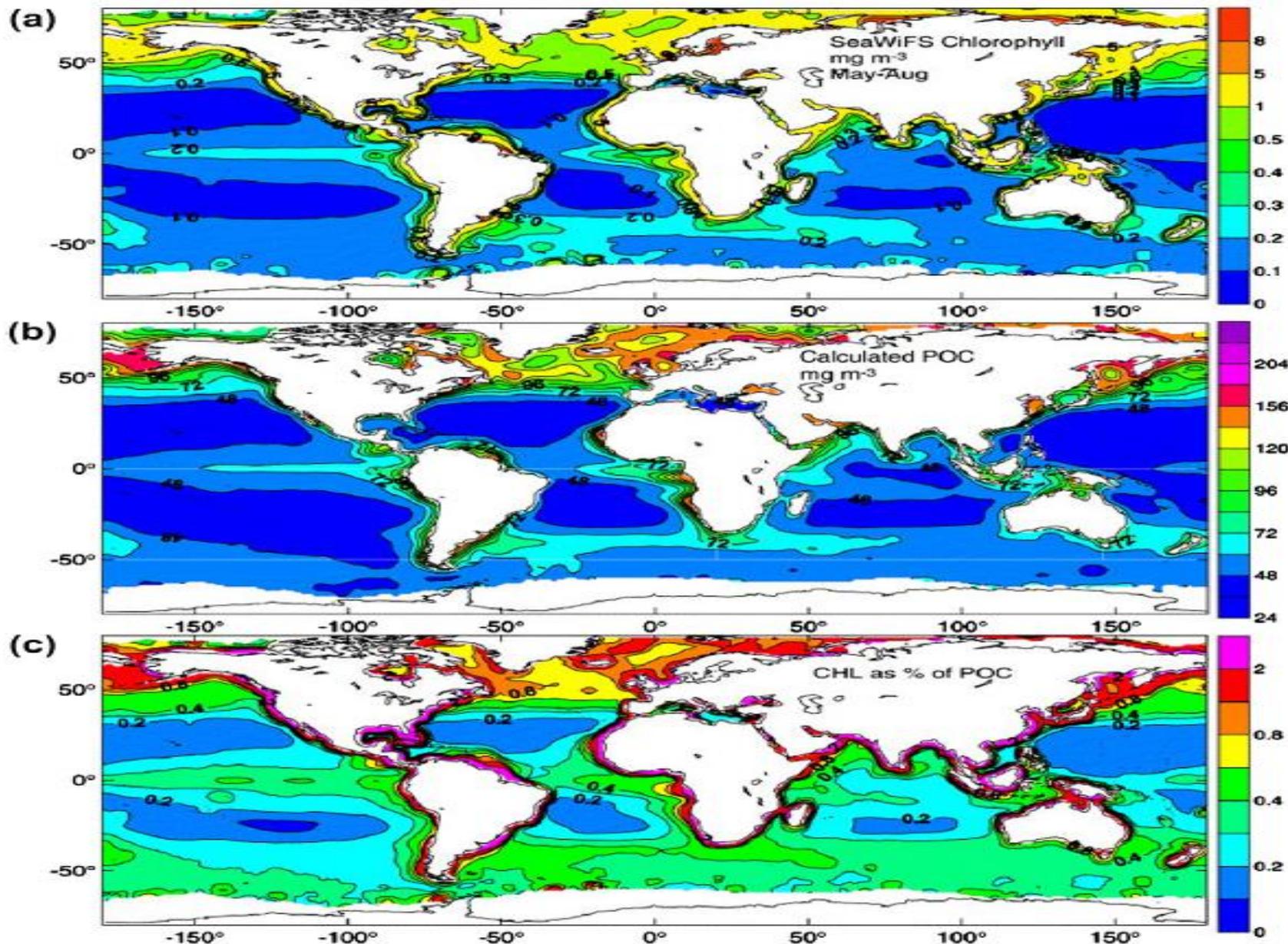
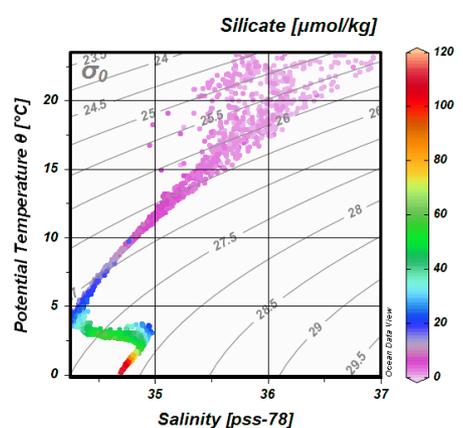
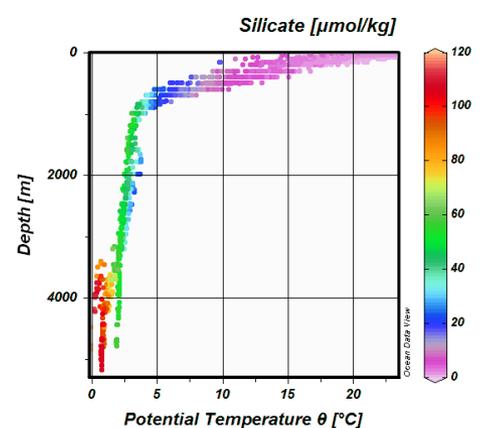
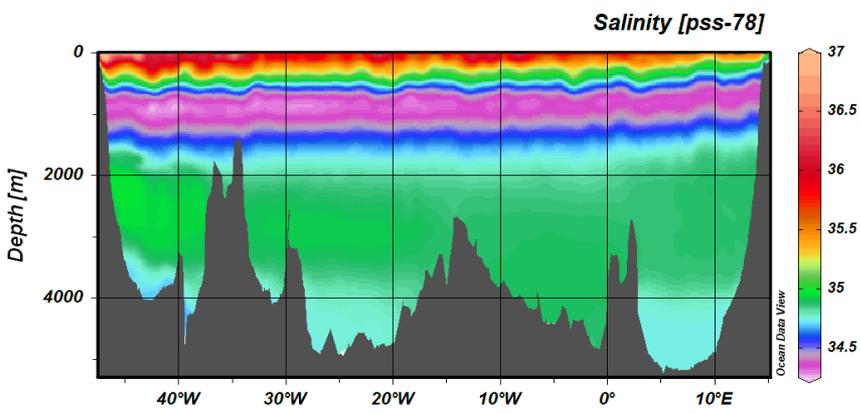
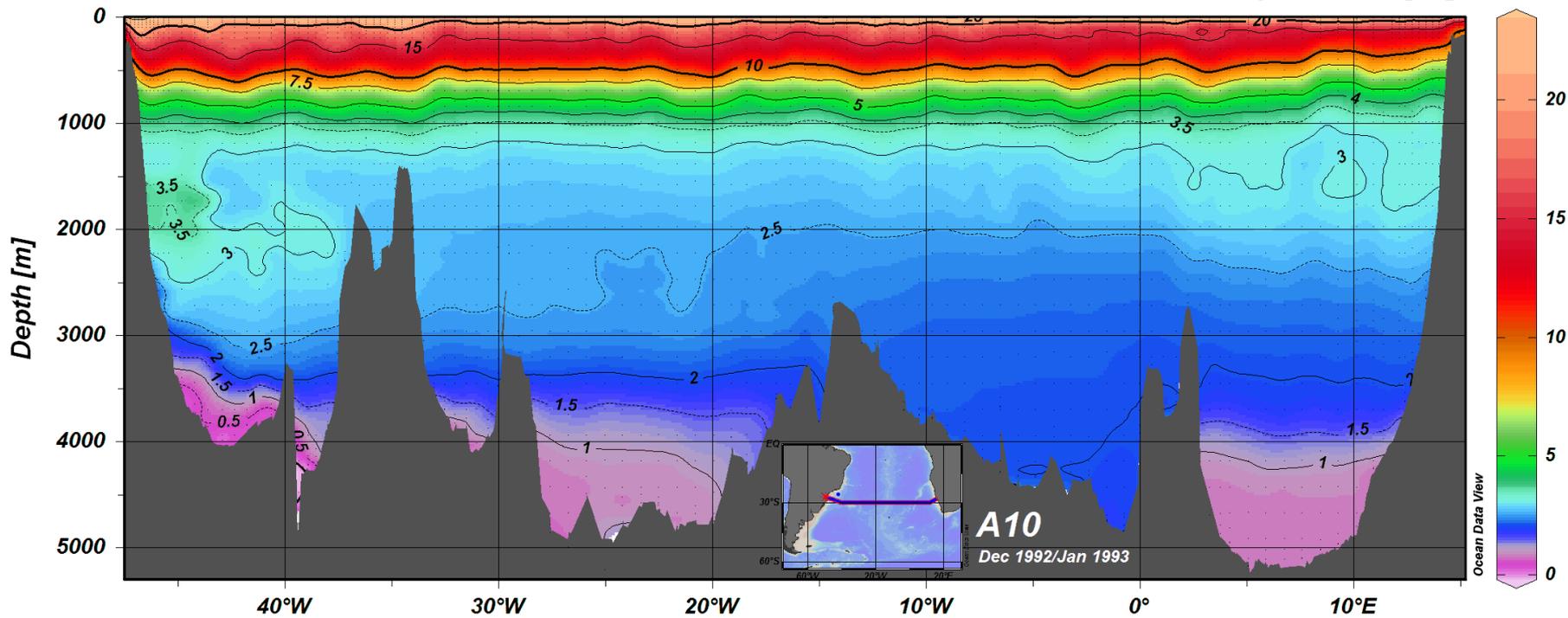
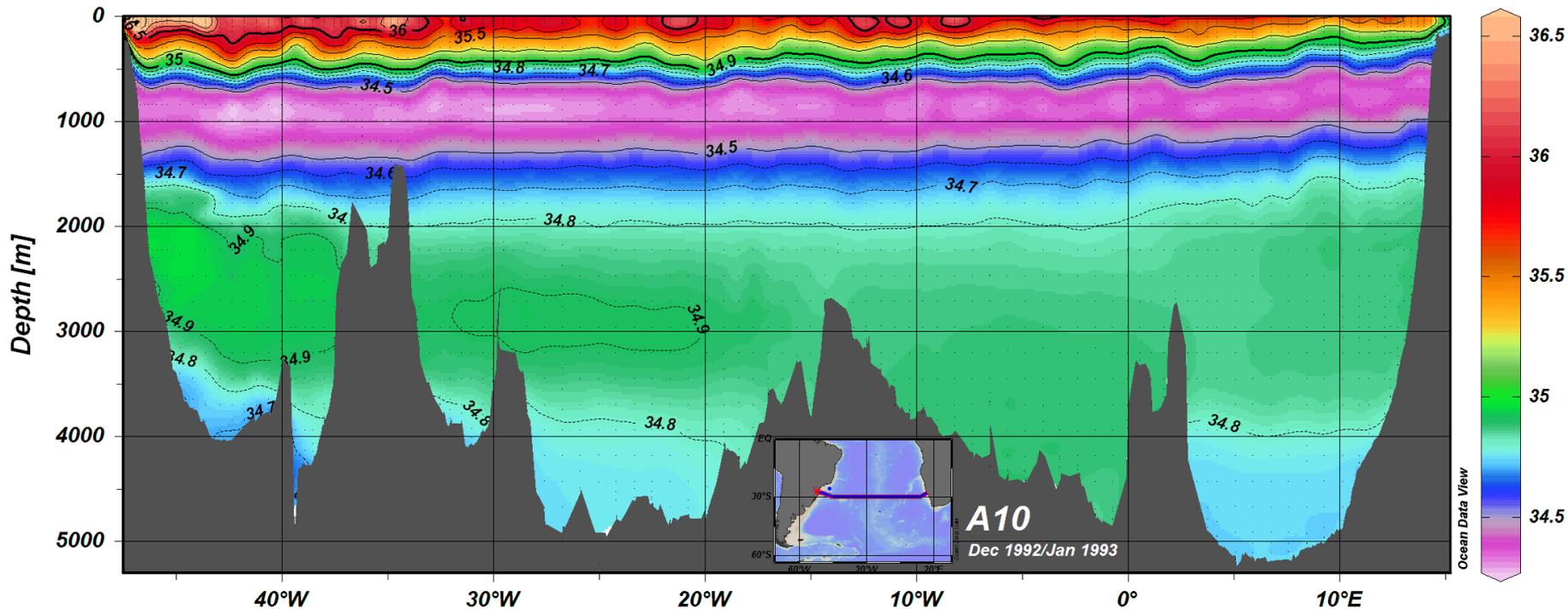


FIGURE S4.2 (a) Chlorophyll (mg m^{-3}), (b) particulate organic carbon (POC; mg m^{-3}), derived from SeaWiFS ocean color data, averaged May-August, 1997–2002, and (c) chlorophyll as % of POC. Source: From Gardner, Mishov, and Richardson, (2006).

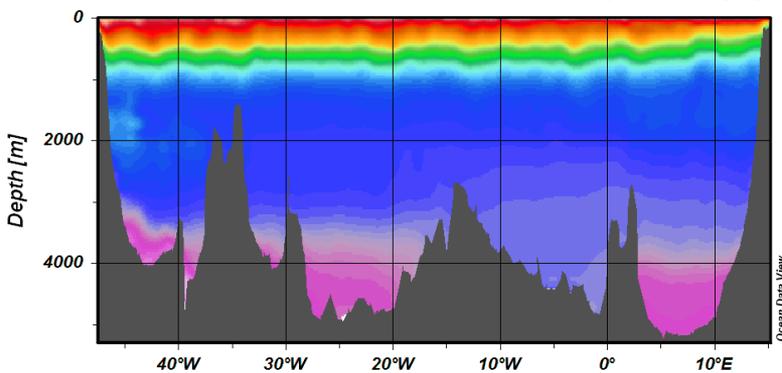


eWOCE

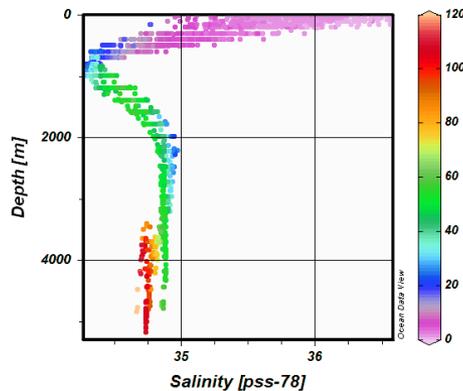
Salinity [pss-78]



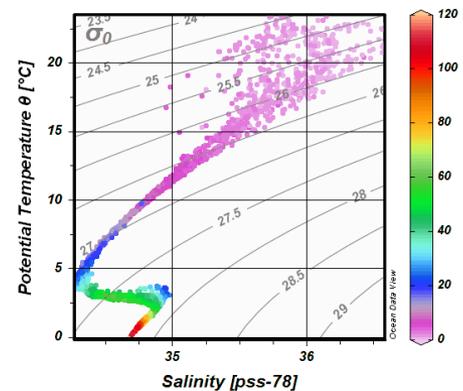
Potential Temperature θ [°C]



Silicate [$\mu\text{mol/kg}$]

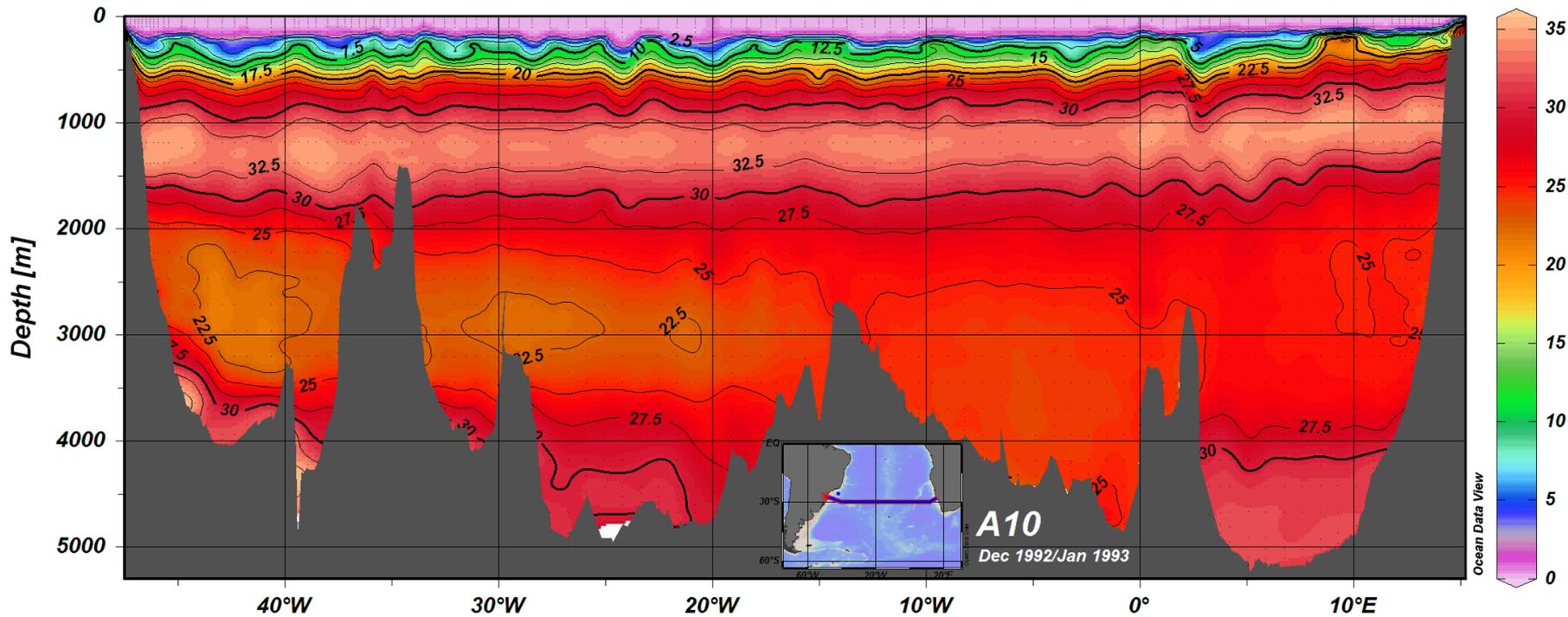


Silicate [$\mu\text{mol/kg}$]

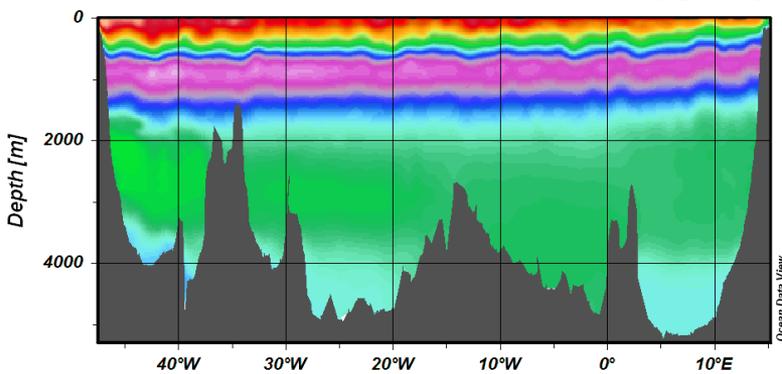


eWOCE

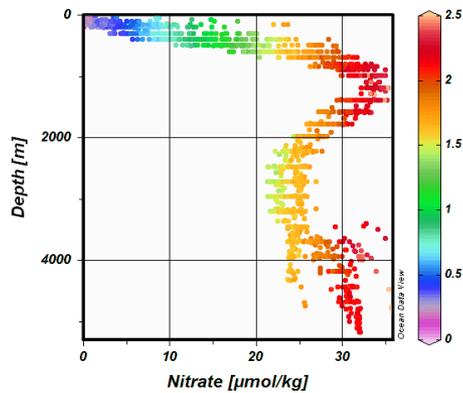
Nitrate [$\mu\text{mol/kg}$]



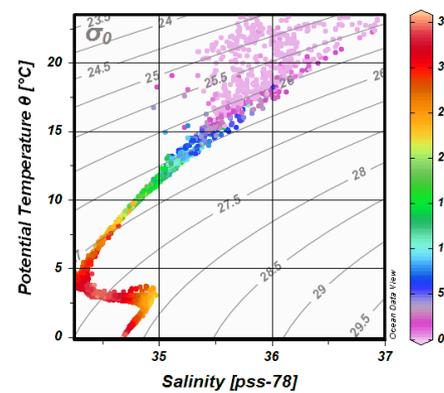
Salinity [pss-78]



Phosphate [$\mu\text{mol/kg}$]

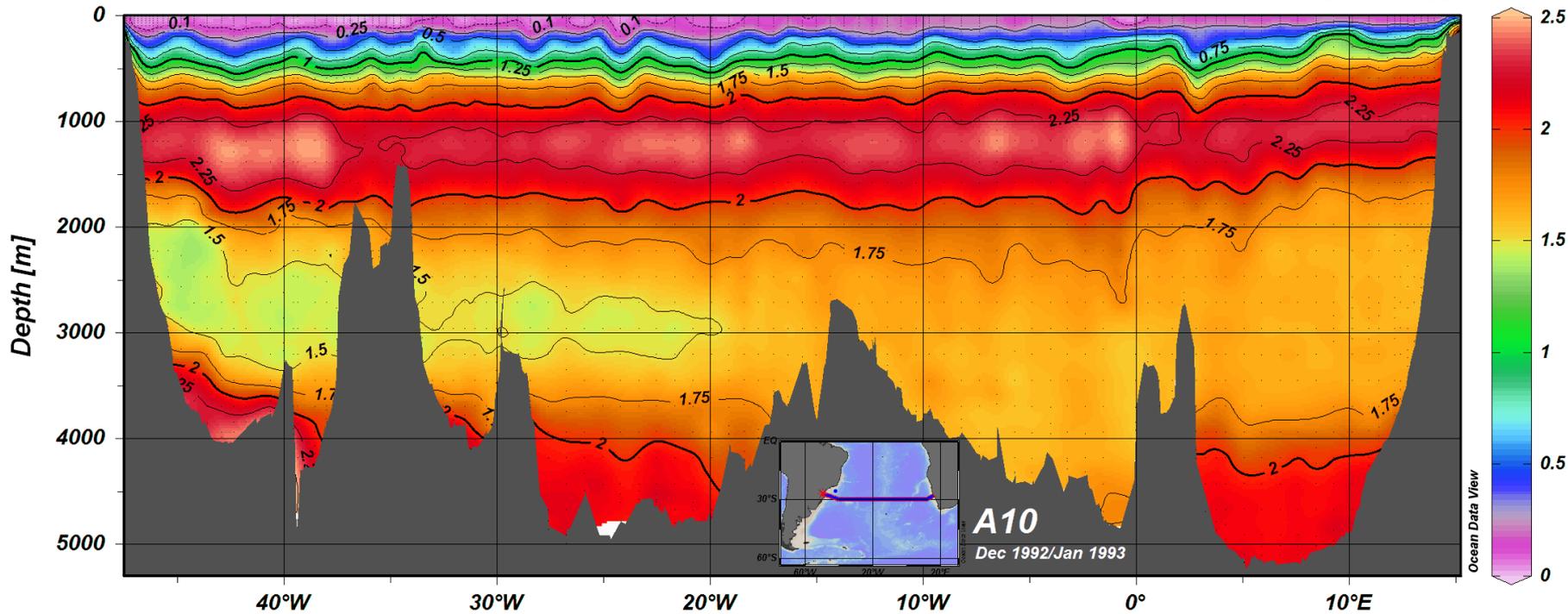


Nitrate [$\mu\text{mol/kg}$]

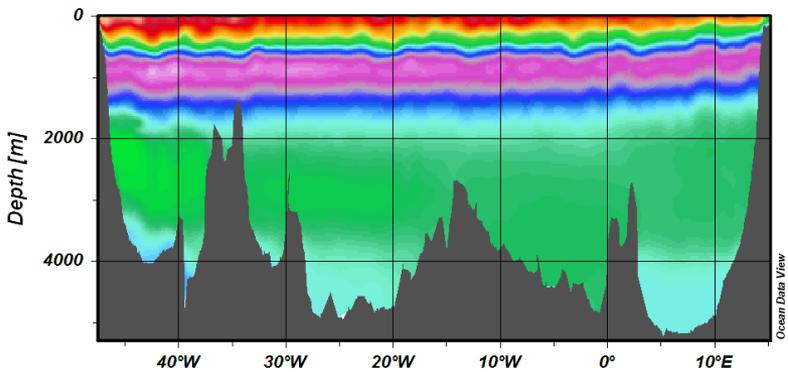


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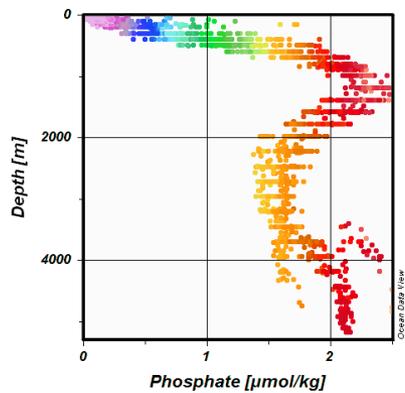
Phosphate [$\mu\text{mol/kg}$]



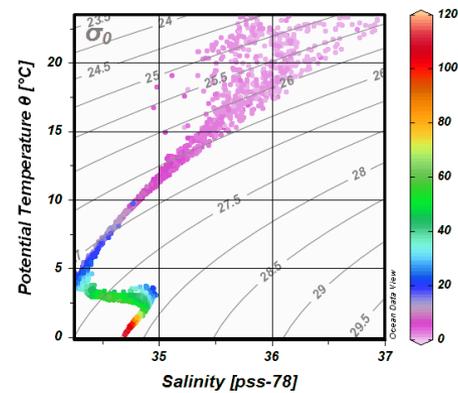
Salinity [pss-78]



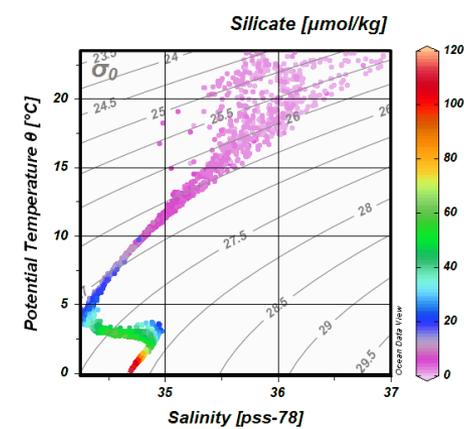
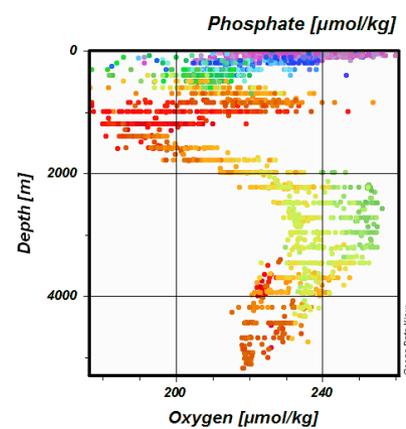
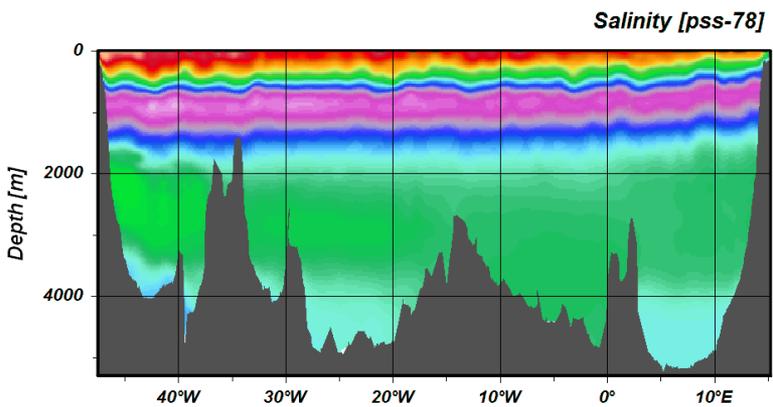
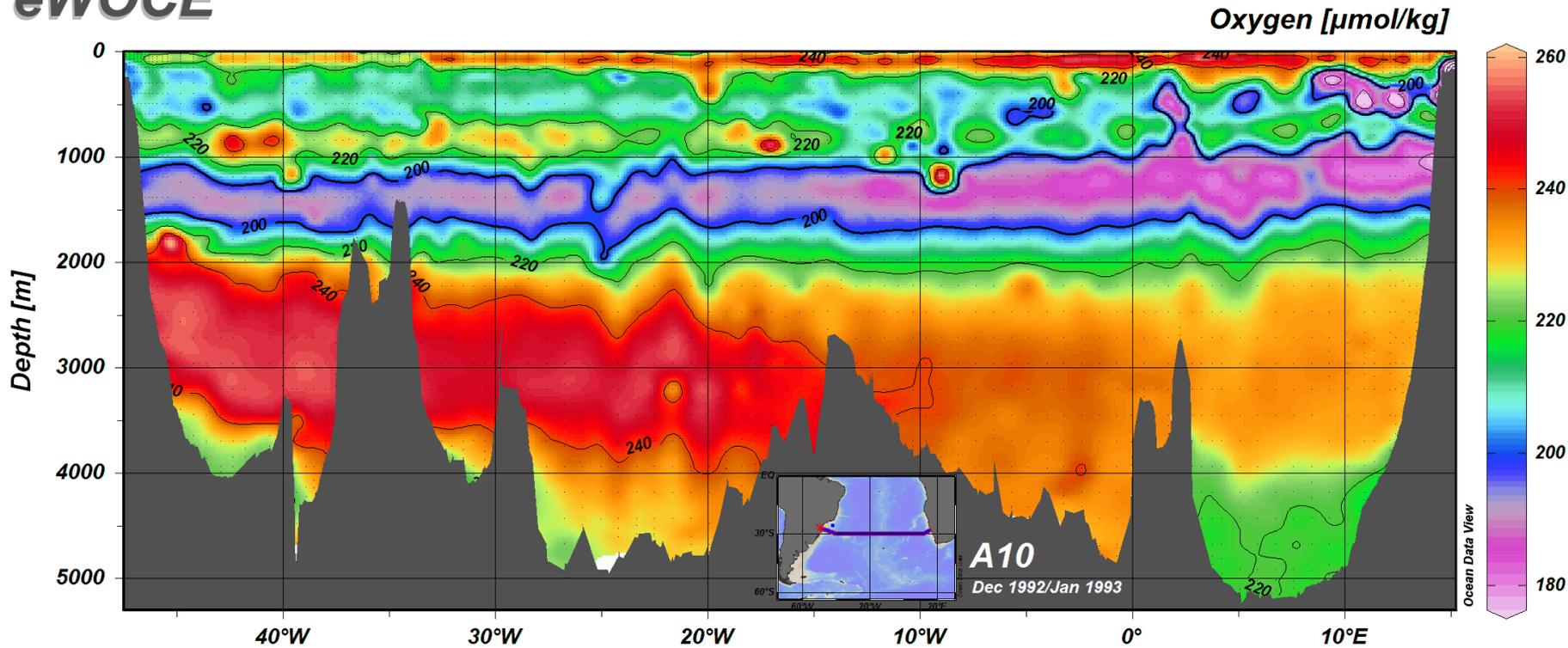
Nitrate [$\mu\text{mol/kg}$]



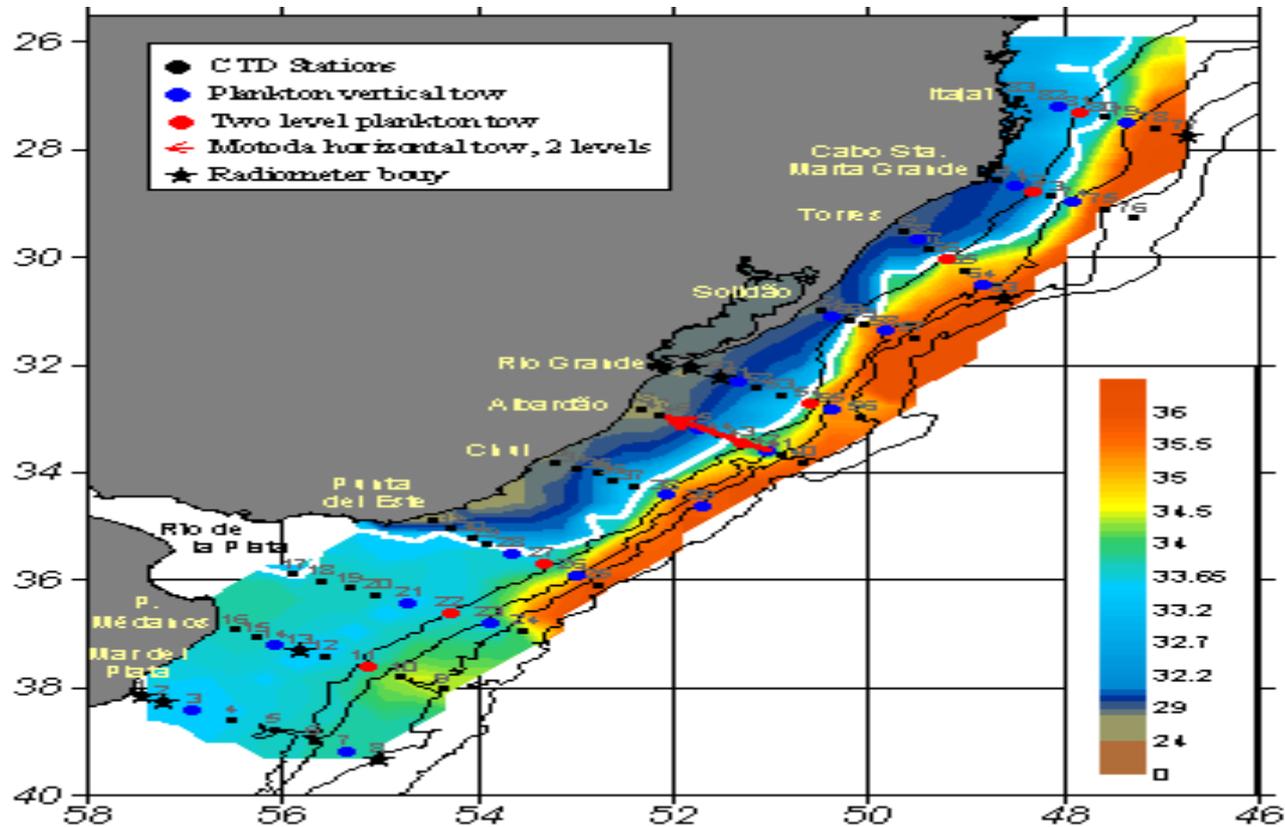
Silicate [$\mu\text{mol/kg}$]

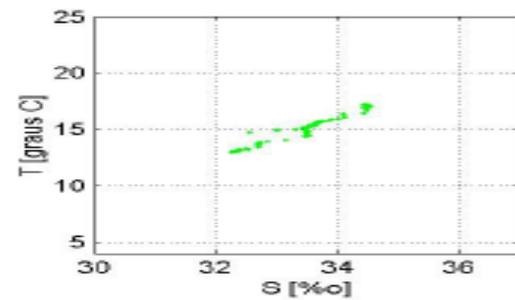
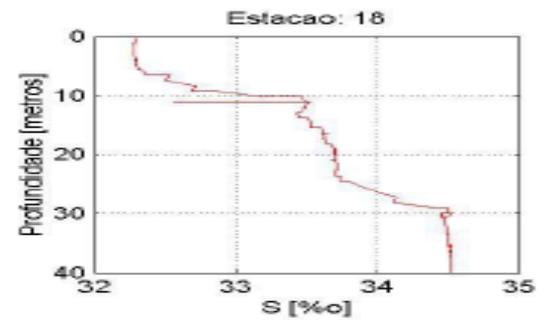
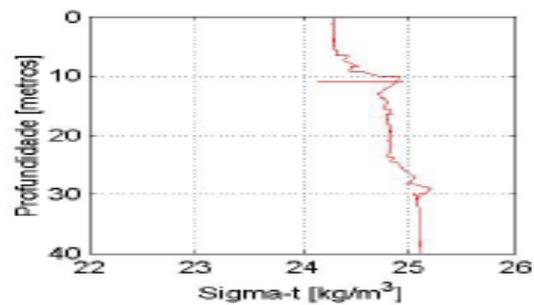
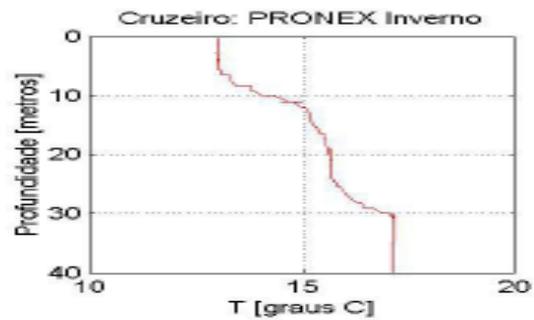
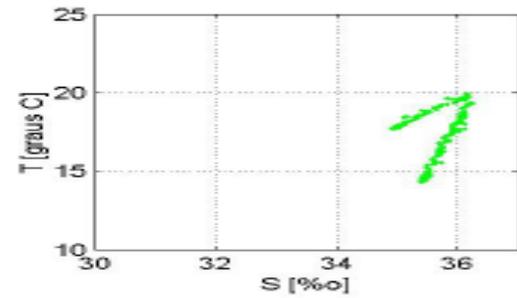
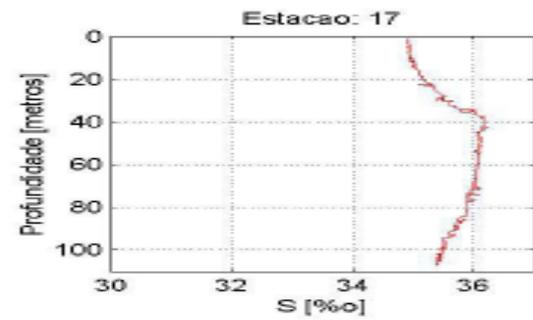
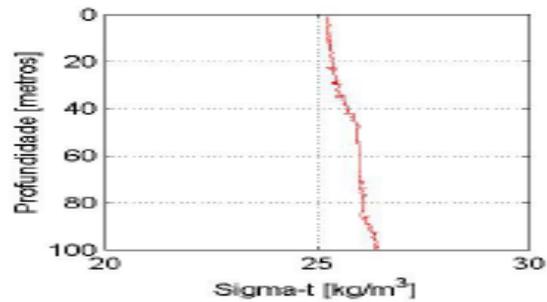
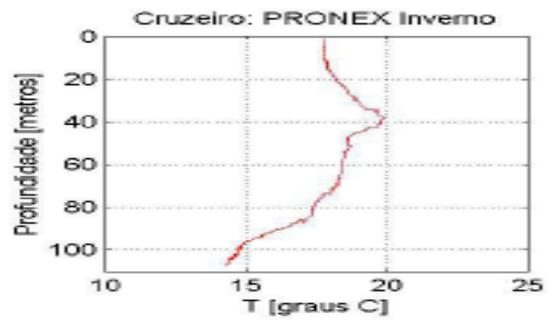


eWOCE



Regiões costeiras onde existe muita descarga fluvial ocorre geralmente uma camada de água superficial de baixa salinidade





DENSIDADE ???

relação entre a massa de qualquer substância e seu volume

Variação 1021 kg/m³ na superfície até 1070 kg/m³ no fundo

Conseqüência da Salinidade, Temperatura e Pressão ($\rho(S,T,p)$ é determinada pela equação de estado da água do mar)

A densidade vai determinar onde uma massa de água vai se estabelecer

Gde parte das correntes oceânicas pode ser adequadamente descrita conhecendo-se o campo de pressão.

A pressão em um determinado ponto do oceano é definida como o peso da água acima deste ponto. Este peso, por sua vez, é função da altura da coluna d'água em questão e da sua densidade.

Cálculo da densidade da água do mar

The density (ρ , kg/m³) of seawater as a function of practical salinity (S), temperature (t , °C) and applied or gauge pressure (p , decibars) is given by

$$\rho(S, t, p) = \rho(S, t, 0) / [1 - p/K(S, t, p)] \quad (7)$$

where $K(S, t, p)$ is the secant bulk modulus. The specific volume ($V = 1/\rho$ m³/kg) of seawater can be obtained from

$$V(S, t, p) = V(S, t, 0) \cdot [1 - p/K(S, t, p)] \quad (8)$$

$$\rho(S, t, 0) = \rho_w + (b_0 + b_1 t + b_2 t^2 + b_3 t^3 + b_4 t^4) S + (c_0 + c_1 t + c_2 t^2) S^{3/2} + d_0 S^2 \quad (13)$$

$$b_0 = + 8.24493E-1$$

$$c_0 = - 5.72466E-3$$

$$b_1 = - 4.0899 E-3$$

$$c_1 = + 1.0227 E-4$$

$$b_2 = + 7.6438 E-5$$

$$c_2 = - 1.6546 E-6$$

$$b_3 = - 8.2467 E-7$$

The density of the reference pure water (SMOW, Craig, 1961) is given by (IUPAC, 1976)

$$\rho_w = a_0 + a_1 t + a_2 t^2 + a_3 t^3 + a_4 t^4 + a_5 t^5 \quad (14)$$

where

$$a_0^* = + 999.842594 [-28.263737]$$

$$a_1 = + 6.793952 E-2$$

$$a_2 = - 9.095290 E-3$$

$$a_3 = + 1.001685 E-4$$

$$a_4 = - 1.120083 E-6$$

$$a_5 = + 6.536332 E-9$$

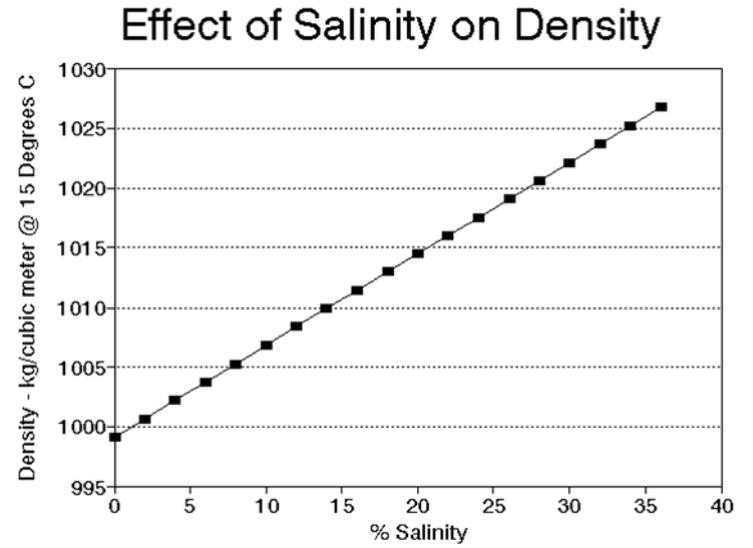
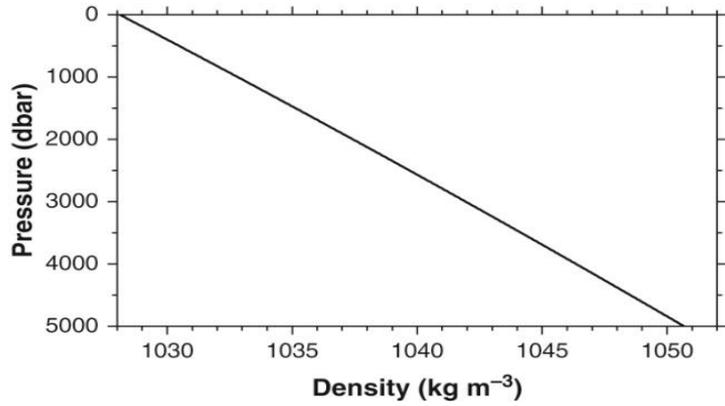
DISTRIBUIÇÃO HORIZONTAL DA DENSIDADE

aumenta de 22 próximo do equador até 26-27 em 50° a 60° de latitude

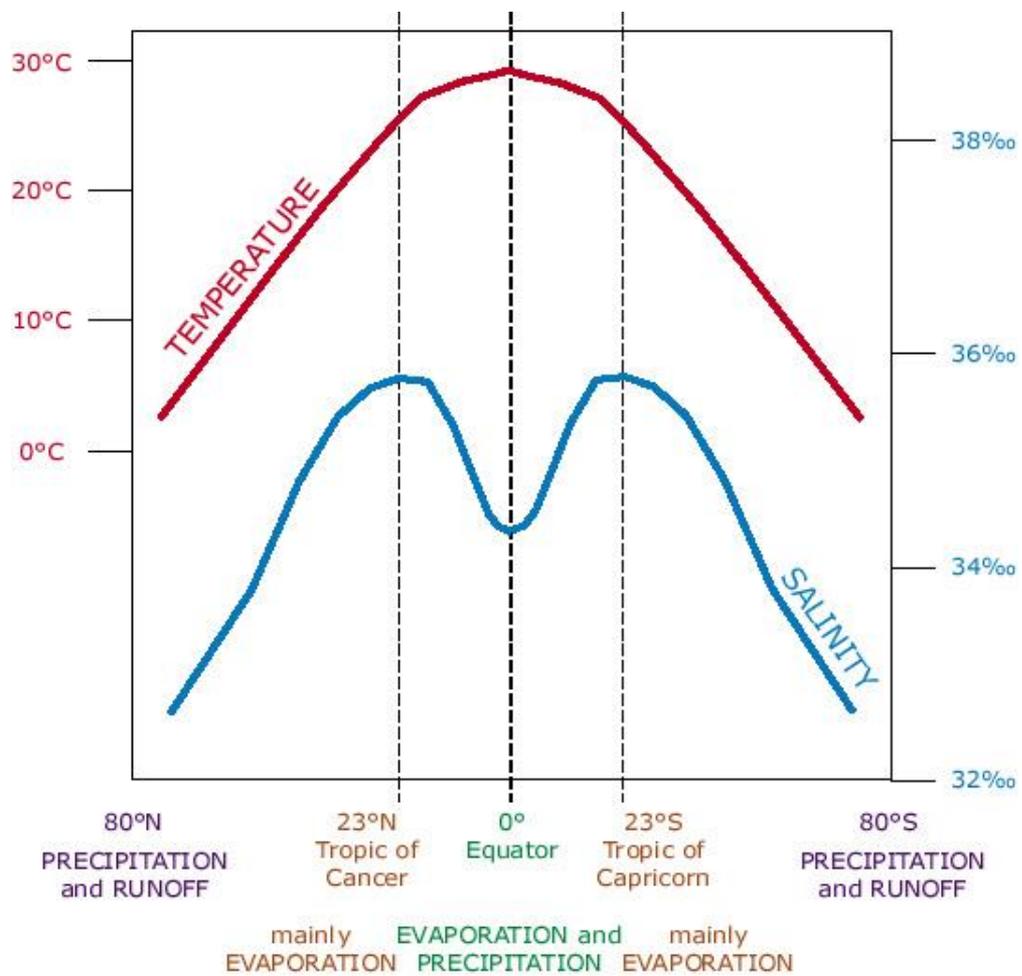
Oceano aberto (Temperatura) águas com $\uparrow T$, \downarrow densidade, são geralmente encontradas em camadas superficiais, e águas com $\downarrow T$, \uparrow densidade, em camadas profundas.

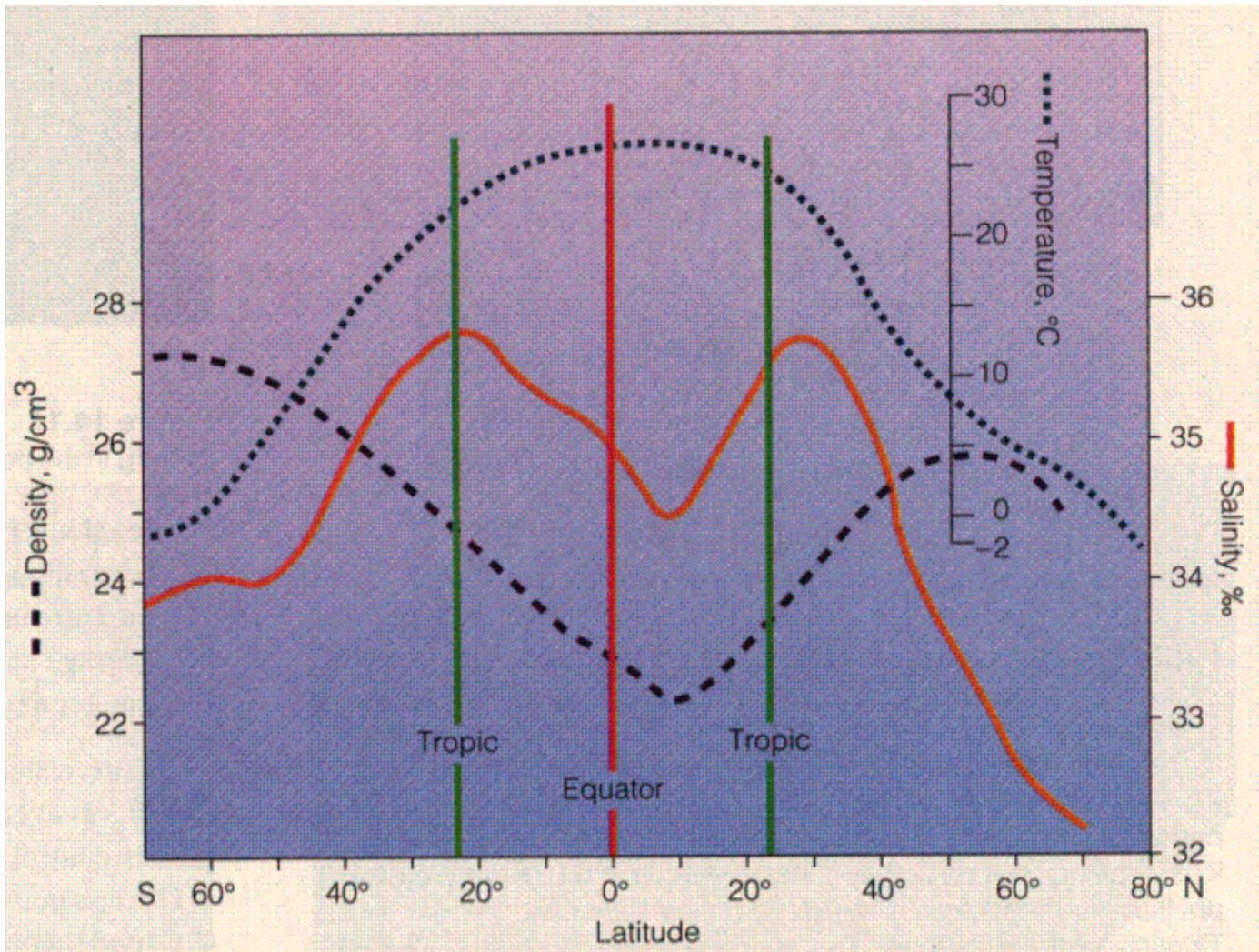
A Δ da S que ocorre em oceano aberto não tem efeitos suficientemente importantes para sobrepor-se ao efeito da temperatura na densidade.

Em águas costeiras (fjords, lagoas e estuários), a S é sempre o fator controlador na determinação da densidade, em todas as profundidades.



Increase in density with pressure for a water parcel of temperature 0°C and salinity 35.0 at the sea surface.



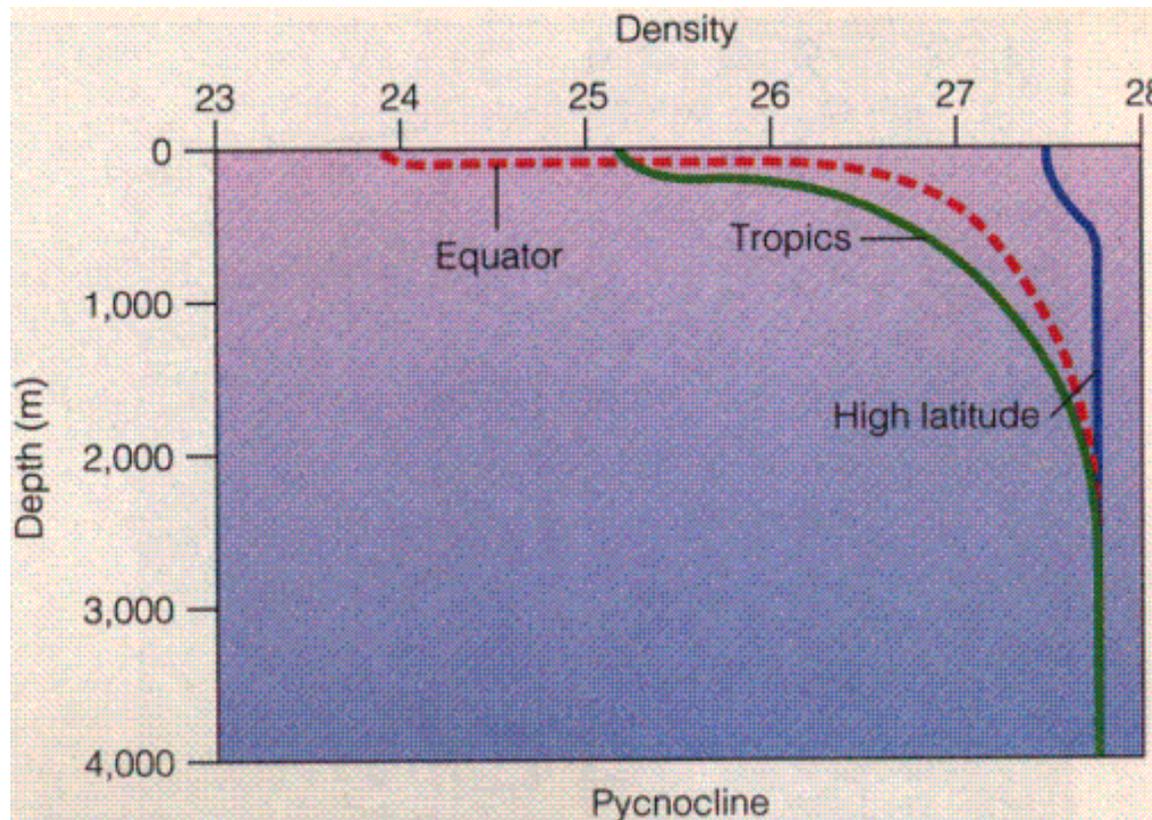


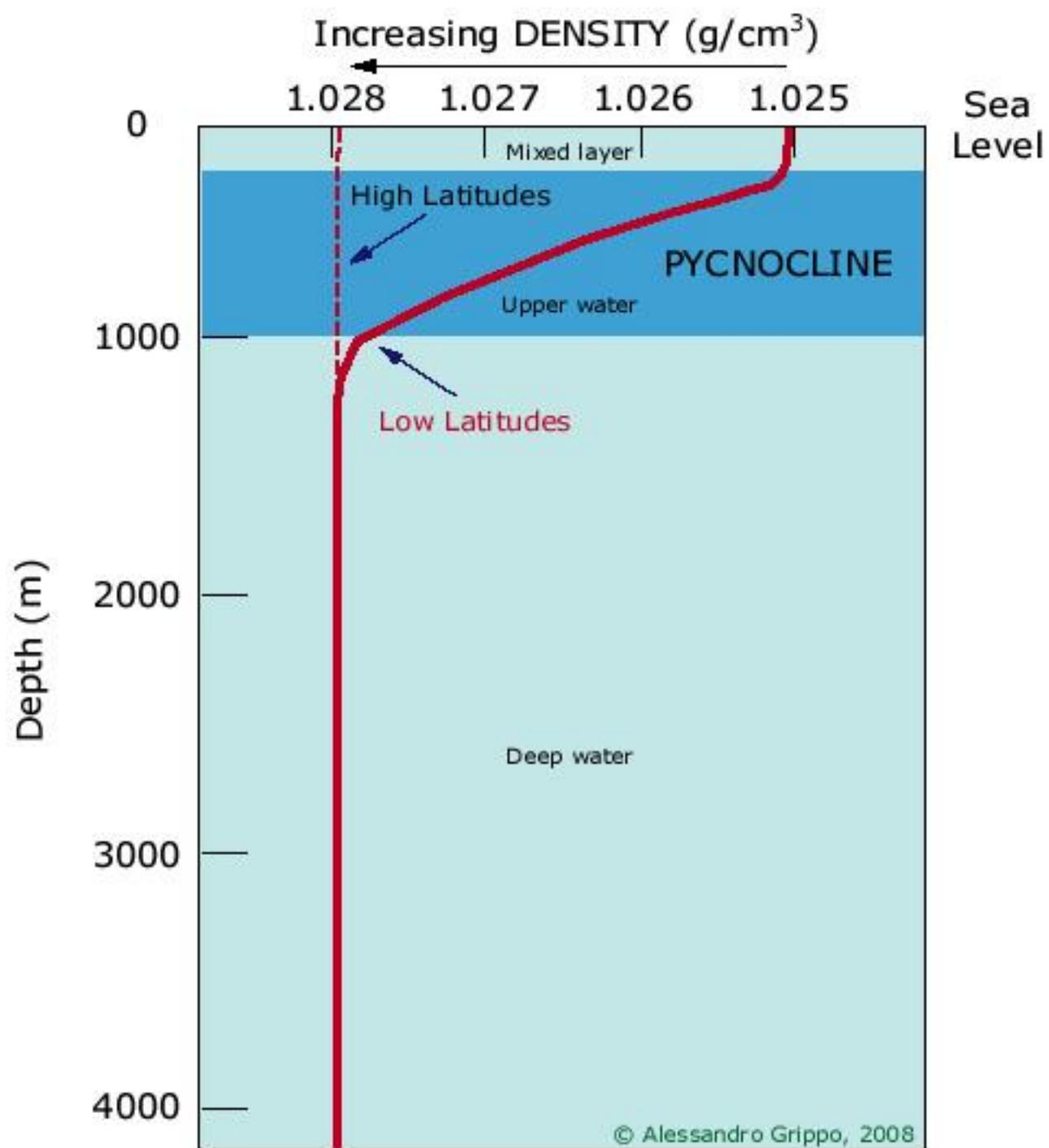
DISTRIBUIÇÃO VERTICAL DA DENSIDADE

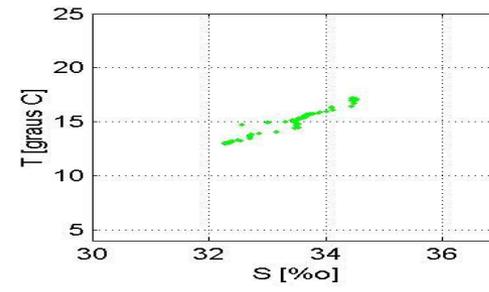
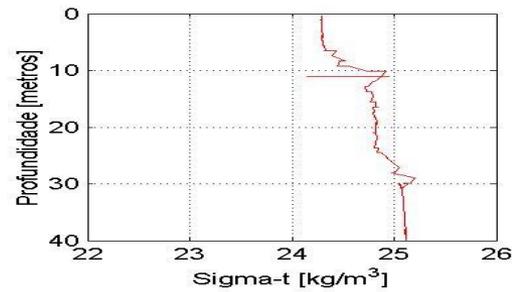
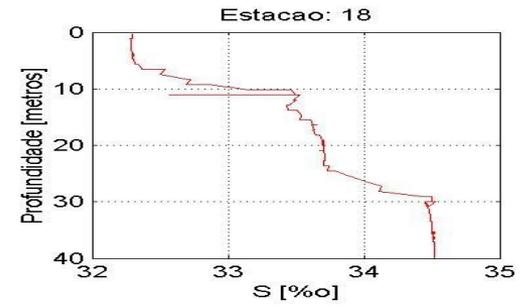
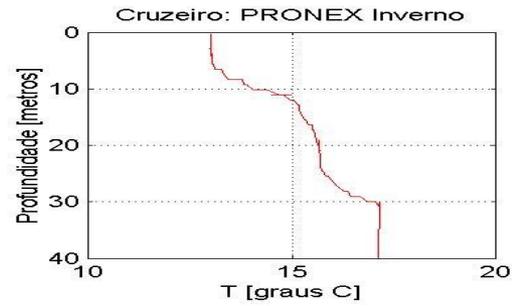
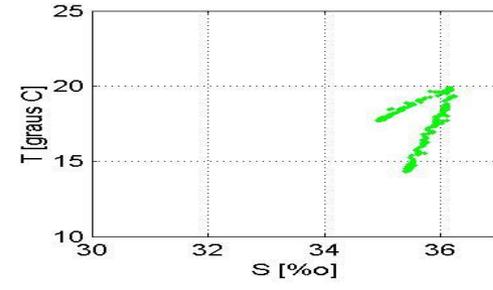
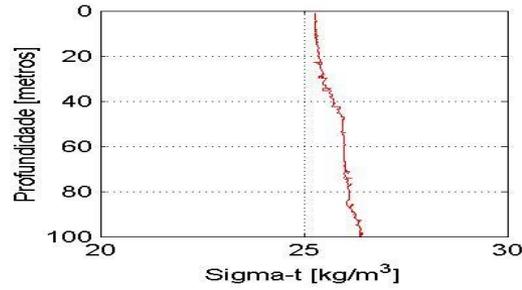
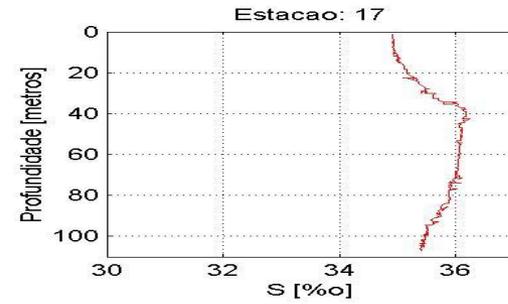
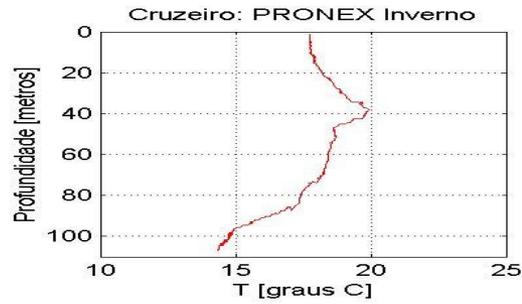
Equador: Pícnoclina saliente/ camada de mistura peq

Trópicos: Camada de mistura maior/pícnoclina pouco saliente

Pólos: Pícnoclina muito peq (apenas devido às variações de S)







Estabilidade/instabilidade na coluna de água

- Conceitos: estável, neutro, instável
- O número E – ondas internas – a frequência de Brünt-Väisälä
- Geração de instabilidade
- Processo de Dupla-difusão

A equação da estabilidade estática:

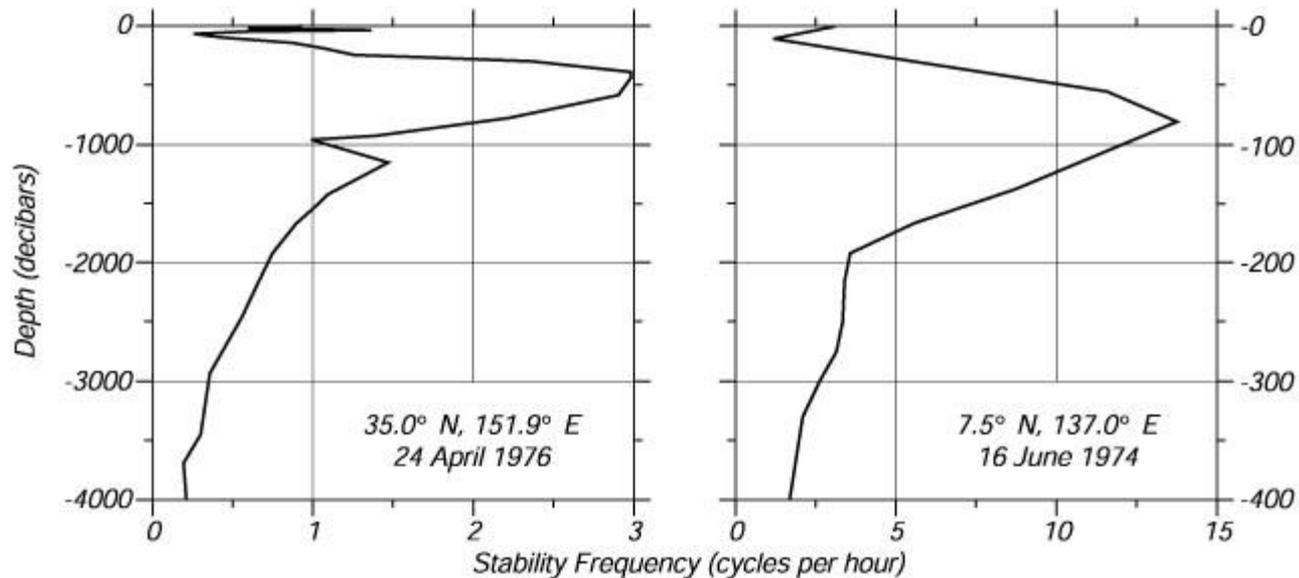
$$E \approx -\frac{1}{\rho} \frac{d\rho}{dz}$$

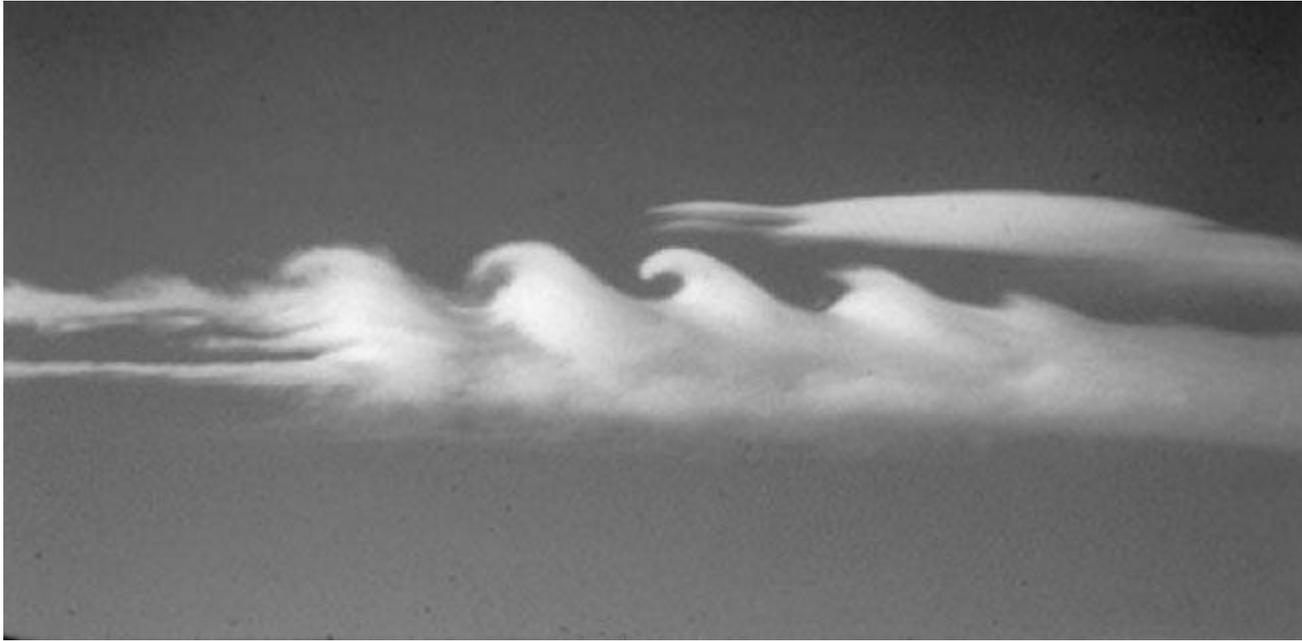
$E > 0$ Estável

$E = 0$ Neutro

$E < 0$ Instável

In the upper kilometer of the ocean, $z < 1,000\text{m}$, $E = (50-1000) \times 10^{-8}/\text{m}$, and in deep trenches where $z > 7,000\text{m}$, $E = 1 \times 10^{-8}/\text{m}$.

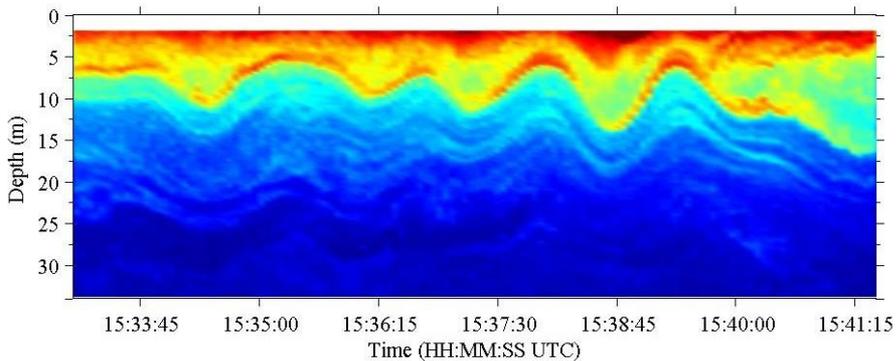




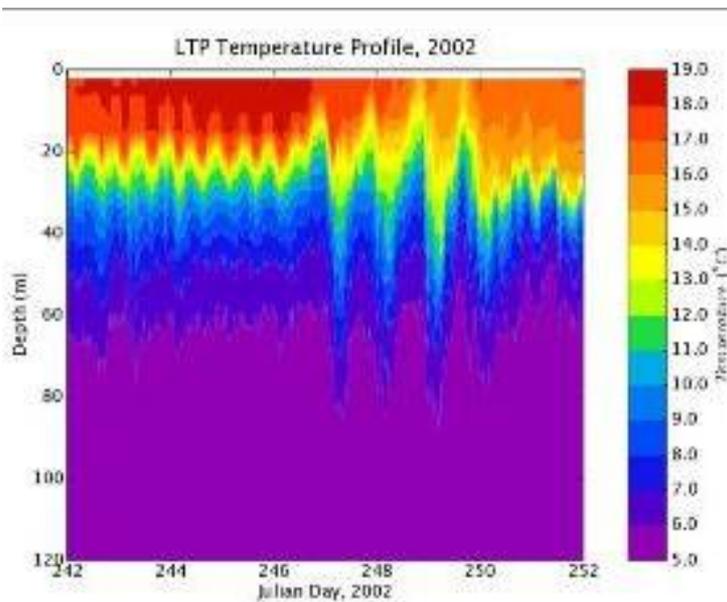
Billow clouds showing a Kelvin-Helmholtz instability at the top of a stable atmospheric boundary layer. Some billows can become large enough that more dense air overlies less dense air, and then the billows collapse into turbulence.

Photography copyright Brooks Martner, NOAA Environmental Technology Laboratory.





Ondas internas no Saint Lawrence Estuary: Large-amplitude internal waves (see image to the right) may be generated by stratified tidal flows over bottom topographic features. It is thought that such waves may be an important source of mixing in coastal waters, which means that they play a key role in coastal ecosystems, circulation, pollutant dispersal, etc.



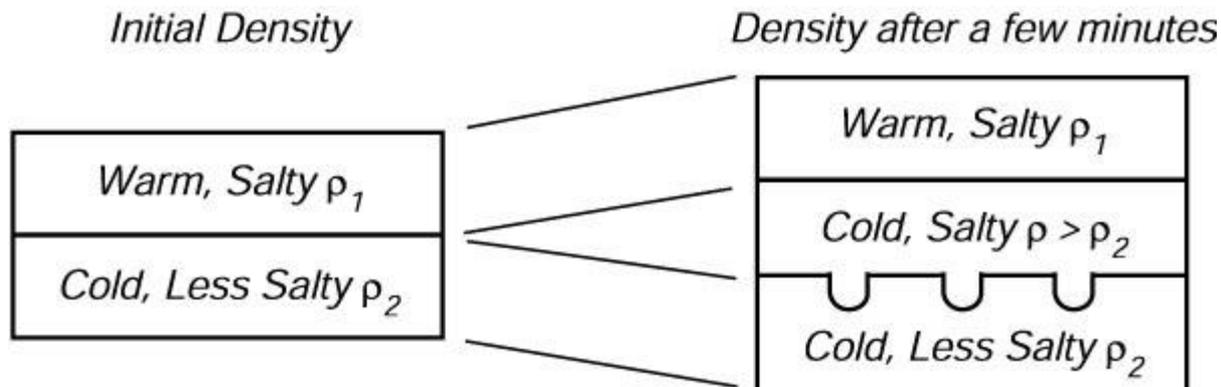
<http://iod.ucsd.edu/research/iwaves/>

<http://www.es.flinders.edu.au/~mattom/IntroOc/notes/figures/animations/fig10a7c.html>

Double Diffusion and Salt Fingers

In some regions of the ocean, less dense water overlies more dense water, yet the water column is unstable even if there are no currents. The instability occurs because the molecular diffusion of heat is about 100 times faster than the molecular diffusion of salt. The instability was first discovered by Melvin Stern in 1960 who quickly realized its importance in oceanography.

Consider two thin layers a few meters thick separated by a sharp interface. If the upper layer is warm and salty, and if the lower is colder and less salty than the upper layer, the interface becomes unstable even if the upper layer is less dense than the lower.



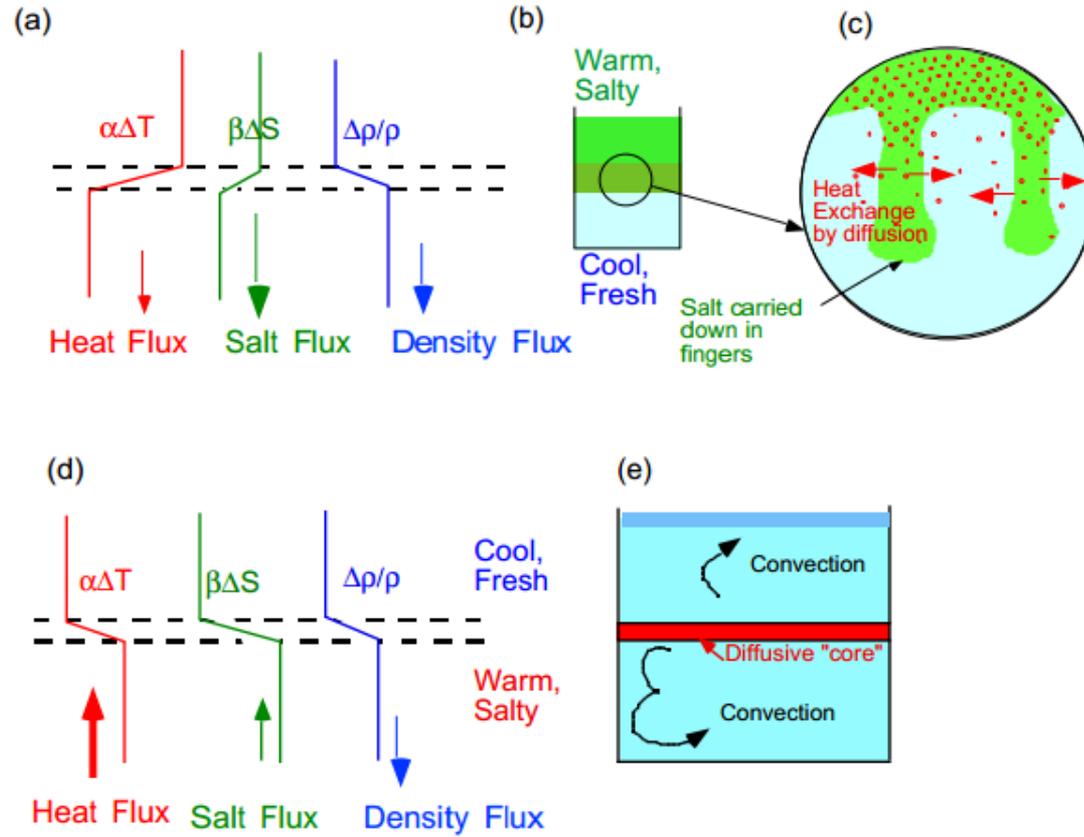
Warm salty over colder less salty. This process is called salt fingering. It occurs in central waters of sub-tropical gyres, western tropical North Atlantic, and the North-east Atlantic beneath the out flow from the Mediterranean Sea. Salt fingering eventually leads to density increasing with depth in a series of steps. Layers of constant-density are separated by thin layers with large changes in density, and the profile of density as a function of depth looks like stair steps. Schmitt *et al.*, (1987) observed 5-30m thick steps in the western, tropical North Atlantic that were coherent over 200-400km and that lasted for at least eight months. Kerr (2002) reports a recent experiment by Raymond Schmitt, James Leswell, John Toole, and Kurt Polzin showed salt fingering of Barbados mixed water 10 times faster than turbulence.

Colder less salty over warm salty. This process is called diffusive convection. It is much less common than salt fingering, and it is mostly found at high latitudes. Diffusive convection also leads to a stair step of density as a function of depth. Here's what happens in this case. Double diffusion leads to a thin, warm, less-salty layer at the base of the upper, colder, less-salty layer. The thin layer of water rises and mixes with water in the upper layer. A similar process occurs in the lower layer where a colder, salty layer forms at the interface. As a result of the convection in the upper and lower layers, the interface is sharpened; and any small gradients of density in either layer are reduced. Neal *et al.* (1969) observed 2-10m thick layers in the sea beneath the Arctic ice.

Cold salty over warmer less salty.

Warmer less salty over cold salty.

Salt Fingering e Cabelling



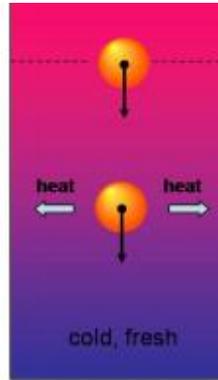


Figure 1: Schematic of the salt finger mechanism.

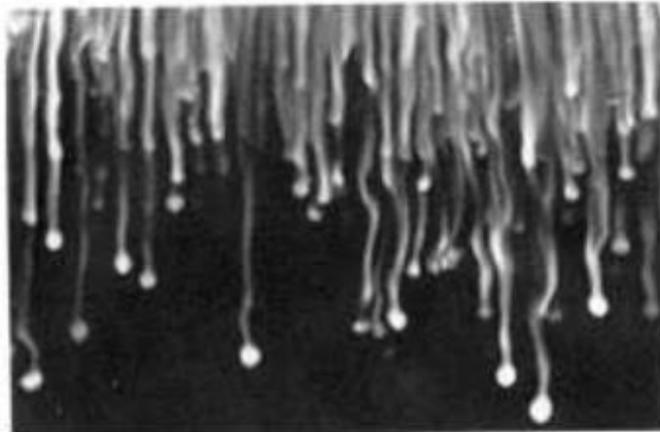


Figure 2: Salt fingers in a laboratory experiment [3].